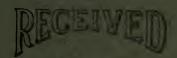
Changes in Stratigraphic Nomenclature by the U.S. Geological Survey 1967

GEOLOGICAL SURVEY BULLETIN 1274-A





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Changes in Stratigraphic Nomenclature by the U.S. Geological Survey 1967

By GEORGE V. COHEE, ROBERT G. BATES, and WILNA B. WRIGHT CONTRIBUTIONS TO STRATIGRAPHY

GEOLOGICAL SURVEY BULLETIN 1274-A



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CONTRIBUTIONS TO STRATIGRAPHY

CHANGES IN STRATIGRAPHIC NOMENCLATURE BY THE U.S. GEOLOGICAL SURVEY, 1967

By GEORGE V. COHEE, ROBERT G. BATES, and WILNA B. WRIGHT

LISTINGS OF NOMENCLATURAL CHANGES

In the following table, stratigraphic names adopted, revised, reinstated, or abandoned are listed alphabetically. The age of the unit, the revision, the area involved, the author's name, and date of publication of the report are given. The publications in which the changes in nomenclature were made are listed in the references at the end of this publication. The capitalization of age terms in the age column follows official usage.

A1

Abrigo Formation	294	Location	Revision and reference
	- Middle and Late Cambrian southeastern Arizona.	southeastern Arizona	Abrigo Formation in report area subdivided (in ascending order) into the Three C Southoun
Achiote Conglomerate	. Late Cretaceous-	west-central Puerto Rico.	Belle, and Peppersuee Members. (Creasey, 1997b.) New name adorted
Adams Argillite	- Early Cambrian	east-central Alaska	(Mattson, 1967.) New name adouted
Admire Group	. Early Permian	northeastern Kansas.	(Brabb, 1967.) (Brabb, 1967.) Age channed from Parmian to Fouly Domeion
Alaska Bench Limestone (of Amsden Group) Early and Middle Pennsyl- western Montana. vanian.	. Early and Middle Pennsyl- vanian.	western Montana.	(Johnson and Warner, 1967.) Alaska Bench Limestone removed from Big Snowy Group and reassigned as middle formation of
Aleman Formation (of Montoya Group)	Late Ordovician	southwestern New Mexico	Arussen tyroup. Age changed from Mississippian or Pennsylvanian to Early and Middle Pennsyl- (Maughan and Roberts, 1967.) Formerly Aleman Cherty Member of Montoya Dolomite and remains a member of this forma-
Allsbury Formation	Early Silurian	northeastern Maine	ton outside the report area. (Pratt, 1967.) New name adopted.
Alta Formation	Early Permian (Wolfcamp) west Texas	west Texas.	(Ekren and Frischknecht, 1967.) Age changed from Permian to Forly Pormion
American Flag Formation	Cretaçeous(?)	south-central Arizona_	(Orie), Myers, and Crosby, 1967.)
Americus Limestone Member (of Foraker Lime- stone)	Early Permian	northeastern Kansas.	(Creasey, 1967b.) Age changed from Permian to Farly Permian
Amsden Group	Early and Middle Pennsyl- central Montana vanian.	central Montana	(Johnson and Wagner, 1967.) Amsden raised to group rank in central Montana where it includes (in ascending order) the Tyler Formation Alseke Baroh Linosettes and horder
Andrecito Member (of Lake Valley Limestone) Early Mississippian (Osage) southwestern New Mexico Anón Formation west-central Puerto Rico Antelope Valley Formation (of Pogonip Group) Early and Middle Ordovician. southeastern California	Early Mississippian (Osage) middle to late Eocene Early and Middle Ordovician.	southwestern New Mexico west-central Puerto Rico	 Potential Andreas Detail Intervoluts, and Devils in good usage elsewhere. (Maughan and Roberts, 1967.) Andrectio Member of Laudon and Bowsher (1949) douber of Laudon and Bowsher (1949) (Jones, Hernon, and Moore, 1967.) Anon Formation of Pessagno (1960) adopted. (Matison, 1967.) Formation in part of Inyo County, Calif. Remains

4 Domnotion	Early Permian	Nevada	Age changed form Permian to Early Permian.
Arcturus Follitation.			(Ketner, 1967.)
Ashlock Formation	. Late Urdovician (Uincinnat- ian).	central Acentucky	Owingsville, the Gilbert and Stingy Creek
			Members of the Ashlock Formation are replaced by the Grant Lake Member.
		•	(Simmons, 1967b.)
Aspinwall Limestone Member (of Onaga Shale) - Early Permian-	. Early Permian	northeastern Kansas	Age changed from Fermian to Early Fermian. (Tobuson and Warner 1967)
Adin Mombor (of Mormonskill Shala) Middle Ordovician	Middle Ordovician	eastern New York	Austin Glen Member of Ruedemann (1942) adopted.
AUSTIN CIERT METHON (01 10 100 100 100 100 100 100 100 100		anthouten Vareas	(Zen, 1967.) A roos T imestone Member of Condra (1997) adonted
Avoca Limestone Member (of Lecompton Lime-	Late Pennsylvanian (Virgil) horuneastern mansas-	INTURBASUATI IN ALISAS	(Johnson and Adkison, 1967.)
Stone) Reken Formation	Late Devonian and Early	Montana and North Dakota	Age changed from Late Devonian(?) and Early
	Mississippian.		Mississippian to Late Devonian and Early Mississippian.
	()limma(?)	A laska	(Sandberg and Klapper, 1967.) Age changed from Oligocene or Miocene to Oligo-
Banjo Foint Formadoll.	- Oligocomo(i)		cene(?).
Baners T'nff Member (of Quichapa Formation)	- Oligocene or Miocene	southwestern Utah	Bauers Tuff Member of Mackin (1960) adopted.
D of Dillone Teams tion	L'ata Silinrian	southeastern Alaska	(Averitt, 1967.) New name adopted.
Day of Fillars Future under the second secon			(Muffler, 1967.) Are absured from Pelsonatio(?) to lata Pelsonatio(?)
Bean Canyon Formation	- late Paleozoic(?)	southeastern California	Age changed itout 1 aleccore(1) to late 1 aleccore(1). (Dibblee, 1967a.)
Bedrock Snring Formation	- middle(?) Pliocene	do	Age changed from middle Pliocene to middle(?)
			Pliocene. (Dibblee, 1967a.)
Beebe Limestone Member (of West Castleton Formation)	Early Cambrian	. western Vermont	Reduced in rank from Beebe Limestone to Beebe Limestone Member.
runnauou). D. r. riesetee Momber /of Loomenton I imo. Lata Danneelvanian (Virgil) - northeastarn Kansas.	I.ata Bannswiwanian (Viroil)	northeastern Kansas	(Zen, 1967.) Beil Limestone Member of Condra (1930) adopted.
Bell Limestone Member (of Lecondpion Line- stone).		southastan Idaho	(Johnson and Adkison, 1967.) Included in Snake River Groun.
Big Hole Basalt	LIeistocene	- 200 Mileaster Trans	(Mundorff, 1967.)
Big Snowy Group-	- Late Mississippian	Montana and North Dakota	Big Snowy Group restricted to (in ascending order) the Withew Sendstone Otter and Heath
			Formations, of Late Mississippian (Chester) age.
	I it. Donorionication (Ilianil)	northoastarn Kansas	(Maughan and Koberts, 1967). Rig Shrings Limestone Member of Condra (1927)
Big Springs Limestone Member (of Lecompton Late reunsylvanian (Yugu) not pressored Automotion Limestone).	Late Feimsylvaman (Vugu)		adopted. (Tobreen and Advison 1967)
Bissell Formation (of Tropico Group)	- Miocene(?)	southeastern California	Age changed from Miocene or Pliocene to Miocene(?)
Black Mountain Basalt	. Pleistocene	do	Black Mountain Basalt of Baker (1912) adopted.
			(DIDDIER, 190/a.)

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Location Revision and reference	western Vermont Bomoseen Graywacke Member of Ruedemann	- Pennsylvania, New Jersey and Age changed from Silurian to Late Silurian (Ca- New York).	 Southwastern Idaho, northwestern New name adopted. Southwastern New Wryconling, and Jennings, 1967.) Southwastern New Warion. Roy Mamber of Sciences (1997) 	1°ă	 weak out of the control of the control	- Alaska	- southeastern Alaska New name adomted			. routhwestern New Mexico Age changed from Late Ordovician to Middle	Ortoovician. (Jones: Herron and Moore, 1967.) Formerly Cable Carryon Sandstone Member of	Montoya Dolomite, which remains good usage elsewhere. (Pratt, 1967.) western Montana	SISSIPPIAN OF Pennsylvanian to Early Pennsyl- vanian (Morrow)
Åge	Cambrian(?)	Late Silurian	Middle(?) Jurassic Late Devonian	Tertiary	Cambrian(?) and Early Cam- western Vermont. brian.	late(?) Oligocene		Late Pennsylvanian (Virgil)	Middle and Late Ordovician.		do	Early Pennsylvanian (Morrow).	
Name	Bomoseen Graywacke Member (of Bull Forma- Cambrian(?)	Bossardville Limestone	Boundary Ridge Member (of Twin Creek Lime- Middle(?) Jurassic. stone). Box Member (of Percha Shale) Late Devonian	Bozeman Group	Bull Formation.	Burls Creek Shale Member (of Katalla Forma- tion).	Burnt Island Conglomerate (of Hyd Group) Late Triassic	Burroak Shale Member (of Deer Creek Lime- stone).	Bushkill Member (of Martinsburg Formation) Middle and Late Ordovician Pennsylvania and New Jersey.	Cable Canyon Sandstone Member (of Montoya Middle Ordovician. Dolomite).	Cable Canyon Sandstone Member (of Seconddo	Cameron Creek Member (of Tyler Formation) Early Pennsylvanian (Morrow).	

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Age changed from Early or Middle Devonian to Early Devonian.	(Oliver, 1967.) In report area divided into (ascending order) limestone member, handed member, gypsiferous member, and Winsor Member. (Coshion (1970))	Carney Lake Greiss of Treves (1960) adopted.	(Pagyey, Dutton, and Lamey, 1900.) New name adopted. (Plafter, 1967.)	New name adopted.	Age changed from Early Mississippian to Late Mississippian.	(Maughan and Roberts, 1967.) Name abandoned, 1967.) (1150-1950-1950-1957.)	China formation divided into (ascending order) Chinitra Formation divided into (ascending order) Tonnie Siltstone and Paveloff Siltstone Members.	(Detterman and Hartsock, 1966.) Age changed from Permian to Early Permian (Wolfcamp and Leonard).	Age changed from Pennsylvanian and Permian to Middle and Late Pennsylvanian.	CUTE, MYELS, and CLOSDY, 1907) Clay Creek Limestone Member of Moore (1932) adopted.	(Johnson and Adkison, 1967.) Age changed from Late Cretaceous(?) and Tertiary to Late(?) Cretaceous or early Tertiary.	(Creasey, 1967b.) New name adopted.	(Epstein, Epstein, Spink, and Jemmiss, 1907) Coachella Fanglomerate of Vaughan (1922) adopted. (Dishibar 1937)	In report area, name changed from Coeymans Linestone teo, Coeymans Formation; includes	(ascending order) Depue Limestone, Peters Valley, Shawnee Island, and Stormville Members. Coeymans Limestone remains in good usage else-	where. (Bystein, Epstein, Spink, and Jennings, 1967.) (Bystein, Epstein, Spink, and Jennings, 1967.) (Costein, Epstein, Spink, and Jennings, 1967.)
east-central New York	southwestern Utah	Upper Peninsula, Michigan	south-central Alaska	northwestern Alaska		southwestern Massachusetts	southwestern Alaska	west Texas	west Texas	northeastern Kansas	south-central Arizona	New Jersey	southeastern California	New Jersey and Pennsylvania		eastern New York
- Early Devonian	. Middle and Late Jurassic	- early Precambrian	- post-early Oligocene(?) to pre-middle Mincene	Pleistocene	. Late Mississippian (Meramec). western Montana.	Late Triassic	Late Jurassic	- Early Permian (Wolfcamp and Leonard),	- Middle and Late Pennsyl- vanian.	Late Pennsylvanian (Virgil) northeastern Kansas .	- Late(?) Cretaceous or early Tertiary.	- Late Silurian	- Miocene(?)	- Early Devonian		. Early Devonian
Carlisle Center Formation	carmel Formation	Carney Lake Gneiss	Cenotaph Volcanics	Chariot Gravel	Charles Formation (of Madison Group)	Chicopee Shale	Chinitna Formation	Cibolo Formation.	Cieneguita Formation	Clay Creek Limestone Member (of Kanwaka shalo)	Cloudburst Formation	Clove Brook Member (of Decker Formation)	Coachella Fanglomerate	Coeymans Formation (of Helderberg Group)		Coeymans Limestone (of Helderberg Group)

A5

Name	Age	Location	Revision and reference
Corbin Sandstone Member (of Lee Formation) Early Pennsylvanian.	. Early Pennsylvanian	southeastern Kentucky	Corbin Sandstone Tongue adopted for Wofford, Hollyhill, and Williamsburg quadrangles. Corbin Sandstone Member remains in good usage else- whene.
Cornwallis Limestone (of Hyd Group)	. Late Triassic	southeastern Alaska	(Smith, 1967.) New name adopted.
Cottonwood Canyon Member (of Madison Limestone). Cottonwood Canyon Member (of Lodgepole	Late Devonian and Early Mississippian. Early Mississippian	northern and west-central Wyo- ming. southern Montana	(Muffler, 1967.) New name adopted. (Sandberg and Kilapper, 1967.) New name adonted
Limestone of Madison Group). Council Grove Group		northeastern Kansas.	(Sandberg and Klapper, 1967.) Age changed from Permian to Early Permian.
Coyote Butte Formation	Early Permian (Wolfcamp and Leonard).	Oregon	(Johnson and Wagner, 1967.) Age changed from Permian to Early Permian (Wolfcamp and Leonard).
Crab Orchard Formation	Early and Middle Silurian central Kentucky	central Kentucky	(Ketner, 1967.) Crab Orchard Formation of report area includes
			(ascending order) Lower Silurian Flum Creek Member, and Middle Silurian Oldham, Lulbe- grud Shale, Waco, and Estill Shale Members.
Crowder Formation	Pliocene.	southeastern California	(Simmons, 1967a.) New name adopted.
Cuba Mesa Member (of San Jose Formation) early Eocene	early Eocene	northwestern New Mexico	(Dibblee, 1967a.) New name adopted.
Curzon Limestone Member (of Topeka Lime- Late Pennsylvanian (Virgil) northeastern Kansas.	Late Pennsylvanian (Virgil).	northeastern Kansas.	(Baltz, 1967.) Curzon Limestone Member of Gallaher (1898)
cutter Dolomite (of Montoya Group)	Late Ordovician	southwestern New Mexico	adopted. (Johnson and Adkison, 1967.) Cutter Dolomite used in report area. Cutter Dolo- mite Member or Cutter Member in rood usare
Danby Formation	Late(?) Cambrian	western Vermont	elsewhere. (Pratt, 1967.) Age changed from Early Cambrian to Late(?)
Daylight Formation	Precambrian and Early Cam- brian.	southern California and southern Nevada.	Cam. 1967. (Zen., 1967.) Age changed from Early Cambrian to Precambrian and Early Cambrian.
De Chelly Sandstone	Early Permian (Leonard)	Colorado, Arizona, and New Mexico.	(Barnes and Christiansen, 1967.) Age changed from Permian to Early Permian (Leonard).
Decker Formation	Late Silurian	New Jersey, New York, and Penn- sylvania.	(Hallgarth, 1967.) Formerly Decker Limestone. Includes Wallpack Center and Clove Brook Members. (Epstein, Epstein, Spink, and Jennings, 1967.)

