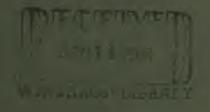
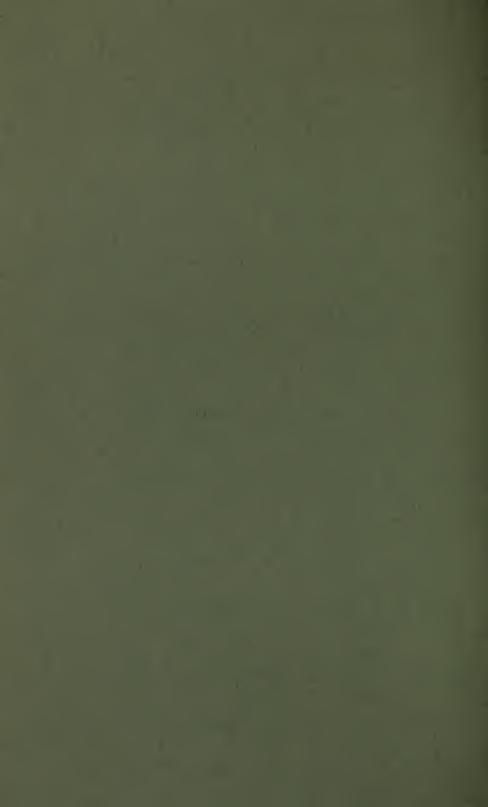
Changes in Stratigraphic Nomenclature by the U.S. Geological Survey 1964

GEOLOGICAL SURVEY BULLETIN 1224-A







Changes in Stratigraphic Nomenclature by the U.S. Geological Survey 1964

By GEORGE V. COHEE and WALTER S. WEST

CONTRIBUTIONS TO STRATIGRAPHY

GEOLOGICAL SURVEY BULLETIN 1224-A



UNITED STATES DEPARTMENT OF THE INTERIOR STEWART L. UDALL, Secretary

GEOLOGICAL SURVEY

William T. Pecora, Director

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CHANGES IN STRATIGRAPHIC NOMENCLATURE BY THE U.S. GEOLOGICAL SURVEY, 1964

By GEORGE V. COHEE and WALTER S. WEST

LISTINGS OF NOMENCLATURAL CHANGES

In the following listings, the changes in stratigraphic nomenclature are grouped together in the categories of (1) new names adopted, (2) previously used names adopted, (3) names revised, (4) changes in age designation, (5) names reinstated, and (6) names abandoned. The stratigraphic names involved in change are listed alphabetically under each category. The age of the unit, the area in which the name is employed, the title of the report, and the publication in which the change is described are given.

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Year of	publication				1964 [1965]	1964		1964	1965	1964	1964		1965	1964	1965
is adopted	Publication (U.S. Geol. Sur- vey except as indicated)	This report, p. A30	This report, p. A50	This report, p. A33	Bull. 1179	Map MF-282.	This report, p. A25	Map GQ-345	Bull, 1194-M	Am. Assoc. Petroleum Geol- ogists Bull., v. 48, no. 9.	do	This report, p. A23	Prof. Paper 477	Prof. Paper 475-D	Bull, 1194-K
Report in which new name is adopted	Title and authorship	Nomenclature and age of formations in the An- sonia quadrangle, Fairfield and New Haven	Counties, Connecticut, by C. E. Fritts. Paintbrush Tuff and Timber Mountain Tuff of Nye County, Nevada, by P. P. Orkild.	Belted Range Tuff of Nye and Lincoln Coun- ties, Nevada, by K. A. Sargent, D. C. Noble,	and E. B. Ekren. Geology of the Frenchie Creek quadrangle, north-central Nevada, by L. J. P. Muffler.	Reconnaissance bedrock geology of the Wabas- sus Lake quadrangle, Maine, by D. M. Lar-	rabee. Denny Creek Granodiorite Gueiss, Browns Pass Quartz Monzonite, and Kroenke Granodiorite Monnt Harvard duadrangle.	Colorado, by Fred Barker and M. R. Brock. Geologic map of the Prospect Peak quad-	rangue, Cannorma, Dy G. A. A. Anaxunau. Lower Mesozoic extrusive rocks in southeast- ern Arizona: the Canelo Hills Volcanics, by	F. T. Hayes, F. S. Simons, and R. S. Raup. New Permian stratigraphic units in the Glass Mountains, west Texas, by G. A. Cooper and P. F. Grant		Denny Creek Granodiorite Gneiss, Browns Pass Quartz Monzonite, and Kroenke	dranouorte, Mouur harvard quadrange, Colorado, by Fred Barker and M. R. Brock. Lake Bonneville: Quaternary stratigraphy of eastern Jordan Valley south of Salt Lake	Thirsty Canyon Tuff of Nye and Esmeralda Counties, Nevada, by D. C. Noble, R. E.	Anderson, B. B. Bkrein and J. T. O'Connor. The Foote Creek and Duton Creek Forma- tions, two new formations in the north part of the Laramie basin, Wyoming, by H. J.
	Location	Connecticut	Nevada	do	do	Maine	Colorado	California.	Arizona	Texas.	do	Colorado	Utah	Nevada.	Wyoming
	Age	Ordovician(?)	Pliocene	Miocene or Pliocene	Mesozoic	Devonian	Precambrian	Early Pleistocene	Triassic and Jurassic	Early Permian (Leonard).	do	Precambrian	Pleistocene	Pliocene	Paleocene
	Name	Allingtown Metadiabase	Ammonia Tanks Member (of Tim- ber Mountain Tuff) (of Piapi	Canyon Group). Belted Range Tuff	Big Pole Formation (of Pony Trail	Bottle Lake Quartz Monzonite	Browns Pass Quartz Monzonite	Burney Basalt	Canelo Hills Volcanics	Cathedral Mountain Formation	Decie Ranch Member (of Skinner	kanch Formation). Denny Creek Granodiorite Gneiss.	Draper Formation (of Lake Bonne- ville Group).	Dry Lake Member (of Thirsty Canvon Tuff).	Dutton Creek Formation