CHANGES IN STRATIGRAPHIC NOMENCLATURE-Continued

Dedham Granodiorite	pre-Carboniferous	eastern Massachusetts	Age changed from early Polooroio(2) to nue Control
Deer Creek Limestone (of Shawnee Group) Late Pennsylvanjan (Virgil)northeastern Kansas	Late Pennsvlvanian (Virgil)	northeastern Kansas	Church Target and the carry 1 are concreted by pre-Carlour- (Church, 1966.) Dear Creek Timestone in area of record indian
			These members (assenting order): Ozavkie Line- stone, Oskaloosa Shale, Rock Bluff Linestone, Larsh Shale, Burroak Shale, Ervine Creek Line-
Depue Limestone Member (of Coeymans For-	Early Devonian	Pennsylvania and New Jersey	stonte. (Johnson and Adkison, 1967.) New name adopted.
Devils Pocket Formation (of Amsden Group)	. Middle Pennsylvanian	western Montana	(Epstein, Epstein, Spink, and Jennings, 1967.) Formerly, of Big Snowy Group. Age changed
			from Mississippian or Pennsylvanian to Middle Pennsylvanian (Aroba and Des Moines(?)).
Diablo Formation	Early and Late Permian (Guadalupe).	Nevada	Age changed from Permian to Early and Late Permian (Guadaluna)
Dimple Limestone	Middle Pennsylvanian	west Texas	(Ketner, 1967.) Age changed from Pennsylvanian to Middle Penn-
Dinosaur Canyon Member (of Moenave Forma- Late Triassic(?) tion).	Late Triassic(?)	southwestern Utah	Oriel, Myers, and Crosby, 1967.) [Oriel, Myers, and Crosby, 1967.] Formerly Dinosaur Canyon Sandstone Member and remains as such outside of aroo of record
Doniphan Shale Member (of Lecompton Lime- Late Pennsylvanian (Virgil) northeastern Kansas.	Late Pennsylvanian (Virgil)	northeastern Kansas	(Wilson and Stewart, 1967.) Doniphan Shale Member of Condra (1927) adonted.
stone). Double Bluff Drift	Pleistocene	northwestern Washington	(Johnson and Adkison, 1967.) New name adopted.
Douglas Group.	Pennsylvanian (Virgil)	northeastern Kansas	(Easterbrook, Crandell, and Leopold, 1967.) In Kansas base of Douglas Group lowered to top
Douglass Mesa Gravel	Pleistocene (Kansan or	central Colorado	of Lansing Group. Johnson and Addrison, 1967.) New name adored
Duttonville Member (of Rondout Formation)	Yarmouth). Late Silurian or Early	New Jersey and Pennsylvania	(Yarnes and Scott, 1967.) New name adopted.
El Paso Dolomite.	Devonian. Early Ordovician	southwestern New Mexico	(Epstein, Epstein, Spink, and Jennings, 1967.) Formerly El Paso Limestone or Formation in
Enitanh Dolomite (of Naco Groun)	Rarly Darmian	A virono	report area. Both terms remain in good usage elsewhere. (Pratt, 1967.)
Enitanh Formation	do	Artizulia	Age changed from Permian to Early Permian. (McKee, 1967.)
Frutha Croath Linnation Manhar at A			Eputapli Formation in report area; Epitaph Dolo- mite in good usage elsewhere. (Creasey, 1967a.)
Creek Linestone).	Late Fennsylvanian (virgil)	northeastern Kansas	Ervine Creek Limestone Member of Condra (1927) adopted.
			Jonnson and Adkison, 1967.)

-Continued	Revision and reference	. Fairview Valley Formation of Bowen (1954) adopted. (1)hblee. 1967a.)	Ea.	. Age changed from Fermian to Early Fermian. (Johnson and Wagner, 1967.) New name adonted.		Age changed from Pennsylvanian and Permian to Middle and Late Pennsylvanian and Barly Dermide	Fraction Tuff in 1967.) Fraction Tuff in report area; Fraction Breecia in	good usage tasewhere. (Ekren, Rogers, Anderson, and Botinelly, 1967.) New name adopted.	(Brabb, 1967.) Age changed from middle Tertiary to middle	t et darsey, 1967b.) (Creasey, 1967b.) . Age charged from Pennsylvanian to Middle and Late Pennsylvania and Barly Permian.	Garlock Series of Dibblee (1952) adopted as Garlock Formation.	(Ulibblee, 1967a.) Age changed from Miocene(?) to Oligocene(?) to middle Miocene(?).	New name adopted. (Imlay, 1967.)	Goler Fromation of Lynoptee (1922) adopted. (Dibblee, 1962a). - Age changed from Cambrian(?) to Early Cam-	Drian(7): 977a.) (Neuman, 1967a.) (Neuman, 1967a.) Grant Lake Member used in area from Richmond Grant Lake Limenorpassivarid to Owingsville. Grant Lake Limenorpassivarid to Owingsville northward to Maysville. (Simmons, 1967b.)
STRATIGRAPHIC NOMENCLATURE-Continued	Location	California	northeastern Kansas	northeastern Kansas New Jersev, New York, and Penn-	sylvania. northeastern Kansas	Wyoming, Montana, North Dako- ta, and South Dakota.	southern Nevada	east-central Alaska	south-central Arizona.	west Texas	southeastern California	do	Wyoming, Idaho, and Utah	southeastern California	eastern Kentucky
IN STRATIGRAPH	Age	Permian and Permian(?)		Early Permian 1 Farly Devonian	1	Middle and Late Pennsylva- nian and Early Permian.	late Miocene	Early Cambrian	middle Tertiary or younger south-central Arizona	Middle and Late Pennsyl- vanian and Early Permian.	Permian and Permian(?)	Oligocene(?) to middle Mi- ocene(?).		Paleocene and Eocene s Early Cambrian(?)	Late Ordovician (Cincinnatian).
CHANGES IN	Name	Fairview Valley Formation.	Falls City Limestone (of Admire Group)	Five Point Limestone Member (of Janesville Shale). Distribution Mombor (of New Southand For-	Figure 1 and	Fountain Formation	Fraction Tuff.	Funnel Creek Limestone	Galiuro Volcanics.	Gaptank Formation	Garlock Formation	Gem Hill Formation (of Tropico Group)	Giraffe Creek Member (of Twin Creek Lime- Late Jurassic. stone).	Goler Formation	Grant Lake Member (of Ashlock Formation) Late Ordovician (Cincinnatian). eastern Kentucky

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CONTRIBUTIONS TO STRATIGRAPHY

 Age changed from Oligocene to late Oligocene and (o) early Miocene. (Monroe. 1967.) 	Age changed from Oligocene or early Miocene to Oligocene(?).	(MacNeil, 1967.) New mane adopted.	(Mumer, 1907.) New name adopted.	(Muffler, 1967.) Age changed from Permian to Early Permian.	Harmony Hills Tuff Member of Mackin (1960)	acoptor 1967.) (Averitt, 1967.) Hartford Limestone Member of Moore, Frye, and Jewet (1944) relistated. Original usage, Hartford Limestone, abandoned in 1912.	(Johnson and Adkison, 1967.) Formery a member of Stranger Formation; now hasal member of revised Lawrence Formation.	(Johnson and Adkison, 1967.) Age changed from Late Cambrian to Middle and Late Cambrian.	(Prinz, 1967.) Hatch Hill Formation of Theokritoff (1959) adopted. (Zen, 1967.)	Age changed from Ordovician to Middle Ordovi-	(Chidester, Hatch, Osberg, Norton, and Hartshorn, 1967.)	Age changed from Permian to Early Permian. (Johnson and Wagner, 1967.)	Age changed from Pennsylvanian to Middle Penn- svlvanian.	(Oriel, Myers, and Crosby, 1967.) Minishik Limestone included in Helderberg Group	(Epstein, Epstein, Spink, and Jennings, 1967.) Age changed from Permian to Early Permian	(Leonard). (McKee, 1967.) New mane adopted. (Ryshin 1967).	Age changed from Mesozoic or older to Mesozoic and Permian (Ochoa?).	 (Ketner, 1967) Hole-In-The-Wall Tuff Member of Mackin (1960) adopted. (Averitt, 1967.)
northwestern Puerto Rico	Alaska	southeastern Alaska	do	northeastern Kansas	southwestern Utah	northeastern Kansas	do	western Montana	New York, Vermont	northwestern Massachusetts		northeastern Kansas	west Texas	New Jersey and Pennsylvania	Utah and Arizona	east-central Alaska	California	southwestern Utah
late Oligocene and (or) early northwestern Puerto Rico- Miocene.	Oligocene(?)	Early Permian.	Late Triassic	Early Permian	Oligocene or Miocene	Late Pennsylvanian (Virgil)	Pennsylvanian (Virgil)	Middle and Late Cambrian	Middle(?) and Upper Cambrian	vician		Early Permian	Middle Pennsylvanian	Early Devonian	Early Permian (Leonard)	Early Cambrian to Early	Mesozoic and Permian (Ochoa?).	Oligocene or early Eocene
Guajataca Member (of Cibao Formation)	Gunners Cove Formation	Halleck Formation	Homitton Island Limestane (of Hvd Groun)	Hamlin Shale Member (of Janesville Shale)	Harmony Hills Tuff Member (of Quichapa	Formation). Hartford Limestone Member (of Topeka Lime- Late Pennsylvanian (Virgil) northeastern Kansas stone).	Haskell Limestone Member (of Lawrence Forma- Pennsylvanian (Virgil)	tion). Hasmark Formation	Hatch Hill Formation	Hawley Schist.		Hawxby Shale Member (of Onaga Shale)	Haymond Formation	Helderberg Group.	Hermit Shale (of Aubrey Group)	Hillard Limestone	Hodge Volcanic Formation	Hole-In-The-Wall Tuff Member (of Isom Forma- tion).

A9

Name	Age	Location	Revision and reference
Hoosac Formation	- Early(?) Cambrian	- western Vermont	Hoosac Formation as redefined by Skehan (1961) adopted in Vernut; Hoosac Schist remains in
Horn Mountain Tuff Member (of Talkeetna Early Jurassic.	Early Jurassic	southwestern Alaska	good usage in Massachuseus and Connecticut. (Zen, 1967.) New name adorted.
Formation). Horned Toad Formation	- middle Pliocene.	southeastern California	(Detterman and Hartsock, 1966.) Age changed from early or middle Pliocene to
Hoskin Lake Granite	- middle Precambrian	Michigan and Wisconsin	(Dibblee, 1967a.) (Dibblee, 1967a.) Age changed from Precambrian to middle Pre-
Houchen Creek Limestone Bed (of Hamlin Shala Membar)	Early Permian	- northeastern Kansas	(Dutton and Linebaugh, 1967.) Age changed from Permian, to Early Permian.
Hound Island Volcanics (of Hyd Group)	- Late Triassic	southeastern Alaska	(Jonnson and Wagner, 1967.) New name adopted.
Hughes Creek Shale Member (of Foraker Lime-	Early Permian	northeastern Kansas	(Muffler, 1967.) Age changed from Permian to Early Permian.
stone). Hunton Formation	Silurian and Devonian	do	(Johnson and Wagner, 1967.) Hunton Formation used in area of report: Hunton
			Limestone or Hunton Group in good usage elsewhere.
Husted Alluvium	Recent	central Colorado	(Johnson and Adkison, 1967.) - New name adopted.
Hyd Formation.	Late Triassic	southeastern Alaska	(Varnes and Scott, 1967.) Raised to groun rank in renort area: includes
			(ascending order) Burnt Island Conglomerate, Cornwallis Limestone, Hamilton Island Lime- stone, and Hound Island Volcanics Hyd Forma-
latan Limestone Member (of Stranger Forma- Pennsylvanian (Vireil)	Pennsvlvanjan (Virøil)	northeastern Konses	tion remains in good usage elsewhere. (Muffer, 1967.)
tion).			- Taken Intrevoite Fourced to Internoer Tain, and Jacod in revised Stranger Formation. (Johnson and Adkison, 1967.)
	Ternary or Quaternary		- New name adopted. (Campbell, 1967.)
Interlake Formation Towa Point Shale Member (of Toneka Lime-	. Late Ordovician and Silurian. Late Pennsylvanian (Viroil)	North Dakota, South Dakota, Wyoming, and Montana.	Interlåke Formation of Baillie (1951) adopted. (Sandberg, 1967) Form Dobief Short, Marchane Conduct (1964)
stone). Ira Formation	Middle Ordovician		- Jowar out strate memory of Contura (1927) auopueu. (Johnson and Adkison, 1967.) Formerly Ira Slate of Rarly Ordovician age
Isom Formation.	Eocene or early Oligocene	southwestern Utah	(Zen, 1967.) Isom Formation of Mackin (1960) adopted.
			(Blank and Mackin, 1967.)

CHANGES IN STRATIGRAPHIC NOMENCLATURE-Continued

. New name adopted.	(Dibblee, 1967a.) Age changed from Permian to Early Permian.	(Jonnson and Wagner, 1967.) New name adopted.	Age changed from Permian to Early Permian.	Jones Point Shale Member of Condra (1927) adopted.	(Johnson and Adkison, 1967.) New name adopted.	(BrabD, 1967.) Kanwaka Shale divided into (ascending order) Lav Creek Limestone and Stull Shale Members	(Johnson and Adkison, 1967.) Katahlin Quartz Monzonite of Hitchcock (1861)	(Neuman, 1967a.) New name adopted.	(Mumer, 1967.) New name adopted.	(Varnes and Scott, 1967.) King Hill Shale Member of Condra (1927) adopted.	Age changed from Oligocene to Oligocene and	Miocene(?). (Muessig, 1967.) Formerly Keneeling Nun Rhyolite Tuff.	(Pratt, 1967.) New name adopted.	(Muffler, 1967.) Age changed from Oligocene or Miocene to Eocene	and Uligocene(?). (Plafker, 1967.) New name adopted.	(Mattson, 1967.) Andrecito Member of Laudon and Bowsher (1949)	adopted and made lowest member of formation. (Jones, Hernon, and Moore, 1967.) . New name adopted.	(Dibblee, 1967a.) Larsh Shale Member of Condra (1927) adopted. (Johnson and Adkison. 1967.)	Formerly Lawrence Shale. Base of formation low- ered to base of Haskell Limestone Member. For- mation includes (ascending order) Haskell Lime- stone, Robbins Shale, Ireland Sandstone, and Amazonia Limestone Members. (Johnson and Adkison, 1967.)
southeastern California	northeastern Kansas	west-central Puerto Rico	northeastern Kansas	do	east-central Alaska	northeastern Kansas	central Maine	southeastern Alaska	central Colorado	northeastern Kansas	northeastern Washington	southwestern New Mexico	southeastern Alaska	southern Alaska	west-central Puerto Rico	southwetsern New Mexico	southeastern California	northeastern Kansas	do-
Oligocene or early Miocene	Early Permian	Early Cretaceous	Early Permian	Late Pennsylvanian (Virgil)do	diddle or Late	Urdovician. Late Pennsylvanian (Virgil) northeastern Kansas.	Early Devonian	Late Triassic	Pleistocene	Late Pennsylvanian (Virgil)	Oligocene and Miocene(?)	Miocene(?)	Late Silurian	Eocene and Oligocene(?)	Late Cretaceous.	Early Mississippian	Pliocene(?)	Late Pennsylvanian (Virgil)	Pennsylvanian (Virgil)
Jackhammer Formation	Janesville Shale	Jayuya Tufi	Johnson Shale	Jones Point Shale Member (of Topeka Lime-	stone). Jones Ridge Limestone	Kanwaka Shale (of Shawnee Group)	Katahdin Quartz Monzonite	Keku Volcanics.	Kettle Creek Alluvium	King Hill Shale Member (of Lecompton Lime- Late Pennsylvanian (Virgil) northeastern Kansas	Klondike Mountain Formation	Kneeling Nun Tuff.	Kuiu Limestone	Kushtaka Formation	Lago Garzas Formation.	Lake Valley Limestone	Lane Mountain Andesite	tone)	Lawrence Formation (of Douglas Group)

Name	Age	Location	Revision and reference
Leach Canyon Tuff Member (of Quichapa Formation).	Oligocene or Miocene	southwestern Utah	Leach Canyon Tuff Member of Mackin (1960)
Lecompton Limestone (of Shawnee Group)	Late Pennsylvanian (Virgil) northeastern Kansas.		auopueu. (Averitt, 1967.) Phe following members have been adonted for the
			Lecompton Limestone (assending order): Spring Branch Limestone, Doniphan Shale, Big Springs Limestone, Quen Hill Shale, Bell Limestone, King Hill Shale, and Aroos Limestono Monhone
Lee Formation	. Early Pennsylvanian	east-central Kentucky	(Johnson and Adkison, 1967.) Includes Livingston Conglomerate Member in area of renort
Leeds Creek Member (of Twin Creek Lime- stone).	early Late Jurassic	Wyoming, Utah, and Idaho	(Weir, 1967.) New name adopted.
Lehman Ridge Gravel	Pleistocene (Nebraskan or Aftonian)	central Colorado	(Imlay, 1967.) New name adopted.
Lincoln Creek Formation	. late Eocene to early Miocene	southwestern Washington	(Varnes and Scott, 1967.) New name adopted.
Livingston Conglomerate Member (of Lee For- mation).	Early Pennsylvanian.	. east-central KentuckyI	(Beikman, Raú, and Wagner, 1967.) Livingston Conglomerate Member of Miller (1910) adonted.
Llaves Member (of San Jose Formation)	early Eocene	northwestern New Mexico	(Weir, 1967.) New name adopted.
Lodgepole Limestone (of Madison Group)	Early Mississippian	Montana, Wyoming, Utah, Idaho, and (subsurface) North Dakota.	(Baltz, 1967.) Cottonwood Canyon Member adopted as basal member of Lodeenole Limestone
Long Creek Limestone Member (of Foraker Limestone). Longmeadow Sandstone.	Early Permian Late Triassic		(Sandberg and Klapper, 1967) Age changed from Permian to Early Permian. Johnson and Wagner, 1967.)
McCloud Limestone	Early Permian (Leonard and Wolfcamp).		(Hatshorn and Koteff, 1967.) Age changed from Permian to Early Permian (Leonard and Wolfeamu).
Madera Limestone (of Magdalena Group)	Middle and Late Pennsyl- vanian and Early Permian (locally).	central New Mexico	Ketner, 1967.) In report area, age changed from Middle and Late Pennsylvanian to Middle and Late Pennsyl- vanian and Early Permian.
Madison Limestone	Early Mississippian and Devonian.	Utah, Idaho, Colorado, Wyoming, (Montana, and South Dakota.	(Myers, 1967.) Cottonwood Canyon Member adopted as basal member of Madison Limestone.
Maravillas Formation	Late Cretaceous	west-central Puerto Rico	(Sandberg and Klapper, 1967.) New name adopted. (Mattson, 1967.)

CHANGES IN STRATIGRAPHIC NOMENCLATURE-Continued

Age changed from Precambrian to middle Pre- cambrian.	Dutton and Linebaugh, 1967.) New name adopted. New name adopted. New name adopted. Tabor, 1967.) New name adopted. Loney, 1967., Loney, 1967.)	Martinisburg Formation in report area divided into (ascending order) Bushkill, Ramseyburg, and Pen Argyl Members. (Drake and Epstein, 1967.) New man adopted.	New name adopted. Epstein, Epstein, Spink, and Jennings, 1967.) New name adopted. Ekren and Frischknecht, 1967.)	New name adopted. (Dibblee, 1967a.) Mesquite Schist of Dibblee (1952) adopted.	(UIDDICE, 1967.4.) New name adopted. (Epstein, Epstein, Spink, and Jennings, 1967.) Age changed from Permian to Early Permian	(Leonard). (Neughan, 1967.) (New name adopted. (Sainsbury, 1967.) Aritzainsbury, 1967.)	Motitary Formatoria divided moves action accounding of the Dinosatr Canyon, Whitmore Point, and Spring- dals Sandstone Ambers. (Wilson and Skewart, 1967.) Morkton Formation used in southwestern Vermont. Morkton Outerstie in good usees elsewhere in	Vermont. (Zen, 1967.) Monstrate Formation of Pessagno (1960) adopted. (Mattson, 1967.) Montoya Group in report area. Montoya Dolomite	in good ussge elsewhere. (Fratt, 1967.) New name adopted. Varmes and Soott, 1967.) Naknek Formation divided into (ascending order) Chisik Conglomerate, lower sandstone, Snug Harbor Silistone, and Ponneroy Arkose Members. (Dotterman and Harksok, 1966.)
Age cam	(Dut New New (Tabo New New	and	(Epst New New (Ekre	New (Dibl Mesou			Din Din dal (Wils Monk	Ver Ver (Zen, Mons (Mati	in (Prat New Nakn Nakn Char Char Char
Michigan and Wisconsin	southwestern Alaska southeastern Kentuckydodo	Pennsylvania and New Jersey New Jersey, New York, Pennsylvania.	northeastern Maine	southeastern California	Pennsylvania and New Jersey Wyoming, Montana, North	Dakota, and South Dakota. northwestern Alaska	western Vermont	west-central Puerto Rico southwestern New Mexico	central Colorado
middle Precambrian	Early Jurassic	Ordovician	do	late Pliocene and (or) early Pleistocene(?). Precambrian(?)	Early Devonian Early Permian (Leonard)	sconsin)	Early Cambrian.	middle(?) Eocene	Pleistocene
Marinette Quartz Diorite n	Member (of Talkeetna me Member (of Lee me Tongue (of Lee	Martinsburg Formation	Maskenozha Member (of New Scotland Forma tion), Mattawamkeag Formation 0	Meeke Mine Formationla Mesquite SchistP	Minisink Limestone (of Helderberg Group) E Minnekahta Limestone E		Moenave Formation	Monserrate Formation m Montova Group M	Alluvium

-Continued	Revision and reference	 Cape Thompson Member adopted as member of Nascrak Formation. Needles Fange Formation of Mackin (1960) adopted. (Blank and Mackin, 1967.) Needles Fange Formation of Mackin (1960) adopted. (Blank and Mackin, 1967.) New name adopted. New name adopted. Consists of (ascending order) Maskenozha and Linestone, which remains in god usage outside treport area. (Epstein, Epstein, Spink, and Jennings, 1967.) Now Includes Austin Glen Member. (Sentistis, 1967.) Now Includes Austin Glen Member. (Sentistis, 1967.) Now Includes Austin Glen Member. (Sentistis, 1967.) Age changed from Permian to Early Permian. (Joinson and Wagner, 1967.) Age changed from Permian to Early Permian. (Dinbole, 1967.a) Age changed from Permian to Early Permian. (Dinbole, 1967.a) Age changed from Permian to Early Permian. (Dinson and Vagner, 1967.) Age changed from Permian to Early Permian. (Dinson and Vagner, 1967.) Age changed from Permian to Barly Permian. (Dinson and Vagner, 1967.) Age changed from Permian to Early Permian. (Dinson and Adrison.) 1967.) Onton and Adrison. (Dinson and Adrison.) 1967.) Onton and Adrison. (Dinson and Adrison.) 1967.) Onton and Adrison. (Dinson and Adrison.) 1967.) (Dinson and Adrison.) 1967.)<th></th>	
CHANGES IN STRATIGRAPHIC NOMENCLATURE-Continued	Location	northwestern Alaska	
ES IN STRATIGRAPH	Age	 Early and Late Mississippian. northwestern Alaska Eocene or early Oligoeene southwestern Utah Oligoeene(?) and lower or southeastern California middle Moceane New Jersey, New York sylyania. Pre-Carbonilerous	
CHANGI	Name	Nasorak Formation (of Lisburne Group)	

	Ozawkie Limestone Member (of Deer Creek Late Pennsylvanian (Virgil) northeastern Kansas.	northeastern Kansas	Ozawkie Limestone Member of Condra (1935)
Lunescone).			
Pablo Formation	- Permian (Ochoa?)	Nevada	Age changed from Permian(?) to Permian
Pavaloff Siltstone Member (of Chinitne Forme-	Lata Lirrassio	couthinizations Alasta	(Uchoa?). (Kether, 1967.)
tion).			New marine adoption. (Detterman and Hartsock, 1966.)
Pawlet Formation	- late Middle Ordovician	western Vermont	Pawlet Formation of Zen (1961) adopted.
Pedee Group	. Pennsylvanian	northeastern Kansas	Pedee Group no longer used in Kansas.
Pen Argyl Member (of Martinsburg Formation). Late Ordovician	- Late Ordovician	eastern Pennsylvania.	(Johnson and Adkison, 1967.) Pen Argyl Member of Behre (1927) adonted
Peppersauce Member (of Abrigo Formation)	- Middle(?) and Late Cambrian southern Arizona	southern Arizona	(Drake and Epstein, 1967.) Pernakes ind Mombur of Structure (1998) adverted
Percha Shale.	Late Devonian	southwestern New Mexico	Creasey, 1967b.) Perchas Shale divided into Ready Pay (lower) and
Perryville Member (of Lexington Limestone)	- Middle Ordovician	Kentuckv	Box (upper) Members. For the state of the st
			(Neuman, 1967b.)
Peters Valley Member (of Coeymans Formation)	Early Devonian	New Jersey and Pennsylvania	New name adopted.
Pickhandle Formation	- Oligocene (?) to middle Mio-	Southeastern California	(Epstein, Epstein, Spink, and Jennings, 1967.) Pickhandle Formation of Bowen (1954) adopted.
Pine Valley Gravel	cene(7). - Pleistocene (Illinoian and	central Colorado	(Dibblee, 1967a.) New name adopted.
Pioneer Formation	Sangamon age). - Precambrian	southern Arizona	967.) mation revised downward
			include the Scanlan Conglomerate Member in area of report.
Plum Creek Member (of Crab Orchard Forma-	Early Silurian	central Kentucky	(Creasey, 1967b.) Plum Creek Member of Foerste (1906) adopted.
Plymouth Marble	- Early Cambrian(?)	western Vermont	Age changed from Early Cambrian to Early
Pomeroy Arkose Member (of Naknek Forma- Late Jurassic.	Late Jurassic	southwestern Alaska	Camoran(7). (Zen, 1967.) Redefined to include Pomeroy Member and upper
tion).			sandstone member of Kirschner and Minard (1949).
Ponaganset Gneiss	- Mississippian(?) or older	northern Rhode Island	(Detterman and Hartsock, 1966.) New name adopted.
Portage Creek Agglomerate Member (of Tal-	Early Jurassic	southwestern Alaska	(Quinn, 1967.) New name adopted.
keenia Formation). Port Ewen Shale (of Helderberg Group)	Early Devonian	New York, New Jersey, and Penn-	(Detterman and Hartsock, 1966.) Port Ewen Shale used in area of report. (Transin, Dranch, Sciel, and Transing, 1027.)
Possession Drift	Pleistocene	northwestern Washington	New name adopted. (Easterbrook, Crandell, and Leopold, 1967.)