1965	1964	1964 1964	1964		1965	1005	COST			1964	1965			1965	1965		1964			1964		
do -	Bull. 1179	Prof. Paper 475-D	Map GQ-303	This report, p. A54	Bull. 1205.	tt 1011 []t		This report, p. A26		Prof. Paper 475-D.	Bull. 1194-0			Prof. Paper 477	Bull. 1194–N		Map GQ-358	This report, p. A31		Bull. 1194-B	This report, p. A53	This report, p. A55
Hyden, H. McAndrews, and R ¹ H Tschudy. do-	Geology of the Frenchie Creek quadrangle,	Thirstur-central revaua, by L.J.F. Anumer. Thirsty Canyon Tuff of Nye and Esmeralda Counties. Nevada, by D.C. Noble, R. E.	Anderson, E. B. Ekren, and J. T. O'Connor. Geology of the Tyrone quadrangle, Kentucky,	by E. R. Cressman. Marinette Quartz Diorite and Hoskin Lake	Granue of northeastern wisconsul, by W.C. Prinz, Geology of the Garns Mountain quadrangle,	Bonneville, Madison and Teton Counties, Idaho, by M. H. Staatzand H. F. Albee.	Alaska, by E. H. Lathram, J. S. Pomeroy,	Denny Creek Granodiorite Gneiss, Browns	Pass Quartz Monzonite, and Kroenke Granodiorite, Mount Harvard quadrangle,	Colorado, by Fred Barker and M. K. Brock. Thirsty Canyon Tuff of Nye and Esmeralda Commiss Nevela by D. C. Nohle, F. F.	Anderson, E. B. Ekren, and J. T. O'Connor. Stratizably and chronology of late interla-	cial and early Vashon glacial time in Seattle	H. H. Waldron, and Meyer Rubin.	Lake Bonneville: Quaternary stratigraphy of eastern Jordan Valley south of Salt Lake	Outy, Utan, by K. D. MULLION.	Forks Formations of southern Montana and	Bedrock geologic map of the Big Lake quad-	Nomenclature and age of formations in the	Ansonia quadrangle, raineld and New Haven Counties, Connecticut, by C. E.	Hovey Group, a redefined stratigraphic name for the Hovey Formation of northeast Maine,	by Louis Pavlides. Marinette Quartz Diorite and Hoskin Lake Granite of northeastern Wisconsin, by W.	O. Prinz. The Mashel Formation of southwestern Pierce County, Washington, by K. L. Walters.
do	Nevada	do	Kentucky	Wisconsin	Idaho		Alaska	Colorado		Nevada	Washington			Utah	Montana and Wyo-	•9mm	Maine	Connecticut		Maine	Wisconsin	Washington
Late Cretaceous and	Paleocene. Mesozoic	Pliocene	Middle Ordovician	Precambrian	middle Pliocene or	younger.	Faleocene Infough Miocene.	Precambrian		Pliocene	Plaistocana			do	Late Devonian		Devonian	Ordovician(?)		Silurian	Precambrian	Miocene
Foote Creck Formation	Frenchie Creek Rhyolite (of Pony	Trail Group). Gold Flat Member (of Thirsty Canvon Tuff).	Grier Limestone Member (of	Lexington Limestone). Hoskin Lake Granite	Kirkham Hollow Volcanics		Kootznahoo Formation	Kroenke Granodiorite		Labyrinth Canyon Member (of	I musty Canyon + unit.			Little Cottonwood Formation (of Lake Bonneville Group).	Logan Gulch Member (of the Three	FORKS FORMALIOU).	Love Ridge Quartz Monzonite	Malthy Lakes Volcanics		Maple Mountain Formation (of Hovey Group).	Marinette Quartz Diorite	Mashel Formation

CHANGES IN STRATIGRAPHIC NOMENCLATURE

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NEW NAMES AL

	Year of	publication	1965	1964				1964	1964 [1066]	1964	1964	1964		1964	1964 [1965]	
	is adopted	Publication (U.S. Geol. Sur- vey except as indicated)	Bull. 1194-F	Bull. 1194-B.	This report, p. A30	This report, p. A49	This report, p. A45	Map GQ-358	Bull. 1179	Am. Assoc. Petroleum Geol- ogists Bull., v. 48, no. 9.	Southeastern Geology, v. 5, no. 4.	Am. Assoc. Petroleum Geologists Bull., v. 48, no. 9	<u> </u>	Am. Assoc. Petroleum Geologists Bull., v. 48, no. 9.	Bull. 1179.	This report p. A27
NEW NAMES ADOF'ED FOR OFFICIAL USE IN U.S. GEOLOGICAL STREET	Report in which new name is adopted	Title and authorship	The Matagamon Sandstone: a new Devonian formation in north-central Maine, by D. W.	Rankin. Hovey Group, a redefined stratigraphic name for the Hovey Formation of northeast Maine,	by Louis Faviaces. Nomenclature and age of formations in the Ansonia quadrangle, Fairfield and New Hayen Counties, Connecticut, by C. E.	Fritts. Paintbrush Tuff and Timber Mountain Tuff of Nye County, Nevada, by P. P. Orkild.	do	Bedrock geologic map of the Big Lake quad-	Geology of the French by D. M. Larrave. Geology of the French Creek quadrangle,	New Permian stratigraphic units in the Glass Mountains, west Texas, by G. A. Cooper	The Pungo River Formation, a new name for middle Miocene phosphorites in Beaufort	. New Permian stratigraphic units in the Glass Mountay, be the contract of the	salue A. D. Wahnonie Formations of south- salver and Wahnonie Formations of south- eastern Nye County, Nevada, by F. G.	Poole, W. J. Carr, and D. P. Elston. New Permian stratigraphic units in the Glass Mountains, west Texas, by G. A. Cooper and P. R. Grant	Geology of the Frenchie Creek quadrangle,	South Pass Formation on the southwest flank of Wind River Mountains, Wyoming, by N. M. Denson, H. D. Zeller, and E. V. Stephens.
NT SOO TOTAL		Location	Maine	do	Connecticut	Nevada.	dodo	Maine.	Nevada	Texas	North Carolina	Texas	Nevada	Texas.	Nevada	Wyoming
N NON MALANA		Age	Early Devonian	Ordovician or Silurian.	Ordovician(?)	Miocene(?) and Pliocene.	do	Devonian	Mesozoic	Early Permian (Leonard).	middle Miocene	Early and Late Per- mian (Guadalupe).	late Miocene	Early Permian (Leonard).	Mesozoic	late Miocene to middle Pliocene.
T CATATON MAN		Name	Matagamon Sandstone	Nine Lake Formation (of Hovey Group).	Oronoque Member (of Derby Hill Schist).	Pah Canyon Member (of Paint- brush Tuff) (of Plapi Canyon	Paintbrush Tuff (of Piapi Canyon	Pocamoonshine Gabbro-Diorite	Pony Trail Group	Poplar Tank Member (of Skinner Ranch Formation).	Pungo River Formation	Road Canyon Member (of Word Formation).	Salyer Formation	Skinner Ranch Formation	Sod House Tuff (of Pony Trail	South Pass Formation