Name	Age	Location	Revision and reference
Poultney Slate	Early and Middle Ordovician. western Vermont.	western Vermont	- Age changed from Early Ordovician to Early and Middle Ordovician
Prida Formation (of Star Peak (Froup)	Early, Middle, and Late Triassic.	northwestern Nevada	(Zen, 1967.) (Zen, 1967.) - Age changed from Middle Triassic to Early, Middle, and Late Triassic.
Punchbowl Formation	late Miocene and early Pliocene.	southwestern California	(Silberling and Wallace, 1967.) - Age changed from late Miocene to late Miocene and early Pliocene.
Pybus Formation	Permian	southeastern Alaska	(Dibblee, 1967a.) - Formerly Pybus Dolomite.
Quail Lake Formation .	late Miocene	southeastern California	(Muffler, 1967.) - New name adopted.
oer (of Lecompton Lime-	Late Pennsylvanian (Virgil)	northeastern Kansas	(Dibblee, 1967a.) Queen Hill Shale Member of Condra (1927) adopted.
stone). Quichapa Formation	Oligocene or early Miocene	southwestern Utah	Johnson and Adkison, 1967.) - Quichapa Formation of Mackin (1960) adopted.
D_0	do	do	(Blank and Mackin, 1967.) - Quichapa Formation divided into (ascending order)
			Leach Canyon Tuff Member, Bauers Tuff Member, and Harmony Hills Tuff Member.
Rainvalley Formation (of Naco Group)	d Late(?) d and Gua	Permian Arizona and New Mexico	AF Changed from Permian to Early and Late(?) - Remian.
Ramseyburg Member (of Martinsburg Forma-	lupe(?)). Late(?) Ordovician	New Jersey and Pennsylvania	(McKee, 1967.) - New name adopted.
Rand Schist	Precambrian(?)	southeastern California	(Drake and Epstein, 1907.) - Rand Schist of Hulin (1925) adopted.
Ravena Member (of Coeymans Limestone)	Early Devonian	southeastern New York	(Dibblee, 1967a.) - Ravena Member of Rickard (1962) adopted.
Ready Pay Member (of Percha Shale)	Late Devonian	southwestern New Mexico	(Epstein, Epstein, Spink, and Jennings, 1967.) - Ready Pay Member of Stevenson (1944) adopted.
Red Buttes Quartz Basalt (of Tropico Group)	early Miocene(?)	southeastern California	Jones, Hernon, and Moore, 1967.) - Age changed from Pliocene(?) to early Miocene(?).
Reeve Meta-andesite	Early and Late Permian (Guadalupe).	California.	(Dibblee, 1967a.) - Age changed from Pennsylvanian to Early and Late Permian (Guadalupe).
Regina Member (of San Jose Formation)	early Eocene	northwestern New Mexico	. New name adopted.
Renegade Tongue (of Wasatch Formation)	Eocene	northeastern Utah	(Bauz, 1907.) - New name adopted.
Rich Member (of Twin Creek Limestone)	Middle Jurassic	Utah. Wyoming and Idaho	(Cashion, 1967b.) New name adonted

Ria Prista Farmation	middle Eorene	west-central Puerto Rico	New name adonted
Robbins Shale Member (of Lawrence Forma- Late Pennsylvanian (Virgil) northeastern Kansas	Late Pennsylvanian (Virgil)	northeastern Kansas.	(Mattson, 1967.) A member of the Lawrence Formation; formerly a
tion).			member of the Stranger Formation. (Ichnson and Adkison 1967)
Robinson Formation	. Early and Late Permian (Guadalupe).	California	Age changed from Pennsylvanian to Early and Late Pernian (Guadalupe).
Robles Formation	Early and Late Cretaceous west-central Puerto Rico.	west-central Puerto Rico	Age changed from Late Cretaceous to Early and
Rockabema Quartz Diorite	Ordovician	northern Maine.	Late Orteraction. (Matson, 1967.) New name adopted.
Rock Bluff Limestone Member (of Deer Creek	Late Pennsylvanian (Virgil) northeastern Kansas	northeastern Kansas	(Ekren and Frischknecht, 1967.) Rock Bluff Limestone Member of Condra (1927)
Limestone). Rondout Formation	Late Silurian and Early Devonian.	New York, New Jersey, and Pennsylvania.	adopted Adkison, 1967.) (Johnson and Adkison, 1967.) Rondout Formation in report area consists of (assenting) Duttorville, Whiteport Dolomite, and Mashipaconz Members.
Saddleback Basalt (of Tropico Group)	early Miocene(?)	southeastern California	(Epstein, Epstein, Spink, and Jennings, 1967.) Age changed from Pliocene(?) to early Miocene(?).
Saginaw Bay Formation	Mississippian and Pennsyl-	southeastern Alaska	(Dibblee, 1967a.) New name adopted.
Salem Gabbro-Diorite	vanian. pre-Carboniferous	eastern Massachusetts	(Mumer, 1967.) Age changed from early Paleozoic(?) to pre-Carbonif-
Salizvik (fravel	Tertiary or Quaternary	northwestern Alaska	erous. (Chute, 1966.) New name adonted.
Salvisa Bed (of Perryville Member).	Middle Ordovician	Kentucky	(Campbell, 1967.) Salvisa Bed of Miller (1913) adopted.
San Andres Limestone	Early and Late Permian (Leonard and Guadalupe)	west Texas and New Mexico	(Neuman, 1967b.) Age changed from Permian to Early and Late Per- mian (Leonard and Guadalupe).
San Angelo Sandstone (of Pease River Group) Early Permian (Leonard)	. Early Permian (Leonard)	west Texas.	Age changed from Permian to Early Permian
San Francisquito Formation	Paleocene and Eocene(?)	southeastern California	(Leonard). (Oriel, Myers, and Crosby, 1967.) New name adopted.
San Jose Formation	early Eocene	northwestern New Mexico	(Diubiee, 1964a.) Name adopted.
Sanpoil Volcanics.	Eocene(?)	northeastern Washington	Age changed from Eocene or Oligocene to Eocene(?).
Santa Margarita Formation	late Miocene and early Pliocene(?).	southeastern California	(Autessig, 1907.) Santa Margarita Formation restricted from area of the report, Quail Lake Formation used instead.
Sawatch Sandstone	Late Cambrian	central Colorado	Controller and a state of the second of Sawatch Sandstone used instead of Sawatch Quartraite in area of report. (Varues and Scott, 1967.)

	INVINITANTA NT CO	CHANGES IN STRATIGRAFHIC NOMENCLATURE-Continued	-Continued
Name	Age	Location	Revision and reference
Scanlan Congomerate Member (of Pioneer Formation or Shale).	Precambrian	southern Arizona	. Raised from bed to member.
Scatter Creek Rhyodacite	- Eocene(?)	northeastern Washington	Age changed from Eocene or Oligocene to
Scherrer Formation (of Naco Group)	Early Permian	Arizona	Eocene(?). (Muessig, 1967.) Age chanced from Permian to Early Permian
Second Value Dolomite (of Montoya Group)	. Middle and Late Ordovician . southwestern New Mexico	southwestern New Mexico	(McKee, 1967.) - Second Value Dolomite of Entwistle (1944)
Sedalia Dolomite	Early Mississippian	northeastern Kansas	 adopted. (Pratt, 1967.) . Edalia Dolomite used instead of Sedalia Limestone in on of the near the nea the near the near the near the near the near the near the ne
Sevy Dolomite	Early and Middle(?) Devonian.	east-central Nevada	(Johnson and Adkison, 1967.) Age changed from Early(?) and Middle Devonian to Early and Middle(?) Devonian.
Sharon Syenite	Carboniferous or older	eastern Massachusetts	(Drewes, 1967.) - Age changed from Devonian(?) to Carboniferous
Shawnee Island Member (of Coeymans Forma-	Early Devonian	Pennsylvania and New Jersey	or older. (Chute, 1966.) - New name adopted.
Sheldon Limestone Member (of Topeka Lime- stone).	Late Pennsylvanian (Virgil) northeastern Kansas.	northeastern Kansas	(Epstein, Epstein, Spink, and Jennings, 1967.) - Sheldon Linestone Member of Condra (1930) - adonted.
Shingle Pass Tuff.	Miocene.	southern Nevada	(Johnson and Adkison, 1967.) - Shingle Pass Tuff of Cook (1960) adopted.
Sixmile Creek Formation (of Bozeman Group)	Miocene and Pliocene	southwestern Montana	(Ekren, Rogers, Anderson, and Botinelly, 1967.) - New name adopted.
Skooner Gulch Formation.	early Miocene	northwestern California	(Robinson, 1967.) - Skooner Gulch Formation of Weaver (1944) adopted.
Sliderock Member (of Twin Creek Limestone)	Middle Jurassic	Wyoming, Utah, and Idaho	(Addicott, 1967.) - New name adopted.
Smoky Member (of Nopah Formation)	Late Cambrian	Nevada	. Used as Smoky Member of Nopak Formation in the Smotor Dame Merida
Snake River Group	Pleistocene and Recent	southeastern Idaho	Ross, 1967. Markey, Avyada. (Ross, 1967.) - Snake River Group includes Big Hole Basalt in eventheastern Joho
Snug Harbor Siltstone Member (of Naknek For- mation). Sonoma Range Formation	Late Jurassic	southwestern Alaska	Anundorff, 1967.) New name adopted. Detterman and Hartsock, 1966.) (Gilluly, 1967.)

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CONTRIBUTIONS TO STRATIGRAPHY

Sooka Formation	early Miocene(?)	Alaska	Age changed from Oligocene or Miocene, or both'
Southern Belle Member (of A hrizo Formation) Middle(?) and Late(?) Cam-	Middle(?) and Late(?) Cam-	southern Arizona	to early Micoene(?). (MacNeil, 1967) Southern Belle Member of Stoyanow (1936)
			adopted. (Creasey, 1967b.)
Spring Branch Limestone Member (of Lecomp- ton Limestone).	Late Pennsylvanian (Virgil) northeastern Kansas.	northeastern Kansas	Spring Branch Limestone Member of Condra (1927) adopted.
Star Peak Group	. Early to Late Triassic	northwestern Nevada	(Johnson and Adkison, 1967.) Star Peak raised to group rank everywhere.
Sterling Plutonic Group.	. Mississippian(?) or older	southeastern Connecticut	(Silbering and wanace, 1967.) Age changed from pre-Pennsylvanian to Mississip-
онно П. – Г. – П. – 10 – 10 – 10 – 10 – 10 – 10 – 10 – 1	Posts Damentronion	westam Montona	pian(?) or older. (Goldsmith, 1967.) New rows of orded
tion).	Poult Doronion	Ponneylrania and Naw Larcav	(Maughan and Protects, 1967.) Stormwills Memberts of White (1882) adouted
Strender Formation (of Douglass Formation)	. Daily Devoluanterretting and the strugging and the Late Pennsylvanian (Virgi) northeastern Kansas.	northeastern Kansas.	(Epstein, Epstein, Spink, and Jennings, 1967.) Base of Stranger Formation lowered to too of
			Stanton Limestone and included in the Douglas Group.
Stull Shale Member (of Kanwaka Shale)	do	do.	(Johnson and Adkison, 1967.) Stull Shale Member of Moore (1932) adopted.
Sugarlump Tuff	Miocene(?)	southwestern New Mexico	(Johnson and Adkison, 1967.) Formerly Sugarlump Tuffs.
Swakane Biotite Gneiss	. pre-Late Cretaceous	north-central Washington	(Pratt, 1967.) Age changed from pre-Tertiary to pre-Late Creta-
Svrena Formation	Late Pennsvlvanian	southwestern New Mexico	ceous. (Cater and Crowder, 1967.) Age changed from Pennsylvanian to Late Penn-
Talkeetna Rommation	Early Jurassic	southwestern Alaska	sylvanian. (Jones, Hernon, and Moore, 1967.) Talkeetna Formation divided into (ascending
			order) Marsh Creek Breccia Member, Portage Creek Aggiomerate Member, and Horn Mountain
Tallman Bandamanata	Parmian(%)	Nevada	. un member. (Detterman and Hartsock, 1966.) Age changed from Permian to Permian(?)
Tanicitos Member (of San Jose Formation)	early Eocene	northwestern New Mexico.	(Gilluly, 1967.) New name adopted.
	Mississippian and Penn-	west Texas	(Baltz, 1967.) Age changed from Pennsylvanian to Pennsyl-
	sylvanian.		vanian and Mississippian. (Oriel, Myers, and Crosby, 1967.)
Three C Member (of Abrigo Formation)	Middle and Late(7) Cambrian.	southern Arizona	New name adopted. (Creasey, 1967b.)
Tindir Group	. Precambrian	east-central Alaska	Age changed from Freeambrian and Early Cam- brian(?) to Precambrian. (Brabb, 1967.)

C NOMENCLATURE-Continued
STRATIGRAPHI
CHANGES IN

Name	Age	Location	Revision and reference
Topeka Limestone (of Shawnee Group)	Late Pennsylvanian (Virgil) northeastern Kansas	northeastern Kansas	 The Topeka Limestone has been divided into the following members: (ascending order) Hartford Limestone, Iowa Point Shale, Curzon Limestone, Jones Point Shale, Sheldon Limestone, and Turner Creek Shale Members.
Topsy Formation Towle Shale Member (of Onaga Shale) Tropico Group	post-early Oligocene(?) to pre-middle Miocene. Early Fermian	southeastern Alaskanortheastern Kansassoutheastern California	 (Johnson and Adkison, 1967.) New name adopted. (Plaffer, 1967.) Age changed from Fernian to Early Permian. (Johnson and Wagner, 1967.) Age changed from Micoene(?) and Pliocene(?) to Oliconanc(?) and Micoene(?)
Turner Creek Shale Member (of Topeka Lime- Late Pennsylvanian (Virgil) northeastern Kansas stone).	Late Pennsylvanian (Virgil)	northeastern Kansas	(Dibblee, 1967a.) - Turner Creek Shale Member of Condra (1927) adopted.
Twin Creek Limestone	Middle and Late Jurassic	Wyoming, Idaho, and Utah	(Johnson and Adkison, 1967.) . Twin Creek Linnestone divided into following members: (ascending order) Gypsum Spring, Sliderock, Rich, Boundary Ridge, Watton Canyon, Leeds Creek, and Giraffe Creek Mem- hore
Tyler Formation (of Amsden Group)	Early Pennsylvanian	Montana. Alaska	(Imlay, 1967.) . Tyler Formation of Freeman (1922) adopted. . Maaghan and Roberts, 1967. - Age charged from Mincens to middle Mincens
Upham Dolomite Member (of Second Value Middle and Late Ordovician southwestern New Mexico . Dolomite).	Middle and Late Ordovician	southwestern New Mexico	(MacNetl, 1967.) - Upham Dolomite Member of Second Value Dolo- mite of Montoya Group in report area. Elsewhere Upham Dolomite Member of Montoya Dolomite remains in coord usare.
Vale Formation (of Clear Fork Group)	Early Permian (Leonard)	west Texas	(Pratt, 1967),
Valmy Formation	Early, Middle, and Late Ordovician.	northwestern Nevada	(Oriel, Myers, and Crosby, 1967.) . Age changed from Early and Middle Ordovician De Early, Middle, and Late Ordovician in report
Viri Quartz Diorite Porphyry	Eocene(?)	west-central Puerto Rico	(Gilluly, 1967.) New name adopted. (Matkaon, 1967.) New name adopted. (Epstein, Epstein, Spink, and Jennings, 1967.)