1965	1964	1964	1964	1965	1964	1964		1964	1964	1964 [1965]
Bull. 1181-R.	Am. Assoc. Petroleum Geologists Bull., v. 48, no 9	Prof. Paper 475-D	This report, p. A49 Prof. Paper 475-D	Bull. 1194-N	Map GQ-358.	Bull. 1164	This report, p. A40	Prof. Paper 501-D	Prof. Paper 501-B	Bull. 1194-D
Reconnaissance geology of Admiralty Island, Alaska, by E. H. Lathram, J. S. Pomeroy, U. C. Borr, and P. A. Lonov,	New Permian stratigraphic units in the Glass Mountains, west Texas, by G. A. Cooper and P. P. Grout	Thirsty Conyon Turf of Nye and Esmeralda Counties, Nevada, by D. C. Noble, R. E. Anderson, E. B. Ekren, and J. T.	P. Connor. P. Connor. Nye County, Nevada, by P. P. Orkida. Thisty Canyor Tutf of Nye and Bsmeaida. Counties. Nevada. by D. C. Noble, R. E.	Anderson, E. B. Ekren, and J. T. O'Connor. Nonexclature and correlation of lithologic subdivisions of the Jefferson and Three Forks Formations of southern Montana and	northern Wyoming, by C. A. Sandberg. Bedrock geologic map of the Big Lake quad-	Tengle, mane, by D. M. Latance. Tertiary geology of the Beaver Divide area, Fremont and Natrona Counties, Wyoming,	Salyer and Wahmonie Formations of south- eastern Nye County, Nevada, by F. G.	Three pre-Bull Lake tills in the Wind River Mountains, Wyoming, reinterpreted, by	Wildcat Valley Sandstone (Devonian) of southwest Virginia, by R. L. Miller, L. D.	The Witts Springs Formation of Morrow age in the Snowball quadrangle, north-central Arkansas, by B. B. Glick, S. B. Frezon, and Mackenzie Gordon, Jr.
Alaska	Texas	Nevada	do	Montana and Wyoming.	Maine	Wyoming	Nevada	Wyoming	Virginia	Arkansas
Late Jurassic and Early Cretaceous.	Early Permian (Leonard).	Pliocene	do	Late Devonian	Devonian	Eocene	late Miocene and early Pliocene.	or Till - Pleistocene	Early and Middle(?) Devonian.	Early Pennsylvanian (Morrow).
Stephens Passage Group	Sullivan Peak Member (of Skinner Ranch Formation).	Thirsty Canyon Tuff.	Timber Mountain Tuff (of Piapi Canyon Group). Trail Ridge Member (of Thirsty	Trident Member (of Three Forks Formation).	Wabassus Quartz Monzonite	Wagon Bed Formation	Wahmonie Formation	Washakie Point Glaciation or Till -	Wildcat Valley Sandstone	Witts Springs Formation

CHANGES IN STRATIGRAPHIC NOMENCLATURE A5

70		Year of publication	1964 [1965]	1964	1965	1964 [1965]	1965	1964 [1965]	1964	1965	1964	1964	1964	$1964 \\ [1965]$
JRVEY REPORT	is adopted	Publication (U.S. Geol. Survey except as indicated)	Prof. Paper 472	Bull. 1181–S	Map GQ-293	Prof. Paper 472	Map GQ-367	Prof. Paper 472	Jour. Palenotology, v. 38, no. 5.	Map GQ-293	Map GQ-303	Prof. Paper 501-C	do	Prof. Paper 472
FKEVIOUSLY USED NAMES ADOFTED FOR OFFICIAL USE IN U.S. GEOLOGICAL SURVEY REPORTS	Report in which name is adopted	Title and authorship	Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by	Geology of the Paradise quad- rangle, Cache County, Utah, by T. E. Mullens and G. A.	Geology of the Charters quad- rangle, Kentucky, by R. H.	Morris. Geology of the Omaha-Council B Uffs area, Nebraska-Iowa, by B D Millon	Geology of the Dunville quad- rangle, Kentucky, by C. H.	Geology of the Omaha-Council Bluffs area, Nebraska-Iowa,	A by t. D. thermet. A new species of the rugose coral genus Naliskinella from the Middle Devonian of eastern Pennsylvania, by W. A. Oliver,	Gology of the Charters quad- rangle, Kentucky, by R. H.	Geology of the Tyrone quad- rangle, Kentucky, by E. R.	Outline of the stratigraphic and tectonic features of northeast Maine, by Louis Pavlides, Fly Mencher, R. S. Naylor, and	A. J. Boucot.	Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.
FFICIAL USE		Original authorship	Newell, 1935	Williams, 1948	Foerste, 1917	Newell, 1935	Foerste, 1906	Jewett, 1932	Clarke, 1903	Linney, 1882	McFarlan and White, 1948.	Boucot, Field, Fletcher, Forbes, Naylor, and Pavlides,	1964. Gregory, 1899	Newell, 1935
PTED FOR O		Location	Ncbraska, Iowa, and Kansas.	Utah	Kentucky	Nebraska, Iowa, and Kansas.	Kentucky	Nebraska, Iowa, and Kansas.	New York	Kentucky	do	Maine	do	Nebraska, Iowa, and Kansas.
ED NAMES ADO		Age	Late Pennsylvanian	Late Devonian	Middle Silurian	Late Pennsylvanian	Middle Devonian	Late Pennsylvanian	Middle Devonian	Early and Middle Silurian.	Middle Ordovician	Early Devonian	do	Late Pennsylvanian.
SU YISUOUSLA US		Name	Argentine Limestone Member (of Wyandotte Limestone) (of Kansas	Beirdneau Sundsytone Member (of Jefferson Formation).	Bisher Limestone	Block Limestone Member (of Cherry- vale Formation) (of Kansas City Ground Ground Section	Boyle Limestone	Canville Limestone Member (of Dennis Limestone) (of Kansas	Cherry Valley Unwestone Member (of Marcellus Shale).	Crab Orchard Formation	Devils Hollow Member (of Cynthi- ana Formation).	Dockendorff Group	Edmunds Hill Andesite (of Dock-	Fontana Shale Member (of Cherry- vale Formation) (of Kansas City Group) (of Missouri Series).