Wastarquark Chert. Middle Octofedial northern Maine. New name adopted. Wastarquark Chert. Presembrian(7) settinged for the constraints. New name adopted. Wastarquark Chert. Presembrian(7) settinged for the constraints. New name adopted. Wastarguark Chert. Presembrian(7) setty and Middle Combrine. New name adopted. Wastarguark Present. Usb. Jable, Jable Member (of Twen Creek Limesent) New name adopted. New name adopted. Wester Shalp Member (of Strauger Formation). Pennsylvanin (Virgi). New name adopted. New name adopted. Withop Formation. Pennsylvanin (Virgi). New New Name adopted. New New Name adopted. New New Name adopted. Withop Marble Mar		Wasatch Formation	Paleocene and Eocene	northeastern Utah	. Includes Renegade Tongue at ton in area of renort
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Member (of Bull Forma- Early Cambrian(?) western Vermont		Yauco Mudstone	Late Cretaceous	west-central Puerto Rico	. Flatker, 1967.) - Yauco Mudstone of Mitchell (1922) adopted.
Member (of Bull Forma- Early Cambrian(?) western Vermont		York Glaciation	Pleistocene (Wisconsin age)		. New name adopted.
			Early Cambrian(?)	western Vermont	- Zion Hill Quartzite Member of Ruedemann (1914) adopted. (Zen, 1967.)

A21

UPPER PALEOZOIC FORMATIONS OF THE MOUNTAIN CITY AREA, ELKO COUNTY, NEVADA

By R. R. COATS

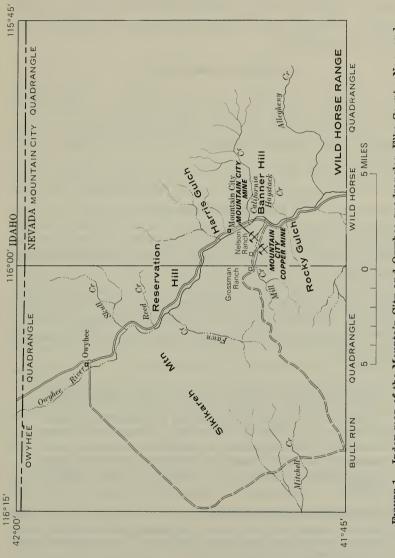
In the Mountain City area of Elko County (fig. 1), the Mountain City copper mine worked ores from a stratigraphically restricted zone of the Valmy Formation of Ordovician age. The Valmy is overlain by a sequence of rocks, predominantly clastic, belonging to the overlap assemblage of Roberts, Hotz, Gilluly, and Ferguson (1958, p. 2838). Some of these formations were described and named in private reports by geologists of the Mountain City Mining Company, during the period after the discovery of the Mountain City copper mine, and these names were later used by T. B. Nolan (unpub. data, 1932). Several of these names were published by Granger, Bell, Simmons, and Lee (1957, p. 116, pl. 14).

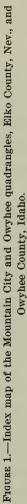
GROSSMAN FORMATION

The Grossman Formation is here named for rocks in the vicinity of the Grossman Ranch, in the valley of Mill Creek, about 1,000 feet west of the Owyhee-Mountain City quadrangle boundary. The type locality is designated as the hill in the SE¹/₄ sec. 4, T. 45 N., R. 53 E., about half a mile west of the Grossman house. Rocks of this formation crop out in the Mountain City and adjacent part of the Owyhee quadrangles in two areas. In the type area, they extend as patches along Mill Creek from the type locality eastward to the south slope of Banner Hill (California Hill on some maps), on the east side of Owyhee River. In the other area, about 2 miles farther south, they crop out in a roughly parallel belt in the valleys of Rocky Gulch and lower Haystack Creek.

The Grossman Formation is dominantly clastic, ranging from a coarse conglomerate to a siltstone and phyllite. The clasts in the conglomerate have been tectonically flattened; they consist chiefly of gray quartzite, black chert, phyllite, and magnetitiferous siltstone. The fresh rock is generally medium gray to greenish gray. In the coarser grained sandstone, a distinct pepper-and-salt appearance results from the contrast between the dark chert and the lighter colored quartzite.

The top and bottom of the formation are nowhere found in the same section, and attitudes are rarely determinable; the thickness of the formation may be as much as 2,000 feet. The formation rests unconformably on the Valmy Formation and is overlain unconformably by the Banner Formation.





As no identifiable fossils have been found, the age of the formation is unknown; because it is post-Valmy and pre-Banner, it must be post-Middle Ordovician and pre-Late Mississippian. The clastic character and obvious derivation from the Valmy suggest that the formation postdates the beginning of the Late Devonian Antler orogeny in the area where it was deposited. It is therefore tentatively assigned to the Devonian or Mississippian.

BANNER FORMATION

The name Banner Limestone was first published by Granger, Bell, Simmons, and Lee (1957, p. 116), who credited the name to T. B. Nolan (unpub. data, 1932). The name Banner Formation rather than Banner Limestone is here employed, because of the varied lithology of the formation.

The south slope of Banner Hill (shown as California Hill on some maps) is here designated the type locality. The formation crops out in a narrow belt which trends nearly east-west and extends from the south slope of Banner Hill (just northeast of the confluence of Mill Creek and the Owyhee River in the Mountain City quadrangle) discontinuously westward to the headwaters of Fawn Creek in the Owyhee quadrangle.

The thickness of the Banner Formation varies widely. The maximum thickness exposed where outcrops are reasonably continuous is about 600 feet. A section measured in some mine workings, now inaccessible (E. C. Stephens, written commun., 1964), shows the following:

Banner Formation:

Thickness (ft)

260
163
143
566

¹This thickness is greater than any seen on the surface. The mention of angular "shale" fragments suggests that conglomerate beds of part of the Grossman Formation were included in the measured section.

The Banner Formation rests with marked angular unconformity on the Valmy Formation and with slighter unconformity on the Grossman Formation. It grades upward, and perhaps also laterally, into the Nelson Formation; transitional material consists of a volcanic breccia, possibly a pépérite, with a calcareous matrix. The upper contact is drawn at the base of the rocks that contain a megascopically visible amount of volcanic material, probably more than 5 percent.

The Banner was originally considered to be of Late Mississippian age on the basis of determinations by G. H. Girty on several poor collections of fossils. Later, the views of paleontologists on the proper age assignment of this material changed, and an age of "probably Pennsylvanian or Permian(?)" was suggested (Granger and others, 1957, p. 116).

I was able to collect better preserved fossils in areas where the Banner is less metamorphosed; study of these permitted Helen Duncan (written commun., 1960) to correlate the Banner with limy shales in the Carlin area from which a Meramec (early Late Mississippian) fauna had been collected.

NELSON FORMATION

The Nelson Formation was originally called the Nelson Amphibolite by Granger, Bell, Simmons, and Lee (1957, p. 116), who credited the name to T. B. Nolan (unpub. data, 1932). It crops out as a westtrending belt of volcanic rocks adjacent to and north of the outcrop of the Banner Formation on Banner Hill and along the north side of Mill Creek. To the west, it is well exposed in the vicinity of the Nelson Ranch (S½ sec. 2, T. 45 N., R. 53 E.), the type locality; the Nelson there is approximately 600 feet thick, but it is somewhat thinner near the Owyhee-Mountain City quadrangle boundary and is much thicker in the headwaters of the east fork of Fawn Creek, where the discontinuous outcrops suggest a thickness of 2,500 feet.

The greater part of the formation consists of flows and tuff-breccias of andesitic and basaltic composition. The formation also includes minor sills of diabase and one lens of rhyolitic tuff. In the Mountain City quadrangle, it is a greenschist composed of tremolite-actinolite, chlorite, epidote, calcite, ilmenite, and andesine, in part altered to albite. In the meta-andesite, the amount of amphibolite is greatest and the recrystallized plagioclase is most calcic near the eastern end of the belt of exposures, where the rocks are closest to the quartz monzonite contact. Westward, the amphibole is lighter in color, the new feldspar is more albitic, and the rock is richer in chlorite. Locally the rock appears to be a metadiabase, probably metamorphosed sill, but exposures are not adequate to permit determination of the contact relations. The upper contact of the Nelson with the overlying Mountain City Formation is placed at the uppermost metavolcanic material. The gradation from limestone of the Banner Formation into pépérite and lava of the Nelson Formation suggests that the Nelson is essentially contemporaneous with the Banner Formation. The Nelson is therefore tentatively assigned a Late Mississippian age.

MOUNTAIN CITY FORMATION

The Mountain City Formation was originally named by Granger, Bell, Simmons, and Lee (1957, p. 116), who credited the name to T. B. Nolan (unpub. data, 1932). They (pl. 14) showed it occurred at the Mountain City mine (not to be confused with the Mountain City copper mine). This long-inactive gold mine in the SE1/4 sec. 2, T. 45 N., R. 53 E. at the crest of a hill north of Mill Creek is here designated as the type locality. The formation has been recognized only west of the Owyhee River, from the Mountain City mine to the headwaters of Fawn Creek. It forms a belt ranging in width from a few hundred feet to 21/2 miles. It consists largely of fine-grained dark-gray to black siliceous schists that have a poorly defined schistosity that commonly parallels the bedding. The matrix of the schist consists mainly of quartz, but includes minor amounts of orthoclase and much very fine carbonaceous matter. Small amounts of sericite and calcite are present. Locally, near the contact with the quartz monzonite, porphyroblasts of andalusite, sometimes greenish, are developed. A few thin limy beds, as much as 2 feet thick, are also present, and small needles of tremolite have formed in these as a result of metamorphism. The total thickness is unknown; however, the minimum original thickness was probably at least 4,000 feet and perhaps as much as 10,000 feet.

The Mountain City Formation rests conformably on the Nelson Formation, but the upper contact has not been seen in the Mountain City area; the formation is limited above by the thrust contact with the Reservation Hill Formation, by the intrusive contact of the quartz monzonite, or by the overlapping Tertiary volcanic rocks.

The age of the formation is unknown, because no identifiable fossils have been found in it. It is questionably assigned to the Carboniferous because of its conformable relationship with the underlying Nelson Formation.

RESERVATION HILL FORMATION

The Reservation Hill Formation is here named for rocks exposed on Reservation Hill, the type locality, in secs. 22 and 23, T. 46 N., R. 53 E., Mountain City and Owyhee quadrangles. The exposures are just east of the highway from Mountain City to Owyhee, Nev., and just inside the Western Shoshone Indian Reservation.

The principal rock type in the Reservation Hill Formation is a fine-grained dolomitic sandstone or siltstone, pale gray on fresh fractures. It weathers creamy white to pale reddish brown $(10R \ 6/4)$; the latter color is quite distinctive. The formation includes thick uninterrupted sequences in which siltstone beds, $\frac{1}{2}-2$ inches thick, alternate rhythmically with thinner phyllite beds. Microscopically, the principal constituents of this rock type are quartz, orthoclase, oligoclase, and dolomite; diopside and themolite, and locally, coarsely prismatic wollastonite have formed as a result of metamorphism. A less important rock type ranges from metagraywacke through micaceous and tremolitic quartzite to graphitic phyllite. Metacherts, biotitic where impure, and pure quartzite are rare. At least one meta-andesite is present, locally attaining a thickness of as much as 200 feet; it is metamorphosed to actinolite-epidote-plagioclase schist, or to hornblende-plagioclase schist. One bed of metarhyolite tuff, about 2 feet thick, was observed; others may be present. Lenses of gray siliceous dolomitic limestone, as much as 50 feet thick, are present locally.

The Reservation Hill Formation rests in thrust contact on the Mountain City Formation along the ridge north of Mill Creek and westward to the valley of Fawn Creek. It is intruded by a Cretaceous quartz monzonite and overlapped by Tertiary and Quaternary deposits.

No fossil evidence for the age of the Reservation Hill Formation has been found. Lithologically, it does not resemble either the known Precambrian rocks of this region or the known Triassic rocks and is therefore assumed to be Paleozoic. R. J. Roberts (oral commun., 1961) suggested that it somewhat resembles the type Havallah Formation of Pennsylvanian and Permian age, and thus it is tentatively assigned an age of Pennsylvanian(?) and Permian(?).

NEW FORMATIONS ON KODIAK AND ADJACENT ISLANDS, ALASKA

By George W. Moore

Formal stratigraphic names have not previously been applied to rocks on Kodiak Island (fig. 2). The island is about 100 km (kilometers) wide and 160 km long and lies south of the center of the Alaskan subcontinent, about 50 km from the mainland. The formations newly named below were studied during the summers of 1962, 1963, and 1965.

UYAK FORMATION

The Uyak Formation is here named for the village of Uyak, which is on the northwest coast of Kodiak Island on the west shore of Uyak Bay (fig. 2). The formation crops out as a belt about 10 km wide along the northwest coast of Kodiak, Uganik, Raspberry, and Afognak Islands. It corresponds to the northwest belt of the greenstone-schist group of Capps (1937).

The type section of the Uyak Formation is designated as the rocks exposed along the west shore of Uyak Bay, from a basal thrust fault 3 km south of Uyak to a point 4 km northwest of Uyak, where the formation goes under a cover of glacial drift. The formation is sheared and cut by faults, but its attitude is fairly uniform; beds strike N. 45° E. and dip 75° NW. The tops of the beds were identified at only a few places, but in each place the beds are upright. If the type section is continuous, a thickness of about 6,000 meters is exposed.

The principal rock types throughout the Uyak Formation are black shale and local schistose green tuff. A few shale and sandstone graded beds also occur. The middle third of the exposed part of the formation contains many beds of pillow basalt and red chert, and the upper third is characterized by thick layers of light-gray chert. Thin limestone lenses occur at two places in the formation : about 700 meters above the base and 400 meters below the top. The chert and basalt underlie ridges, whereas the tuff and shale underlie valleys.

The Uyak Formation is the oldest formation exposed on the Kodiak group of islands. It is thrust over younger rocks to the southeast; elsewhere its relationship to younger sedimentary formations is obscured by the water of Shelikof Strait.

Marine fossils collected from a limestone lens 700 meters above the base of the type section at 57°37.7′ N., 153°59.0′ W., and identified by N.J. Silberling (written commun., 1966) are of Late Triassic age. The Uyak Formation is therefore considered to be Triassic. It has been intruded by penecontemporaneous ultramafic rocks and by a middle Tertiary quartz diorite batholith.

The Uyak Formation correlates with Triassic rocks on the Kenai Peninsula, including chert- and basalt-bearing rocks that extend to Nuka Island Passage (Grant, 1915). It is equivalent to somewhat less deformed and more fossiliferous Triassic rocks directly across Shelikof Strait on Cape Kekurnoi.