PREVIOUSLY USED NAMES ADOPTED FOR OFFICIAL USE IN U.S. GEOLOGICAL SURVEY REPORTS

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1964 [1965]	1964	1965	1964 [1965]	1964		1964	1964 [1965]	1964	1964	1965	1965	1964	1964	
do	Map I-426	Map GQ-379	Bull. 1179	Prof. Paper 501-C		Bull. 1181-S	Prof. Paper 472	Map GQ-306	Bull. 1181–S	Bull. 1194-J	do	Map I-426	Bull. 1194-E.	
do	Geologic map of the Lucerne Valley quadrangle, San Bernar- dino. County, California, by	T. W. Dibblee, Jr. Geology of the Mount Aire quad- range, Utah, by M. D. Critten-	Geology of the Frenchie Creek quadrangle, north-central	Nevada, by L. J. F. Mumer. Outline of the stratigraphic and tectonic features of northeast	Mencher, R. S. Naylor, and	Geology of the Paradise quad- rangle, Cache County, Utah, by T. E. Mullens and G. A.	Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by	Geologic map of the Santa Rita quadrangle, New Mexico, by R. M. Hernon, W. R. Jones,	and S. L. Moore. Geology of the Paradise quad- rangle, Cache County, Utah, by T. E. Mullens and G. A.	Meeter, Meeter, and Loyd Sandstone Members of the Mancos Shale, Moffat and Rio Blanco Coun- tise Colondo Riv T D Duni	and H. L. Cullins.	Geologic map of the Lucerne Valley quadrangle, San Bernardino County, California.	by T. W. Dibblee, Jr. Upper Miocene and Pilocene marine stratigraphy in southern Salinas Vallev, California, by	D.,L. Durham and W. O. Addicott.
do	Vaughan, 1922	Thomas and Krueger, 1946.	Regnier, 1960	Gregory, 1900		Williams, 1948	Newell, 1935	Kuellmer and others, 1953.	Holland, 1952	Konishi, 1959	do	Shreve in Rich- mond, 1960.	Reed, 1925	
do	California	Utah	Nevada	Maine		Utah	Nebraska, Iowa, and Kansas.	New Mexico	Utah	Colorado	do	California	do	
do	Paleozoic	Late Triassic	middle Pliocene to middle Pleistocene.	Early Devonian		Late Devonian	Late Pennsylvanian.	Miocene(?)	Early Mississippian	Late Cretaceous	do	Tertiary	early Pliocene	
Frisbie Limestone Member (of Wyandotte Limestone) (of Kansas	City Group) (of Missouri Series). Furnace Limestone	Gartra Grit Member (of Ankareh Formation).	Hay Ranch Formation	Hedgehog Formation (of Docken- dorff Group).		Hyrum Dolomite Member (of Jeffer- son Formation).	Island Creek Shale Member (of Wyandotte Limestone) (of Kansas	Uny Group) (of Missouri Series). Kneeling Nun Rhyolite Tuff	Leatham Formation	Loyd Sandstone Member (of Man- cos Shale).	Meeker Sandstone Member (of	Mancos Shale). Old Woman Sandstone	Pancho Rico Formation	

CHANGES IN STRATIGRAPHIC NOMENCLATURE

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ED NAMES ADOPTED FOR OFFICIAL USE IN U.S. GEOLOGICAL SURVEY REPORTS-Continued	and the second se
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PREVIOUSLY	

	Year of publication	1964	1964 [1965] 1964	[1965] 1964 [1965]	1964	1964	1964	1964	1964 [1965]	1964	1964
is adopted	Publication (U.S. Geol. Survey except as indicated)	Prof. Paper 501-C	Prof. Paper 472do	Map I-412.	do	Map GQ-306	Map I-426	Map I-412	Prof. Paper 472	Map GQ-306	Prof. Paper 501-C
Report in which name is adopted	Title and authorship	Outline of the stratigraphic and tectonic features of northeast Maine, by Louis Paylides, Ely Mencher, R. S. Naylor, and	A. J. Boucot. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller. do	Geologic map and sections, Deep Creek area, Stevens and Pend Oreille Counties, Washington,	by R. G. Yates. do	Geologic map of the Santa Rita quadrangle, New Mexico, by R. M. Hernon, W. R. Jones,	and S. L. MOORE. Geologic map of the Lucerne Valley quadrangle, San Bernar- dino County, California, by T.	W. DUDDUEG, JT. Geologic map and sections, Deep Creek area, Stevens and Pend Oreille Counties, Washington,	by K. G. Yates. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by	k. D. Muller Geologic map of the Santa Rita quadrangle, New Mexico, by R. M. Hernon, W. R. Jones,	and S. L., Moore. Outline of stratigraphic and tectonic leatures of northeast Maine, by Louis Pavildes, Ely Mencher, R. S. Naylor, and A. J. Bouoct.
	Original authorship	Boucot, Field, Fletcher, Forbes, Naylor, and Pavlides,	1964. Newell, 1935	Fyles and Hew- lett, 1959.	Frebold and	Little, 1962. Hernon, Jones, and Moore, 1953.	Vaughan, 1922	Daly, 1912	Jewett, 1932	Kuellmer and others, 1953.	Fletcher, 1960
	Location	Maine	Nebraska, Iowa, and Kansas. do	Washington	do	New Mexico	California	Washington	Nebraska, Iowa, and Kansas.	New Mexico	Maine
	Age	Late Silurian	Late Pennsylvanian	Cambrian	Jurassic	Miocene (?)	Paleozoic	Tertiary	Late Pennsylvanian	Miocene(?)	Early Devonian
	Name	Perham Formation	Quindaro Shale Member (of Wyan- dotte Limestone) (of Kanasa City Group) (of Missouri Series). Ouivira Shale Member (of Cherv-	vale Formation) (of Kansas City Group) (of Missouri Series). Reeves Linnestone Member (of Mait- len Phyllite).	Rossland Group	Rubio Peak Formation	Saragossa Quartzite	Sheppard Granite	Stark Shale Member (of Dennis Limestone) (of Kansas City	Group) (of Missouri Series). Sugarlump Tuffs	Swanback Formation (of Docken- dorff Group).