KODIAK FORMATION

The Kodiak Formation is here named for Kodiak Island. It crops out along the center of the island in a northeast-trending belt about

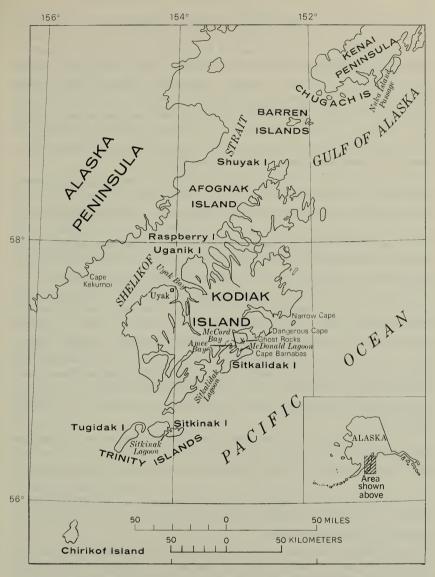


FIGURE 2.-Index map of Kodiak and adjacent islands.

60 km wide that follows an anticlinorium overturned toward the southeast. The Kodiak Formation is the slate-graywacke group of Capps (1937).

The northwest flank of the anticlinorium is a fairly regular homocline in which the attitude of the bedding averages N. 45° E., 45° NW. The type section of the Kodiak Formation is designated as the section along the west shore of Uyak Bay, a fiord which nearly bisects Kodiak Island. The base of the section is 4 km south of the head of the bay, where the lower part of the formation is intruded by a quartz diorite stock. The top of the section is 3 km south of the village of Uyak, where Triassic rocks have been thrust over the youngest exposed part of the Kodiak Formation. Faults and dikes cut the unit, but except locally near the base, the beds everywhere in the type section are upright. A calculation based on the average dip indicates that the Kodiak Formation in its type section is 30,000 meters thick.

This exceedingly thick formation consists of a very regular geosynclinal sequence of graded beds that average about 1 meter thick. The fine-grained layer at the top of each graded bed is generally slate, and the coarser layer is medium-grained sandstone. In some places in the lower part of the formation, the fine-grained layers are phyllitic, and in approximately the upper 4,000 meters, they are shaly. The Kodiak Formation is resistant to erosion and provides excellent outcrops.

Along the axis of the anticlinorium, the Kodiak Formation has been intruded by a middle Tertiarv quartz diorite batholith, and no stratigraphic units older than the Kodiak occur there. On the southeast side of Kodiak Island, a continuous sequence with younger formations seems to exist, but the rocks are so sheared and deformed that all mapped formation boundaries are faults. The Kodiak Formation is distinguished from the next younger formation there primarily by the Kodiak Formation's lack of basalt and secondarily by its slaty foliation.

No fossils diagnostic of age were found in the formation during the present investigation, but Imlay and Reeside (1954, p. 228) reported that Ulrich (1904, pl. 12–13) described two species of *Inoceranus*—from approximately the middle of the formation—that occur elsewhere in the world in rocks of Late Cretaceous age. The Kodiak Formation is therefore considered here to be Cretaceous on the basis of superposition and this fauna.

The Kodiak Formation also underlies the southeastern part of Afognak Island. It is lithologically correlative on the Kenai Peninsula with the slate and gravwacke between Nuka Island Passage and Resurrection Bay (Grant, 1915), and to the southwest on the Shumagin Islands, it is correlative with the Shumagin Formation of Burk (1965). The deep-water sedimentary deposits of the Kodiak Formation correlate on the Alaska Peninsula to the northwest with Cretaceous shallow marine and continental coal-bearing rocks studied by Jones and Detterman (1966) and by Burk (1965).

GHOST ROCKS FORMATION

The Ghost Rocks Formation is named here for Ghost Rocks, which lie on the southeast coast of Kodiak Island directly north of Sitkalidak Island (fig. 2). The unit is equivalent to the southeast belt of the greenstone-schist group of Capps (1937). It crops out in a belt about 10 km wide near the southeast shore of Kodiak Island and through the middle of Sitkalidak Island.

The Ghost Rocks Formation occurs approximately along the axis of an isoclinally folded synclinorium that trends northeast. The synclinorium has been cut by block faults in such a way that wherever the rocks have been carefully studied, older rocks to the northwest and younger rocks to the southeast rest against rocks of the Ghost Rocks Formation in fault contact.

The type section of the Ghost Rocks Formation is designated as the exposures along the north coast of Sitkalidak Island directly opposite Ghost Rocks across Sitkalidak Strait. The type section extends from a major fault on the west side of the head of Amee Bay to a fault on the east side of the mouth of McDonald Lagoon. Specifically, in this isoclinally folded section, the Ghost Rocks Formation includes (1) zeolite-bearing tuffaceous sandstone that crops out at the heads of McCord Bay and Sitkalidak Lagoon, (2) all beds of basalt lying along the synclinorium, and (3) all intervening rocks, consisting of hard claystone, sandstone, tuff, and graded beds, locally in the form of wildflysch. The formation is sheared, faulted, and folded. The internal stratigraphy has not been completely worked out, but the thickness appears to be approximately 5,000 meters. The formation is lithologically distinct, as it is a coherent belt of rocks that contains pillow basalt and tuff, which are not found in either the underlying or the overlying formations.

Capps (1937) correlated this belt of rocks with my Uyak Formation on the northwest side of Kodiak Island, which similarly contains pillow basalt; hence he considered it to be older than my Kodiak Formation. Where the Ghost Rocks Formation rests in fault contact with the Kodiak Formation, however, lithologies that were originally similar in each are more highly metamorphosed in the Kodiak Formation. Moreover, the lack of chert and limestone in the Ghost Rocks Formation makes it different from the Uyak Formation.

No fossils were found in the Ghost Rocks Formation. The similarity in geosynclinal character to formations that are older and younger suggests that the sequence of formations lacks unconformities. Also, an analysis of the thicknesses of similar lithologies in the successive formations and a hypothetical correlation of the volcanic rocks in the Ghost Rocks Formation with nearby intrusive rocks of known age suggest that the formation straddles the boundary between the Paleocene and the Eocene.

To the west, on the Alaska Peninsula, the Ghost Rocks Formation is believed to correlate with the lower part of the basalt-bearing Tolstoi Formation of Burk (1965), which contains fossils of Paleocene age. To the east in Prince William Sound, it correlates with the lower part of the basalt-bearing Orca Group, in which fossils of early Tertiary age have been found (Plafker and MacNeil, 1966, p. B67).

SITKALIDAK FORMATION

The Sitkalidak Formation is named here for Sitkalidak Island; off the southeast coast of Kodiak Island. The formation occurs mainly in a series of patches at the southeastern tips of points on Kodiak, Sitkalidak, and Sitkinak Islands. These points are generally separated by fords, so the individual patches are about 20 km apart. The formation has been deformed into a series of tight folds that are commonly overturned.

The type section of the Sitkalidak Formation lies along the north coast and near the east end of Sitkalidak Island. It extends in a chiefly overturned section from the axis of an anticline, 3 km northwest of Cape Barnabas, to the base of a 20-meter conglomerate bed, 7 km northwest of Cape Barnabas. The formation consists of a rather uniform sequence of sandstone and siltstone graded beds about 3,000 meters thick that formed under geosynclinal conditions of deposition. A few conglomerate beds also occur in the unit. All mapped contacts with the underlying Ghost Rocks Formation follow faults, and the basal contact of the Sitkalidak Formation has not been reached at the base of the type section. The base is defined to be directly above the uppermost basalt or tuffaceous sandstone bed that marks the top of the Ghost Rocks Formation. The upper contact, except in the type section, where it is deliberately specified, is a transitional zone in which the graded beds of the Sitkalidak Formation (below) alternate with crossbedded sandstone or conglomerate (above) that contains coal fragments.

A fossil clam collected from about 300 meters below the top of the Sitkalidak Formation in the type area at 57°11.1' N., 152°56.5' W., is a new genus of Vesicomyidae, and a fossil crab, *Callianassa* aff. C. *porterensis*, from the same locality, indicates an Oligocene age (F. S. MacNeil, written commun., 1963). Evidence from superposition suggests that the Sitkalidak Formation is Eocene and Oligocene. To the northeast in Prince William Sound, the formation correlates with the upper part of the Orca Group (Plafker and MacNeil, 1966); on the Alaska Peninsula, it correlates with the upper part of the Tolstoi Formation of Burk (1965).

SITKINAK FORMATION

The Sitkinak Formation is named here for Sitkinak Island, which is about 30 km long and which lies 15 km southwest of Kodiak Island. The formation is in isolated patches along a belt about 250 km long, extending from Chirikof Island at the southwest to Dangerous Cape on Kodiak Island at the northeast. The type section is along the south shore of Sitkinak Island. The basal part includes beach and shallow-marine deposits, but the bulk of the formation is continental and consists of coal-bearing siltstone, sandstone, and conglomerate. In the type section, several half-meter-thick coal beds occur in association with well-preserved fossil leaves.

A complete, though folded, section of the Sitkinak Formation occurs at its type locality, where the formation is about 1,500 meters thick. The basal contact there, which is locally disturbed by faulting, intersects the south shore of Sitkinak Island at a small lagoon 800 meters east of the south entrance to Sitkinak Lagoon at the base of an alternating zone, where graded beds of the underlying Sitkalidak Formation are succeeded by conglomerate and crossbedded sandstone and siltstone that contain coal fragments. At the type section of the Sitkalidak Formation on Sitkalidak Island, the contact is at the stratigraphic base of an overturned 20-meter conglomerate bed 7 km northwest of Cape Barnabas.

At many of its known areas of occurrence, the Sitkinak Formation is the youngest bedrock unit present. In its type section, however, coalbearing sandstone and siltstone are conformably overlain by marine siltstone containing lower Miocene fossils. The upper contact in the type section is about 200 meters west of the southernmost point of Sitkinak Island.

Identifiable plant fossils are abundant throughout the continental part of the Sitkinak Formation on Sitkinak Island. J. A. Wolfe (written commun., 1968) stated that collections from near the middle and near the top of the section are middle or late Oligocene in age. The Sitkinak Formation is considered here to be Oligocene. It correlates on the Alaska Peninsula with the coal-bearing Stepovak Series of Palache (1904; Stepovak Formation of Burk, 1965). On the east side of the Gulf of Alaska, it correlates with the lower parts of the Poul Creek and Katalla Formations (Plafker, 1967).

NARROW CAPE FORMATION

The Narrow Cape Formation is here named for Narrow Cape, near the east end of Kodiak Island. The type section is along the southwest coast of the cape, from its end northwestward about 1 km to the axis of a syncline where the youngest part of the formation is exposed. At its type locality, the formation is 700 meters thick. The lower twothirds consists of sandstone and a few conglomerate beds; the upper third consists of siltstone. On Sitkinak Island, 150 km southwest of the type locality, about 150 meters of siltstone is preserved along a synclinal axis.

At Narrow Cape, the formation rests unconformably on the Sitkalidak Formation of Eocene and Oligocene age. On Sitkinak Island, it rests conformably on the Sitkinak Formation of Oligocene age. The Narrow Cape Formation is the youngest bedrock formation exposed at each of these two areas of outcrop.

The Narrow Cape Formation contains a rich marine fauna. A collection from near the middle of the section at the type locality was determined by F. S. MacNeil (written commun., 1963) to be middle Miocene. A collection from near the base of the formation on Sitkinak Island was determined by MacNeil to be early Miocene. The age of the formation therefore is considered to be Miocene. The Narrow Cape Formation correlates on the Alaska Peninsula with the Bear Lake Formation of Burk (1965). On the east side of the Gulf of Alaska, it correlates with the upper parts of the Poul Creek and Katalla Formation and the lower part of the Yakataga Formation (Plafker, 1967).

TUGIDAK FORMATION

The Tugidak Formation is here named for Tugidak Island, which is about 10 km wide and 20 km long; this island lies approximately 25 km southwest of Kodiak Island. The Tugidak Formation is the only bedrock unit underlying the island, and it occurs there in a homocline dipping approximately 5° NE. The formation also occurs at the north end of Chirikof Island, where it dips about 10° N.

The type section of the Tugidak Formation is designated as the exposures along the west coast of Tugidak Island, from the south tip of the island to the northernmost exposure of bedrock. In its type section, the formation is approximately 1,500 meters thick. It consists of interbedded sandstone and siltstone characterized by randomly distributed pebbles and cobbles. A 1-meter-thick cobble-conglomerate bed occurs about 350 meters above the base. On Chirikof Island, the Tugidak Formation lies in fault contact with older rocks and is overlain with apparent conformity by an unnamed Pleistocene marine formation.

The Tugidak Formation is richly fossiliferous on Tugidak Island. Three collections of marine fossils spaced stratigraphically through the formation were determined to be of Pliocene age by F. S. MacNeil (written commun., 1963). A fossil snail, *Nassarius* cf. N. Andersoni, in float from near the base of the formation on Chirikof Island is also considered to be Pliocene by W. O. Addicott (oral commun., 1966). The Tugidak Formation correlates with the Tachilni Formation on the Alaska Peninsula (Burk, 1965). The upper part of it correlates with the "lower" part of the of the Yakataga Formation as exposed on Middleton Island, about 400 km northeast of Kodiak Island (George Plafker and F. S. MacNeil, oral commun., 1963); only the uppermost part of the Yakataga Formation is exposed on Middleton Island.

THREE NEWLY NAMED JURASSIC FORMATIONS IN THE MCCARTHY C-5 QUADRANGLE, ALASKA

By E. M. MACKEVETT, Jr.

This report names and describes three Jurassic formations in the McCarthy C-5 quadrangle, Alaska: the Lubbe Creek, Nizina Mountain and Root Glacier Formations. Brief descriptions of the stratigraphy of these formations were included in a report by MacKevett and Imlay (1962), and their distributions were shown in a preliminary geologic map of the quadrangle (MacKevett, 1963). Earlier investigators either did not recognize the Jurassic rocks or lumped them with Cretaceous or Triassic rocks. Moffit (1938, pl. 2), however, showed two small patches of undifferentiated Jurassic rocks within the C-5 quadrangle.

The present report supplements previous reports by naming the formations, revising lithologic descriptions and age assignments, and providing additional stratigraphic information. This report is based on field investigations during the summers of 1961 and 1962, on pertinent laboratory studies, and on paleontologic studies by R. W. Imlay. M. C. Blake, Jr., ably assisted in the field during 1961, and the writer is grateful to him and to Imlay for their contributions.

The McCarthy C-5 quadrangle is on the south flank of the rugged and strongly dissected Wrangell Mountains (fig. 3). It is bounded by the 61°30' and 61°45' parallels and by the 142°30' and the 142°52'30'' meridians. Alpine conditions reflected by glaciers, perennial snowfields, diverse surficial deposits related to glacial activity, and several arêtelike ridges characterize a large part of the quadrangle. The quadrangle is uninhabited, although the two largest of the famous Kennecott mines, the Bonanza and the Jumbo, are near its southwestern corner and formerly were thriving mining camps. The most practical mode of travel in the quadrangle is the helicopter.