1964 [1965] 1964 [1965] 1964				
Prof. Paper 472 dodo				
Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller. 				
Newell, 1935 Bain, 1898				
Nebraska, Jowa, and Kansas. do				
Late Pennsylvanian dododo				
Wea Shale Member (of Cherryvale Formation) (of Kansas City Group) (of Missouri Series). Westerville Linestone Member (of Cherryvale Formation) (of Kansas City Group) (of Missouri Series). Willimantic Gneiss.				

CHANGES IN STRATIGRAPHIC NOMENCLATURE A9

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				Report in which usage is revised	ge is revised	
Name	Age	Location	Revision	Title and authorship	Publication (U.S. Geol. Survey except as indicated)	Year of publication
Alpine Member (of Little Cottonwood Formation) (of Lake Bonneville Group).	Pleistocene	Utah	Formerly Alpine Formation of Lake Bonneville Group in Jordan Valley. Alpine Formation in good usage	Lake Bonneville: Quaternary stratigraphy of eastern Jordan Valley south of Salt Lake City, Utah, by R. B.	Prof. Paper 477	1965
Benton Shale	Early and Late Cre- taceous.	Colorado	eisewhere in Utah. Mowry Shale Member in- cluded in Benton Shale in	Morrison. Geology of northwestern North Park, Colorado, by	Bull. 1188	1965
Birdbear Member (of Jeffer- son Formation).	Late Devonian	Montana	Birdbear made member of Jefferson Formation in report area. Birdbear For- mation in good usage else-	Bedrock geology of the Saw- tooth Ridge quadrangle, Montana, by M. R. Mudge.	Map GQ-381.	1965
Bonneville Member (of Little Cottonwood For- mation) (of Lake Bonne- ville Group).	Pleistocene	Utah	where Formerly Bonneville Forma- tion of Lake Bonneville Group in Jordan Valley. Bonneville Formation in good usage elsewhere in	Lake Bonneville: Quaternary stratigraphy of eastern Jordan Valley south of Salt Lake City, Utah, by R. B. Morrison.	Prof. Paper 477	1965
Boulder River Sandstone Member (of Frontier For- mation).	Late Cretaceous	Montana	Lean. Reassignment from member of Colorado Shale to mem- ber of Frontier Formation.	Correlation of Cretaceous and lower Tertiary rocks near Livingston, Montana, by	Prof. Paper 525-B	1965
Brothers Volcanics (of Stephens Passage Group).	Late Jurassic and Early Cretaceous.	Alaska	Brothers Volcanics assigned to Stephens Passage Group.	A. E. KODERS. Reconnaissance geology of Admiralty Island, Alaska, by E. H. Lathram, J. S. Pomeroy, H. C. Berg, and	Bull. 1181–R.	1965
Buffalo Wallow Formation	Late Mississippian	Kentucky	Vienna Limestone made a member of the Buffalo Wallow Formation in	K. A. JOURY. Geology of the Rome quad- rangle, Kentucky, by M. D. Crittenden, Jr., and R. K.	Map GQ-362	1962
Cane Hill Formation	Mississippian and Early Pennsylvanian.	Arkansas	kentucky. Formery Cane Hill Member of the Hale Formation in report area. Age changed from Pennsylvanlan to Mississippian and Early Pennsylvanlan.	The Witts Springs Formation of Morrow age in the Snowball quadrangle, north-central Arkanass, by B. E. Glick, S. E. Freon, and Mackenzie Gordon, Jr.	Bull. 1194-D.	1964

1964	1964	1964	1964	1964	1964	196 4 [1965]		1965	1964	1965
Prof. Paper 501-C	Prof. Paper 475-D	do	Am. Jour. Science, v. 262, no. 5.	Map GQ-303	Basic Data Report 15	Prof. Paper 472	This report, p. A30	Water Supply Paper 1730-A.	Bull. 1194-B	Map GQ-381
Outline of the stratigraphic and tectonic features of northeast Maine, by Louis Parides, Ely Mencher, B., S. Naylor and A. J.	Surface and subsurface stratigraphic sequence in southeastern Mississippi,	ру D. H. Łargle. do	Reaction between mafic magma and pelitic schist, Cortlandt, New York, by	Fred Barker. Geology of the Tyrone quad- rangle, Kentucky, by E. R.	Records of Wells and test holes, water snalyses, and physical properties of water-bearing materials for the Denver basin, Colorado by J. A. McConaghy, G. H. Chase, A. J. Boettcher, and T. J.	Geology of the Ornaha-Coun- cil Bluffs area, Nebraska-	10wa, py k. D. Anuer. Nomenclature and age of for- mations in the Ansonia quadrangle, Fairfield and New Haven Counties, Con-	nection, by C. B. Fritis, Ground and surface water in the Mesabi and Vermilion fron Range area, north- eastern Minnesot, by R. D. Cotter, H. L. Young, L. R. Petri, and C. H.	Frior. Hovey Group, a redefined stratigraphic name for the Hovey Formation of north- east Maine, by Louis	ravitaes. Bedrock geology of the Saw- tooth Ridge quadrangle. Montana, by M. R. Mudge.
Chapman Sandstone included in Dockendorff Group.	Formerly Clayton Formation, which remains in good usage outside of report area.	Formerly Cook Mountain Formation, which remains in good usage outside of	Formerly Cortlandt Series	Formerly Curdsville Lime- stone of Lexington Group.	Dawson Formation in report area. Dawson Arkose in good usage elsewhere.	Winterset Linestone Member placed in Dennis Limestone.	Oronoque Member made a member of the Derby Hill Schist.	Formerly Duluth Gabbro	Formerly Dunn Brook Mem- ber (Early Silurian age) of Hovey Formation.	Formerly Flood Member of Blackleaf Formation.
Maine	Mississippi	do	New York	Kentucky	Colorado	Nebraska and Iowa.	Connecticut	Minnesota	Maine	Montana
Early Devonian	Paleocene	middle Eocene	unknown	Middle Ordovician	Late Cretaceous and Paleocene.	Late Pennsylvanian	Ordovician(?)	Late Precambrian	Ordovician or Sil- urian.	Early Cretaceous
Chapman Sandstone (of Dockendorff Group).	Clayton Limestone (of Midway Group).	Cook Mountain Limestone (of Claiborne Group).	Cortlandt Complex	Curdsville Limestone Member (of Lexington	Dawson Formation	Dennis Limestone (of Kan- sas City Group) (of Mis-	sourt series). Derby Hill Schist	Duluth Gabbro Complex	Dunn'Brook Formation (of Hovey Group).	Flood Shale Member (of Blackleaf Formation) (of Colorado Group).

CHANGES IN STRATIGRAPHIC NOMENCLATURE

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	Year of publication		1964	1964	1964	1965	1965	1964	1964	1964
ige is revised	Publication (U.S. Geol. Survey except as indicated)	This report, p.A36	Am. Assoc. Petro- leum Geologists Bull., v. 48, no. 9.	Bull. 1194-B	Bull. 181-S	Map GQ-381	Prof. Paper 477	Map GQ-303	do	Map I-412
Report in which usage is revised	Title and authorship	Belted Range Tuff of Nye and Lincoin Counties, Ne- vada, by K. A. Sargent, D. C. Noble, and E. B. Ekren.	New Permian stratigraphic units in the Glass Moun- tains, west Peras, by G. A. Cooner and P. E. Grant.	HOVEN GROUP, a redefined stratigraphic name for the HOVEN Formation of north- east Maine, by Louis	Pavildes. Pavildes. Geology of the Paradise quadrangle, Cache County, Utah, by T. E. Mullens and G. A. Izett.	Bedrock geology of the Saw- tooth Ridge quadrangle, Montana, by M. R.	Mudge. Lake Bonneville: Quaternary stratigraphy of eastern Jordan Valley south of Salt Lake City, Utah, by R.	B. Morrison. Geology of the Tyrone quad- rangle, Kentucky, by E. R. Cressman.	do	Geologic map and sections, Deep Creek area, Stevens and Pend Oreillo Countics,
	Revision	Member of Belted Range Tuff except in eastern part of Nevada Test Site where it remains a member of the Indian Trail Formation of Oak Spring Group. Re- defined to include only the	upper part of tue origination unit throughout its extent. Formerty Hess Linnestone Member of Leonard Formation.	Formerly Hovey Formation of Early Silurian (?) and Early Silurian age.	Jefferson Formation divided into the Hyrum Dolomite Member (at bottom) and the Beirdneau Sandstone Member (at top) in report	Formation includes Birdbear Member in report area.	Lake Bonneville Group redefined to include Draper Formation (top) and Little Cortenwood Formation	(bottom). Formerly Lexington Group	Formerly Logana Formation	of Lexington Group. Reeves Limestone Member made a member of Maitlen Phylite.
	Location	Nevada	Texas	Maine	Utah	Montana	Utah	Kentucky	do	Washington
	Age	Miocene or Pliocene	Early Permian (Leonard).	Ordovician or Silurian.	Late Devonian	do	Pleistocene	Middle Ordovician	do	Early or Middle Cambrian.
	Name	Grouse Canyon Member (of Belted Range Tuff),	Hess Formation	Hovey Group	Jefferson Formation	Do	Lake Bonneville Group	Lexington Limestone	Torana Member (of	Lexington Limestone). Maitlen Phyllite

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Bull. 1194-J	Bull. 1188	Map GQ-282	This report, p. A44	Prof. Paper 477	Prof. Paper 525-B	Map I-417	This report, p. A50	Map GQ-282	do	Bull. 11 94 -N
Washington, by R. G. Yates, and Loyd Sandstone Meeker and Loyd Sandstone Members of the Mancos Blanco Counties, Colorado, Dy J. R. Dyni and M. L.	Cullins. Geology of northwestern North Park, Colorado, by	W. J. Hall, Jr. Geology of the Shopville quadrangle, Kentucky, by N. L. Hatch, Jr.	Paintbrush Tuff and Timber Mountain Tuff of Nye County, Nevada, by P. P.	Urkud. Lake Bonneville: Quater- nary stratigraphy of eastern Jordan Valley south of Salt Lake City, Utah, by R. B.	Morrison. Correlation of Cretaceous and lower Tertiary rocks near Livingston, Montana,	by A. E. Roberts. Preliminary geologic map of the Tepee Creek quad- rangle, Montana-Wvoming.	by I. J. Witkind. Paintbrush Tuff and Timber Mountain Tuff of Nye County, Nevada, by P. P. Orkild.	Geology of the Shopville quadrangle, Kentucky, by N. L. Hatch, Jr.	do	Nomenclature and correlation of lithologic subdivisions of the Jefferson and Three Forks Formations of south- ern Montana and northern Wyoming, by C. A. Sandberg.
Meeker Sandstone Member and Loyd Sandstone Mem- ber made members of Man- cos Shale in report area.	Mowry Shale made a mem- ber of Benton Shale in	report area. Ste. Genevieve and St. Louis Limestones made members of Newman Limestone in	report area Formerly Piapi Canyon Formation of Oak Spring Group. Age was early	r nocene or younger. Formerly Provo Formation of Lake Bonneville Group in Jordan Valley. Provo Formation in good usage	elsewhere in Utah. Reassignment—is a member ber of the Kootenai Forma- tion as well as of the Clov-	erly Formation. Quadrant Sandstone used in report area. Quadrant Formation or Quartzite in	good usage elsewhere. Formerly Rainier Mesa Mem- ber of Piapi Canyon For- mation of Oak Spring Group. Age was Plicene	or younger. Made member of Newman Limestone in report area. Ste. Genevieve Limestone	in good usage elsewhere. Made member of Newman Limestone in report area. St. Touis Limestone in	good usage elsewhere. Formerly Saphington Sand- stone Member of Three Forks Formation.
Colorado	do	Kentucky	Nevada	Utah	Montana	Montana and Wyoming.	Nevada	Kentucky	do	Montana and Wyoming.
Late Cretaceous	Early Creataceous	Late Mississippian	Miocene(?) and Pliocene.	Pleistocene	Early Cretaceous	Pennsylvanian	Pliocene	Late Mississippian	-do	Late Devonian and Barly Mississip- pian.
Mancos Shale	Gr Mowry Shale Member (of Benton Shale).	G Newman Limestone	Piapi Canyon Group	Provo Member (of Little Cottonwood Formation) (of Lake Bonneville Group).	Pryor Conglomerate Mem- ber (of Kootenai Forma- tion).	Quadrant Sandstone	Rainier Mesa Member (of Timber Mountain Tuff) (of Piapi Canyon Group).	Ste. Genevieve Limestone Member (of Newman Limestone).	St. Louis Limestone Mem- ber (of Newman Limestone).	Sappington Member (of Three Foris Formation).