The Lubbe Creek, Nizina Mountain, and Root Glacier Formations occupy part of a belt of Triassic and Jurassic rocks that trends

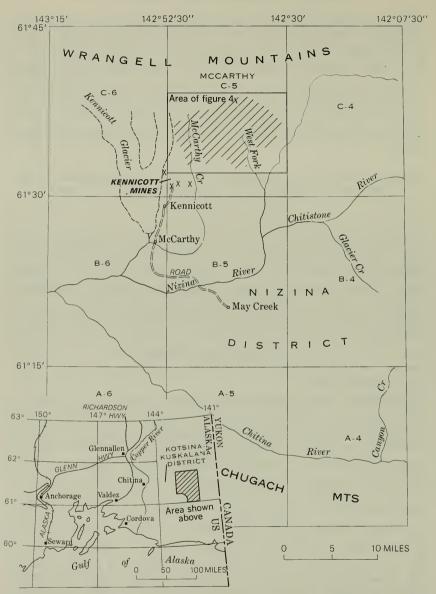


FIGURE 3.-Index map showing the McCarthy C-5 and nearby quadrangles.

northwestward across the middle of the quadrangle (fig. 4). These formations represent the upper part of an essentially concordant stratigraphic sequence whose sediments were deposited from the Late Triassic into the Late Jurassic. The major orogeny in the region was near the end of the Jurassic and (or) in the Early (pre-Albian) Cretaceous; the Jurassic rocks have been folded and faulted and are overlain unconformably by younger deposits. Despite modifications caused by the numerous folds and faults, the Jurassic rocks simulate a gentle northeast-dipping homocline in gross structure. The Lubbe Creek, Nizina Mountain, and Root Glacier Formations consist largely of very fine grained and fine-grained epiclastic rocks.

LUBBE CREEK FORMATION

Name and distribution

The Lubbe Creek Formation is named here from the excellent exposures at its type locality, along Lubbe Creek, a westward-flowing tributary of McCarthy Creek (figs. 4, 5). The formation also crops out on Bonanza Ridge, in the canyon walls that border the upper reaches of McCarthy Creek, and on both sides of the West Fork (fig. 4), all within the McCarthy C-5 quadrangle.

Two small patches of the Lubbe Creek Formation are exposed near the southwest corner of the McCarthy C-4 quadrangle (MacKevett and others, 1964), and the formation also extends into the McCarthy C-6 quadrange (MacKevett, 1965). No similar coeval rocks have been described from other nearby parts of Alaska.

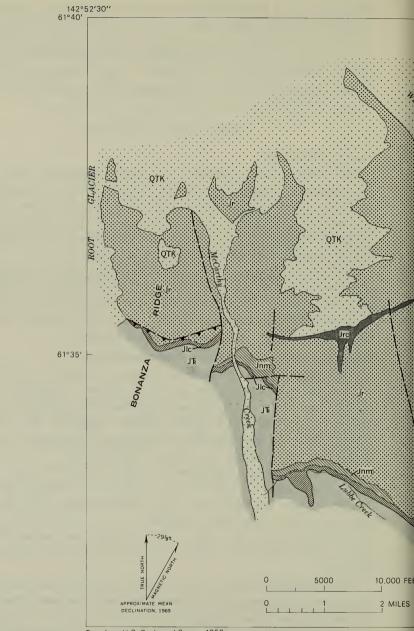
General character, stratigraphic relations, and thickness

The Lubbe Creek Formation consists of impure spiculite and subordinate amounts of coquina. It constitutes an excellent marker horizon, being thin and lithologically distinctive, forming bold outcrops, and containing abundant fossils, including the diagnostic Weyla. Strata in the formation commonly are between $\frac{1}{2}$ and 3 feet thick and, exceptionally, as much as 8 feet thick.

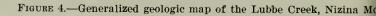
The Lubbe Creek Formation conformably overlies the upper member of the McCarthy Formation (figs. 5, 6), which is largely or entirely Early Jurassic. Its upper contact is a disconformity that separates it from the Nizina Mountain Formation or, in some places, from the Root Glacier Formation (figs. 5, 6). Locally, the Lubbe Creek is in fault contact with other rocks or is overlain by Quaternary surficial deposits.

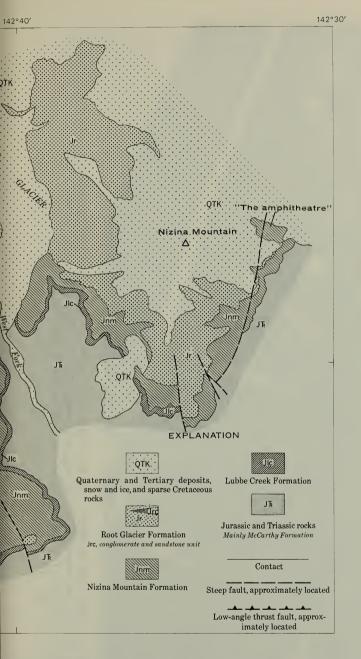
The formation is about 300 feet in maximum thickness, but throughout most of its extent it is thinner because of erosion. Southeast of Nizina Mountain, the stratigraphic interval normally occupied by the Lubbe Creek Formation is marked by a hiatus that separates the upper member of the McCarthy Formation from the Nizina Mountain Formation.

CONTRIBUTIONS TO STRATIGRAPHY



Base from U S Geological Survey, 1959





and Root Glacier Formations in the McCarthy C-5 quadrangle.

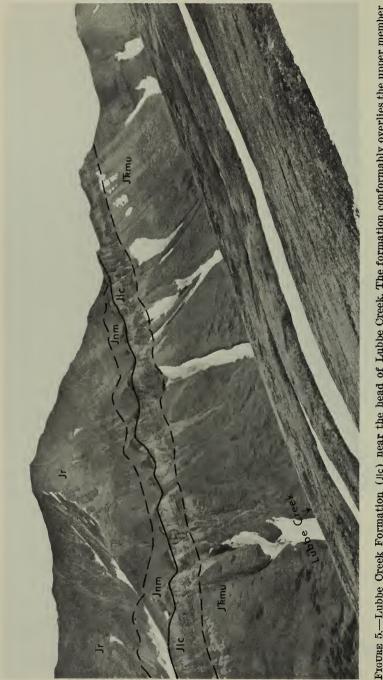


FIGURE 5.-Lubbe Creek Formation (JIc) near the head of Lubbe Creek. The formation conformably overlies the upper member of the McCarthy Formation (JFmu) and is overlain disconformably by the Nizina Mountain Formation (Jnm) or the Root Glacier Formation (Jr). The Lubbe Creek Formation here is about 150 feet thick.

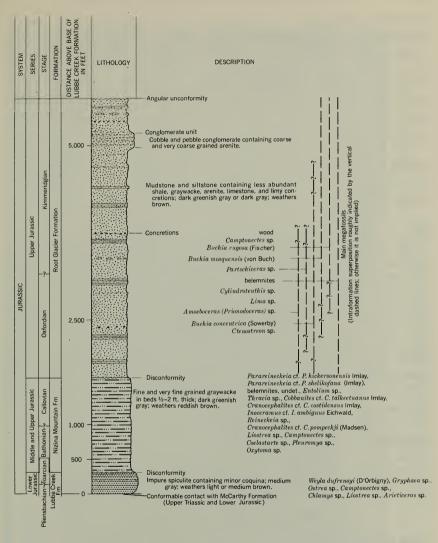


FIGURE 6.—Composite columnar section showing the Lubbe Creek, Nizina Mountain, and Root Glacier Formations.

Petrology and petrography

The Lubbe Creek Formation consists of impure spiculite and minor amounts of coquina; the rocks are medium gray where fresh and light or medium brown where weathered. Chert lenses as much as 10 feet long and 6 inches thick occur in some of the spiculites. The impure spiculites are fine-grained silica-rich rocks that contain organic and inorganic clasts in a chalcedony matrix. The clasts include spicules, Radiolaria, fragments of pelecypods and belemnites, calcite, quartz, dolomite, and plagioclase. Calcite is the dominant mineral of the clastic grains and of most shell fragments. The spiculites also contain minor amounts of chlorite, hematite, pyrite, carbonaceous material, ilemnite, biotite, and apatite. Most of the clasts are ragged to subangular and less than 0.2 millimeter in maximum dimension. The spicules are generally less than 0.3 mm long; commonly they are chalcedonic, and uncommonly, calcareous. A few have chloritic cores.

The coquina contains abundant poorly sorted bioclastic material, chiefly shells and shell fragments of megafossils, in a chalcedony matrix. Its subordinate clastic constituents are similar to those of the impure spiculites.

Age

The Lubbe Creek Formation is late Early Jurassic in age. The ammonite Arieticeras (USGS Mesozoic loc. 28531) provides excellent evidence for a late Pliensbachian age. The early Pliensbachian ammonite Uptonia (Mesozoic loc. 28675), less than 100 feet below the base of the Lubbe Creek Formation in the C-4 quadrangle, shows that no part of the formation is older than Pliensbachian. That the age is no younger than Toarcian is shown by an abundance of the Early Jurassic pelecypod Weyla in the upper part of the formation. Identification of this pelecypod as Weyla dufrenoyi (D'Orbigny) (Prof. S. W. Muller, oral commun., 1964) suggests a Toarcian age for the upper part of the formation.

In addition to Weyla, the fauna of the Lubbe Creek Formation, as identified by R. W. Imlay of the U.S. Geological Survey (written commun., 1963), includes: Prodactylioceras? sp., Arieticeras? sp., Arieticeras sp., Gryphaea sp., Ostrea sp., Camptonectes sp., Astarte sp., Eopecten sp., Chlamys sp., Liostrea sp., and brachiopods.

NIZINA MOUNTAIN FORMATION

Name and distribution

The Nizina Mountain Formation is named here for its type locality outcrops that partly girdle the ridge that extends southward from Nizina Mountain (fig. 4). The formation is well exposed on canyon walls of the West Fork. From these exposures it extends eastward around the nose of the ridge south of Nizina Mountain. Northeastward from there, it crops out almost continously along the east side of the ridge and extends to near "the amphitheatre," where it is overlapped by younger rocks (figs. 4, 7). The formation is intermittently exposed along the north side of Lubbe Creek and near the head of McCarthy Creek.



Except for one isolated outcrop in the McCarthy C-6 quadrangle (MacKevett, 1965), the Nizina Mountain Formation has been recognized only in the McCarthy C-5 quadrangle.

General character, stratigraphic relations, and thickness

The Nizina Mountain Formation consists dominantly of fine-grained to very fine grained graywacke 1 in distinct beds 1/2-2 feet thick. Most outcrops of the formation underlie moderate slopes that have reddishbrown weathered surfaces. The formation's upper and lower contacts are unconformities, commonly disconformities, but east of Nizina Mountain its upper contact is an angular unconformity with post-Jurassic rocks (fig. 7). In most places, the Nizina Mountain Formation disconformably overlies the Lubbe Creek Formation (figs. 4, 5, 6), but southeast of Nizina Mountain it disconformably overlies the upper member of the McCarthy Formation. At most places, the Nizina Mountain Formation is overlain disconformably by the Root Glacier Formation (figs. 5, 6). In a few places, it is separated by an angular unconformity from overlying Cretaceous (Albian) sedimentary rocks or from Tertiary rocks that are part of the Wrangell Lava (fig. 7). Quaternary surficial deposits locally cover the Nizina Mountain Formation

The Nizina Mountain Formation is about 1,350 feet in maximum thickness at its type locality. Elsewhere, because of extensive erosion, it is thinner, and many of its outcrops are only a few tens of a few hundreds of feet thick.

Petrology and petrography

The dominant fine-grained to very fine grained graywackes of the formation are associated with sparsely distributed shaly partings and a few limy lenses and concretions. The graywackes are mainly dark greenish gray and weather reddish brown. They consist of poorly sorted subangular to subrounded clasts in extremely fine-grained chalcedonyrich matrices. The clasts range from 0.05 to 0.5 mm in maximum dimension. They are composed of plagioclase, generally sodic labradorite, and quartz, along with less abundant cherty lithic fragments, calcite, opaque minerals, and rare biotite and chlorite; they also contain clay minerals, opaque dust, and sparse laumontite (leonhardite), epidote, prehnite, and calcite. Opaque minerals in the Nizina Mountain Formation include pyrite, hematite, ilmenite, and magnetite.

¹Nomenclature of the sandstones in this report follows the usage of Williams, Turner, and Gilbert (1954, p. 289-321).

Age

The Nizina Mountain Formation is Middle and Late Jurassic and probably includes strata representative of both the Bathonian and Callovian stages. Its large fauna chiefly consists of belemnites that are as much as 5 centimeters long and 0.5 cm wide and of poorly preserved ammonites. Fossils from the formation that were identified by R. W. Imlay (written commun., 1963) include the following: *Parareineckeia* cf. *P. hickersonensis* Imlay, *P.* cf. *P. shelikofana* (Imlay), *P.* sp., *Entolium* sp., *Thracia* sp., *Cobbanites* cf. *C. talkeetnanus* Imlay, *Cranocephalites* cf. *C. costidensus* Imlay, *C.* cf. *C. pompeckji* (Madsen), *Arctocephalites* ? sp., *Inoceramus* cf. *I. ambiguus* Eichwald, *Plesiopecten*? sp., *Coelastarte* sp., *Pleuromya* sp., *Liostrea* sp., *Camptonectes* sp., *Oxytoma* sp., *Quenstedtia*? sp., and undetermined belemnites, aptychus, fish scales, and a crustacean appendage.

ROOT GLACIER FORMATION

Name and distribution

The Root Glacier Formation is applied here to the thick dominantly very fine grained and clastic marine sedimentary rocks that are well exposed on the arêtelike ridges that border Root Glacier (Bonanza Ridge in the McCarthy C-5 quadrangle and the ridge that extends northward from Donoho Peak in the McCarthy C-6 quadrangle). Because of structural complications on these ridges, including a thrust fault, the type locality of the Root Glacier Formation is designated as the hillside east of the upper part of McCarthy Creek (fig. 4), where the Root Glacier rocks are also fairly well exposed but less deformed. The formation also includes a facies conglomerate and very coarse grained sandstone.

The Root Glacier Formation extends eastward from the margins of Root Glacier to the slopes southeast of Nizina Mountain (fig. 4). It is well exposed on Bonanza Ridge, the environs of upper McCarthy Creek, the ridge between McCarthy Creek and the West Fork, and near the West Fork Glacier (figs. 4, 5, 8). The conglomerate crops out on the ridge between McCarthy Creek and the West Fork and forms strike-controlled spurs that extend laterally from the main ridge (figs. 4, 8).