CHANGES IN STRATIGRAPHIC NOMENCLATURE A13

		Year of publication	1965	1964		1965	1964	1965		
	ge is revised	Publication (U.S. Geol. Survey except as indicated)	Bull. 1181–R.	Prof. Paper 475-D	This report, p. A45	Map GQ-381	Prof. Paper 475-D	Bull. 1194-N	This report, p. A45	This report, p. A49
nentininoo	Report in which usage is revised	Title and authorship	Reconnaissance geology of Admiralty Island, Alaska, by E. H. Lathram, J. S. Pomeroy, H. C. Berg, and	R. A. Loney. Thirsty Canyon Tuff of Nye and Esmeralda Counties, Neyada, by D. C. Noble, R. F. Anderson, F. F.	Ekren, and J. T. O'Connor. Paintburah Tuff and Timber Mountain Tuff of Nye County, Nevada, by P. P. Orkild.	Bedrock geology of the Saw- tooth Ridge quadrangle, Montana, by M. R.	Muqge. Surface and subsurface stratigraphic sequence in southeastern Mississippl,	by D. H. Bargta Nomenclature and correlation of lithologic subdryisions of the Jefferson and Three Forks Formations of southern Montana and	C. A. Sandberg. C. A. Sandberg. Paintbrush Tuff and Timber Mountain Tuff of Nye County, Nevada, by P. P. Orkild.	do
		Kevision	Seymour Canal Formation assigned to Stephens Pas- sage Group.	Formerly Spearhead Rhyo- lite of Pliocene(?) age.	Formerly Stockade Wash Member of Piapi Canyon Formation of Oak Spring Group. Former age was	Formerly Taff Hill Glau- contic Member of Black- leaf Formation.	Formerly Tallahatta For- mation. Tallahatta For- mation in good usage ex-	cept in area of report. Three Forks Formation di- vided into (ascending order): Logan Guich, Trident, and Sappington Members in report area.	Formerly Tiva Canyon Member of Plapi Canyon Formation of Oak Spring Group. Are formerly early	Plicence or younger. Formerly Topopabl Spring Member of Plapi Canyon Formation of Oak Spring Group. Age was Pliceene or younger.
	:	Location	Alaska.	Nevada	do	Montana	Mississippi	Montana and Wyoming.	Nevada	do
-		Age	Late Jurassic and Early Cretaceous.	Pliocene	Miocene(?) and Pliocene.	Early Cretaceous	middle Eocene	Late Devonian and Early Mississippian.	Miocene(?) and Pliocene.	do
		Лаше	Seymour Canal Formation (of Stephens Passage Group).	Spearhead Member (of Thirsty Canyon Tuff).	Stockade Wash Member (of Paintbrush Tuff) (of Piapi Canyon Group).	Taft Hill Member (of Blackleaf Formation) (of Colorado Group).	Tallahatta Siltstone (of Claiborne Group).	Three Forks Formation	Tiva Canyon Member (of Paintbrush Tuff) (of Piapi Canyon Group).	Topopah Spring Member (of Paintbrush Tuff) (of Piapi Canyon Group).

STRATIGRAPHIC NAMES REVISED-Continued

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	1965	1965	1965	1964	196 4 [1965]	1964	
This report, p. A34	Bull. 1194-0	Map GQ-381	Map GQ-362	Prof. Paper 475-D	Prof. Paper 472	Am. Assoc. Petro- leum Geologists Bull., v. 48, no. 9.	This report, p. A49
Belted Range Tuff of Nye and Lincoln Counties, Nevada, by K. A. Sargent, D. C. Noble, and E. B. Ekren.	Stratigraphy and chronology of last intergratiant and early Vashon glacial titme in Seattle area, Washington, by D. R. Mullineaux, H. H. Waldron, and Meyer	Bedrock geology of the Saw- tooth Ridge quadrangle, Montana, by M. R.	Geology of the Rome quad- rangle, Kentucky, by M. D. Crittenden, Jr., and	Surface and subsurface stratigraphic sequence in southeastern Mississippi,	Geology of the Omala- Council Bluffs area, Ne- braska-Iowa, by R. D. Miller.	New Permian stratigraphic units in the Glass Moun- tains, west Texas, by G. A. Cooper and R. E.	Paintbrush Tuff and Timber Mountain Tuff of Nye County, Nevada, by P. P. Orkild.
Member of Belted Range Tuff except in eastern part of Nevrada Test Site where it remains a member of the Indian Trail Formation of	Vashon Drift divided into Lawton Clay Member (at base) and Esperance Sand Member (at top).	Formerly Vaughn Bento- nitic Member of Blackleaf Formation.	Formerly Vienna Limestone	Formerly Winona Sand or Formation and remains as such outside of report area.	Formerly Wintersct Lime- stone of Pennsylvanian age.	Road Canyon Member replaces the "First Lime- stone" member of P. B. King (1931).	Formerly Yucca Mountain Member of Piapi Caryon Formation of Oak Spring Group. Age was early Pliocene or younger.
do	Washington	Montana	Kentucky	Mississippi	Nebraska and Iowa.	Texas	Nevada
Miocene or Pliocenedodo	Pleistocene	Early Cretaceous	Late Mississippian	middle Eocene	Late Pennsylvanian	Early and Late Permian (Guada- lupe).	Miocene (?) and Pliocene.
Tub Spring Member (of Belted Range Tuff).	Vashon Drift	Vaughn Member (of Blackleaf Formation) (of Colorado Group).	Vienna Limestone Member (of Buffalo Wallow Formation).	Winona Marl (of Clai- borne Group).	Winterset Limestone Member (of Dennis Limestone) (of Kansas City Group) (of Mis-	word Formation	Yucca Mountain Member (of Paintbrush Tuff) (of Plapi Canyon Group).