The Root Glacier Formation extends northwestward across the McCarthy C-6 quadrangle (MacKevett, 1965) toward the northeast corner of the C-7 quadrangle. The formation has not been recognized in other nearby quadrangles, but it may correlate with some of the Jurassic or Cretaceous rocks of Moffit (1938, p. 66-68) from the Kotsina-Kuskalana region.

General character, stratigraphic relations, and thickness

The formation mainly consists of poorly sorted and poorly bedded pelitic and fine psammitic rocks. Characteristically these rocks form moderately smooth slopes (figs. 5, 8) that locally are breached by a few resistant limy beds and lenses and by outcrops of coarse clastic rocks. The conglomerate forms bold outcrops that protrude from the surrounding rocks (fig. 8). The formation is cut by andesitic dikes that locally are sufficiently numerous to constitute dike swarms and by a few sandstone dikes. Parts of the formation are a series of open folds that commonly plunge gently northwestward.

The Root Glacier Formation disconformably overlies the Nizina Mountain Formation or, locally, the Lubbe Creek Formation (figs. 4, 5, 6). It is overlain unconformably by the Wrangell Lava and in some places is mantled by Quaternary surficial deposits. The conglomerate forms an intraformational lens stratigraphically high in the formation (figs. 4, 5, 8). Both of its contacts are broadly gradational from conglomerate through sandstone to dominantly pelitic rocks. Snow and ice cover parts of the conglomerate.

Accurate estimates of the thickness of the formation are precluded because of the angular unconformity that marks the upper contact, because of the lack of persistent marker beds, and because of numerous folds and faults. The large disparities in estimated thicknesses probably reflect fairly abrupt changes in thickness attributable to such factors as unequal erosion and differences in original thickness. The estimated thicknesses are probably reasonable minimum values. They range from 1,300 to about 4,000 feet. The conglomerate lens is about 200 feet in maximum thickness.

Petrology and petrography

The formation consists mainly of diverse clastic rocks, chiefly mudstone and siltstone and less abundant graywacke, arenite, and shale. It also contains some limy beds, lenses, and concretions. The conglomerate unit comprises well-indurated pebble and cobble conglomerate and coarse-grained and very coarse grained arenite. Reworked shaly chips are constituents of a few of the dominantly pelitic rocks. Most rocks of the formation are dark greenish gray or dark gray where fresh and diverse shades of brown where weathered.

The mudstone, siltstone, and shale contain similar mineral assemblages. They are closely related rocks and are distinguished by the sizes of their clasts or, in the case of the shale, by the development of fissility and pencil structure. Some of the very fine grained rocks contain carbonaceous trash, fossil wood, minute calcitic or chalcedonic spherical remnants of microfossils, and calcite veinlets. The shales and some of the siltstones are finely laminated because of the preferred



FIGURE 8.—Conglomerate (Irc) of the Root Glacier Formation east of McCarthy Creek. The conglomerate is about 200 feet in maximum thickness. A downfaulted segment of conglomerate is in the left foreground. The Root Glacier Formation (1/) is cut by several dikes. The Wrangell Lava (QTw) dominates the high terrain in the background. orientation of their platy minerals or because of the parallelism of trains of carbonaceous material or other rock-forming elements. The mudstone, siltstone, and shale consist of clasts of very fine grained quartz, calcite, and plagioclase, cemented by calcite, along with chlorite and clay minerals. Less common minerals in these rocks are chalcedony, sericite, biotite, illite, and opaque minerals, chiefly pyrite and secondary hydrous iron oxides. The rare detrital constituents are apatite, zircon, and epidote. A few of the mudstones contain moderate quantities of the zeolite, laumontite (leonhardite).

The graywacke and arenite commonly are very fine grained to medium-grained rocks. The graywacke is feldspathic. It characteristically is poorly sorted and is composed of subangular clasts in a calcite-chlorite-clay matrix that constitutes 10–30 percent of the rock. The clasts include quartz and plagioclase and less commonly clinopyroxene, lithic fragments, biotite, hornblende, calcite, K-feldspar, and opaque minerals. The matrix consists chiefly of calcite and chlorite and subordinate amounts of sericite, chalcedony, and clay minerals. The arenite resembles the graywacke, but it either is cemented by calcite or contains less than 10 percent matrix.

The few limy rocks in the formation are chiefly impure calcarenites and represent a phase of sedimentation marked by the dominance of calcite, both in the detrital fraction and in the matrix or cementing material. The concretions form spherical masses $\frac{1}{2}-\frac{11}{2}$ feet in diameter that are composed of dense, very fine grained calcite.

The conglomerates contain well-rounded pebbles, cobbles as much as 6 inches in diameter, and a few blocky intraformational fragments of mudstone and siltstone, all in a sandy matrix. About two-thirds of the cobbles and pebbles are limestone, probably derived from the Chitistone and Nizina Limestones. The other cobbles and pebbles include medium-grained granite that is rich in pink K-feldspar, chert, altered basalt derived from the Nikolai Greenstone, and quartz.

Except for a few extraneous lithic pebbles and granules, the very coarse grained and coarse-grained arenite consists of subangular clasts of quartz and plagioclase that are cemented by calcite, along with minor chlorite and clay minerals. It also carries minor amounts of opaque iron minerals and their alteration products. Fragments of wood are widely scattered in some of the sandstone.

Age

Paleontologic studies by R. W. Imlay indicate that the Root Glacier Formation is Late Jurassic or, more specifically, late Oxfordian and Kimmeridgian in age. Its main fossils include pelecypods of the genus *Buchia* [formerly *Aucella*], belemnites as much as 12 cm long, ammonites, and wood fragments. Fossils identified by Imlay include:

Lima sp., Camptonectes sp., Ctenostreon sp., Buchia rugosa (Fischer), B. concentrica (Sowerby), B. mosquensis (von Buch), Partschiceras sp., Amoeboceras (Prionodoceras) sp., "Turbo" sp., and Cylindroteuthis sp. This assemblage is similar to that of the Naknek Formation, a widely distributed Upper Jurassic unit of the Alaska Peninsula, Cook Inlet region, and Talkeetna Mountains (Martin, 1926, p. 132, 133, 168–180, 203–218; Grantz, 1960a, b).

THE WEHUTTY FORMATION OF NORTH CAROLINA

By Robert M. Hernon²

The Wehutty Formation is named for the settlement of Wehutty, southwestern North Carolina, in the north-central part of the Isabella quadrangle, Tennessee and North Carolina.

The mostly poor roads leading northwest and southwest from Mt. Olive church, northeast of the center of the Isabella quadrangle, provide the best and least interrupted exposures of the formation in the Wehutty area. The old State Road through Postell, N.C., has fair exposures of part of the Wehutty, and interrupted exposures may be seen along the tributary Allen Branch road and nearby lumber roads. The upper beds are well exposed along the wagon road of upper Persimmon Creek, southeast of Wolfpen Gap, Isabella quadrangle. The cuts along U.S. Highway 64 and minor tributary roads, southeast of Franklin Gap, Isabella quadrangle, also expose the upper strata.

The Wehutty Formation underlies an irregular area that constitutes above one-third of the Isabella quadrangle along its east boundary and also underlies approximately the northwestern third of the adjoining Persimmon Creek quadrangle. The formation extends an unknown distance northeast and southwest of the area studied (Hernon, 1964).

The Wehutty Formation was formerly included in the Copperhill Formation defined by Hurst (1955, p. 9) on the basis of studies in the Mineral Bluff quadrangle, Georgia. He delimited the Copperhill Formation in the northwestern part of the Mineral Bluff quadrangle and regarded the dominantly schistose unit along the southeast margin of its outcrop area as transitional to his Hughes Gap Formation. This schist zone is narrow in the Mineral Bluff quadrangle, but widens rapidly northeastward where it underlies about one-third each of the Isabella and Persimmon Creek quadrangles. Because of the great area extent in these quadrangles and the lithologic differences originally described by Hurst, this schist unit is here separated from the Copperhill Formation and named Wehutty Formation. The transi-

 $^{^2\,{\}rm Prepared}$ by Frank S. Simons from rough draft left by Hernon at the time of his death in June 1965.

tional beds of Hurst constitute the upper part of the Wehutty but apparently much of the lower part is cut out in the southern part of the Isabella quadrangle, probably by faulting, and does not extend into the Mineral Bluff quadrangle. Hurst (1955, p. 8-9) considered the Wehutty Formation as part of the Great Smoky Group of the late Precambrian Ocoee Series (King and others, 1958, p. 951).

STRATIGRAPHIC RELATIONS AND THICKNESS

The base of the Wehutty Formation is arbitrarily placed where the available information indicates change from the predominant metasandstone of the Copperhill Formation of Hurst (1955, p. 9) to an alternating sequence in which argillaceous and silty beds are abundant. Other workers would undoubtedly place the contact somewhat higher or lower, but the difference would probably be minor.

The upper contact is clear cut and is placed at the top of a conglomeratic member that can be traced with confidence. This contact corresponds to the base of the Hughes Gap Formation of Hurst (1955, p. 21–35) and marks a change from a sequence to the northwest having dominantly dark schists, locally graphitic, and phyllites to a sequence to the southeast characterized by light-colored sericitic and muscovitic schists.

Folding and probable faulting of this generally incompetent formation make estimates of its thickness uncertain. Its minimum thickness probably exceeds 4,000 feet; the total may be as much as 100 percent greater.

LITHOLOGY

The Wehutty Formation is composed of alternating beds of schist, ranging from fine- to medium-grained phyllitic varieties, and metasandstones, ranging from impure quartzite to metagraywacke. The arenaceous rocks may be pebbly and have thin conglomerate layers. Quartz-pebble conglomerates are conspicuously more common and better sorted in the upper 3,000 feet or so of the formation. The proportion of originally argillaceous and silty sediment is generally about half or more, but some zones hundreds of feet thick are composed largely of metamorphosed sand and gravelly sand.

SCHIST

The Wehutty Formation is distinguished from the underlying Copperhill Formation by a much greater abundance of schist and from the overlying similar Hughes Gap Formation by a greater abundance of light- to dark-gray phyllites and schists.

The schists range in color from light silvery gray to nearly black, and from typical schist to beds that could be described as glossy

phyllite, usually with visible porphyroblasts. Schistosity is usually conspicuous, but in some areas and beds the orientation of mineral grains is much less evident macroscopically, particularly where the rock is crowded with garnet or with micaceous minerals in cross position. Some of the dark graphitic layers are slaty in appearance except for warping of the cleavage. Deformation of the schistosity ranges from slight warping to acute crinkling, in places in larger folds.

The minerals of the gray schists are biotite, sericite, plagioclase, quartz, garnet, chlorite, tourmaline, apatite, zircon, opaque matter, and staurolite. Biotite in many beds is of two sizes: small flakes oriented along the schistosity and much larger books in cross position. Biotite is typically brown, but a few small flakes are olive green. Muscovite or sericite is abundant and with biotite defines the schistosity. Plagioclase is a major constituent in some laminae and also occurs infrequently as grains in quartz-rich laminae. It ranges from andesine to sodic oligoclase, but sodic oligoclase is most common. In the thin sections studied, quartz is much more abundant than plagioclase. Garnet is commonly in euhedral crystals with a low percentage of inclusions. Some chlorite is a pale-greenish low-birefringent variety, commonly intimately associated with biotite. A second type of chlorite, in part dusty, is pleochroic from olive or brownish green to colorless or to pale tan or green. The index of its cleavage flakes is near 1.62, and the mineral may be either positive or negative. This type of chlorite is associated closely with staurolite as large masses. It also forms smaller masses, perpendicular to schistosity, that may contain rare uniformly oriented remnants of biotite. Green chlorite, possibly original, forms prominent porphyroblasts in cross position in some biotitic schists.

Tourmaline is in rare olive-green or olive-brown grains. Zircon is in small grains and presumably is the mineral at centers of pleochroic halos in biotite and chlorite. The opaque matter includes iron sulfide, commonly pyrrhotite, dark-gray tabular masses of faintly magnetic iron ore believed to be ilmenite, and dusty to somewhat platy material, part of which is graphite or carbonaceous matter, and the rest of which is unidentified.

Euhedral staurolite is in schist layers an inch to a few inches thick that are interlayered with other schists in the middle and upper part of the formation. Some of the schist is thinly but markedly banded and resembles a unit described by Hurst (1955, p. 17) as varved schist. The staurolite crystals range in length from a fraction of an inch to perhaps 3 inches. Twins are not abundant, and most single crystals are well formed, particularly in finer grained rocks. Staurolite is poikiloblastic in varying degrees. Single crystals may show nearly clear parts and other parts thickly set with inclusions and dark opaque dust, probably graphite. These phases have no constant relation to borders of well-formed crystals and seem to represent differences in replaceability of two or more varieties of schist. Other crystals show rather evenly spaced inclusions throughout, mainly of quartz. Alteration of staurolite to sericite or muscovite is apparently rare except in thin zones along the margins of crystals. The sericite of such zones is in part oriented at high angles to the crystal faces.

METASANDSTONE

The most characteristic metasandstone is micaceous feldspathic quartzite in which biotite predominates over muscovite. Muscovite may nearly equal biotite in some layers but ordinarily is much less abundant and may be present in only trace amounts. The feldspar is sodic oligoclase to medium andesine. Much of it is untwinned and may not show cleavage; a few grains show zoning. In the typical micaceous quartzite, feldspar composes less than 25 percent. Mica, mainly biotite, may be as low as 5 percent, and carbonate commonly ranges from 1 to 8 percent. Other minerals are tournaline, zircon, sphene, garnet, leucoxene, chlorite, and apatite, in small or trace amounts. Potash feldspar was not found despite careful search. Radioactive material of low birefringence is in part metamict zircon, but other grains may be metamict monazite.

Typical feldspathic quartzite varies toward orthoquartzite by decrease in feldspar or toward graywacke by increase in mica but without showing any content of rock fragments. Thin conglomerate and scattered granules and small pebbles of quartz are other variations of the quartzite.

Some beds of deeply weathered metasandstone are apparently metamorphosed graywacke and arkose like those in the Copperhill Formation. These seem to be scarce, however.

METACONGLOMERATE

Well sorted quartz-pebble conglomerates are common in the upper part of the formation. The pebbles of these conglomerates are rather evenly sized and well rounded, but are flattened by flowage. The matrix shows a large proportion of quartz and lacks the rock fragments of metagraywackes. These conglomerates are interbedded with schists and with feldspathic mica quartzites. Similar conglomerates are also characteristic of, but more abundant in, the overlying Hughes Gap Formation.

Conglomerates in the lower half of the Wehutty commonly show a larger proportion of feldspar pebbles, continuous gradation in size of fragment from small pebbles to sand, a lesser degree of rounding, and much less persistence in thickness and extent of individual beds than do the quartz-pebble conglomerates; these conglomerates resemble those of the underlying Copperhill Formation.

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