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ation is changed	Publication (U.S. Geol. Survey except as indicated)	This report, p. A63	This report, p. A32	Bull, 1183	Map GQ-293	do	Bull. 1183	Prof. Paper 472	Bull, 1183	Science, v. 145, no. 3633.	Prof. Paper 472
Report in which age designation is changed	· Title and authorship	Precambrian and Lower Cam- brian formations in the Last Chance Range area, Inyo County, California, by J. H.	Stewart. Nomenclature and age of forma- tions in the Ansonia quadran- gle, Fairfield and New Haven Connetient, by	C. E. Fritts. Provisional geologic map of the crystalline rocks of South Car- olina, by W. C. Overstreet and	Henry Bell, 3d. Geology of the Charters quad- rangle, Kentucky, by R. H. Morris.	dodo	Provisional geologic map of the crystalline rocks of South	Carouna, Dy W. C. Overstreet and Henry Bell, 3d. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	Provisional geologic map of the crystalline cocks of South Caroline Juy W. C. Overeraot	and Henry Bell, 3d. and Henry Bell, 3d. Potassium-argon and lead-alpha ages of some plutonic rocks, Bokan Mountain area, south- eastern Alaska, by M. A. Lan-	phere, B. M. MacKevett, Jr., and T. W. Stern. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.
	Location	California	Connecticut	North Carolina and South Carolina.	Kentucky, Ohio Michigan, Pennsyl- vania, and West	Virginia.	North Carolina and South Carolina.	Nebraska and Iowa	North Carolina and South Carolina.	Alaska	Nebraska and Iowa
Age	Former	Precambrian(?) and Early Cambrian.	Ordovician or Devonian.	Precambrian(?) or Paleozoic(?).	Mississippian	do	Precambrian	Pennsylvanian	Cambrian (?)	Cretaceous(?) and Tertiary(?).	Pennsylvanian
A	New	Precambrian and Early Cambrian,	Devonian	Ordovician to Mis- sissippian.	Devonian or Mississippian.	do	Ordovician to Mississippian.	Late Pennsylvanian.	Ordovician to Mississippian.	Late Triassic or Early Jurassic.	Late Pennsylvanian Pennsylvanian
	Name	Andrews Mountain Mem- ber (of Campito For- mation).	Ansonia Gneiss	Battleground Schist	Bedford Shale or Formation.	Berea Sandstone or	Bessemer Granite	Bethany Falls Limestone Member (of Swope Limestone) (of Kanass City Group) (of Missouri	Blacksburg Schist	Bokan Mountain Granite	Bonner Springs Shale (of Kansas City Group) (of Missouri Series).

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This report p. A63	Bull. 1181-R	Prof. Paper 472	Prof. Paper 501-C	Prof. Paper 472	Bull. 1201-A	Prof. Paper 472	Bull. 1183.	This report, p. A62	Bull. 1181-R	Bull. 1183	Prof. Paper 472	This report, p. A51
Precambrian and Lower Cam- brian formations in the Last Chance Range area, Inyo County, California, by J. H.	Recommanserate geology of Admiralty Island, Alaska, by E. H. Lathram, J. S. Pomeroy,	R. C. Derg, and K. A. Louey. Geology of the Omaha-Council Bluffs aree, Nebraska-Iowa, by R. D. Miller.	Chemistry of greenstones of the Catootin Formation in the Blue Ridge of central Virginia, by	J. C. Keeu, Jr. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa,	Petrography of evaporites from the Wellington Formation near Hutchinson, Kansas, by C. L.	Geology of the Omaha-Council Bluffs area, Nebraska-Iowa,	by K. D. Muler. Provisional geologie map of the crystalline rocks of South Caro- lina, by W. C. Overstreet and	Henry Bell, 3d. Precambrian and Lower Cam- brian formations in the Last Chance Range area, Inyo County, California, by J. H.	Stewart. Reconnaissance geology of Ad- miralty Island, Alaska, by E. H. Lathram, J. S. Pomeroy,	H. C. Berg, and R. A. Loney. Provisional geologic map of the crystalline rocks of South Caro- lina, by W. C. Overstreet and	Henry Bell, 3d. Geology of the Omaha-Council Bluffs area, by R. D. Miller.	Age of the Eleana Formation (Devonian and Mississipian) in the Nevada Test Site, by F. G. Poole, P. P. Orkild, Mackenzie Gordon, Jr., and Helen Duncan.
California	Alaska	Nebraska and Iowa	Virginia	Nebraska and Iowa	Kansas	Nebraska and Iowa	North Carolina and South Carolina.	California	Alaska	North Carolina and South Carolina.	Nebraska and Iowa	Nevada
Precambrian(?) and Early Cambrian.	Permian.	Pennsylvanian	late Precambrian	Pennsylvanian	Permian	Pennsylvanian	Devonian(?)	Precambrian(?)	Jurassic(?) to Early Cretaceous (?).	Cambrian	Pennsylvanian	Mississippian to Early Pennsyl- vanian.
Precambrian and Early Cambrian.	Early Permian	Late Pennsylvanian	late Precambrian(?) late Precambrian	Late Pennsylvanian	Early Permian	Late Pennsylvanian	Mississippian(?) to Permian(?).	Precambrian	Late Jurassic and Early Cretaceous.	Ordovician to Missis- sippian.	Late Pennsylvanian	Devonian and Miss- issippian.
Campito Formation	Cannery Formation	Captain Creek Limestone Member (of Stanton Limestone) (of Lansing Group) (of Missouri Series)	Catoctin Formation	Chanute Shale (of Kansas City Group) (of Missouri Series)	Chase Group	Cherryvale Formation (of Kansas City Group) (of Miscouri Sories)	Cherryville Quartz Mon- zonite.	Deep Spring Formation	Douglas Island Volcanics (of Stephens Passage Group).	Draytonville Conglomer- ate Member (of Kings Mountain Quartzite).	Drum Limestone (of Kansas City Group) (of Missouri Series)	Eleana Formation

CHANGES IN STRATIGRAPHIC NOMENCLATURE

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	Year of publication	1964	1964	1964 [1965]	1964 [1965]	1965	1964 [1965]	1965	1965	1964 [1965]	
ttion is changed	Publication (U.S. Geol. Survey except as indicated)	Bull. 1180-C	Prof. Paper 483-B	Prof. Paper 472	do	Bull. 1183.	Prof. Paper 472	Bull. 1183.	Prof. Paper 474-B	Prof. Paper 472	This report, p. A62
Report in which age designation is changed	Title and authorship	Middle and lower Ordovician for- mations in southermost Ne- vada and adjacent California,	by K. J. 4085 The Devonian colonial coral genus Billingustrate and its earliest known species, by W. A. Oliver,	Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	do	Provisional geologic map of the crystalline rocks of South Carolina, by W. C. Overstreet	Geology of the Omaha-Council Bluffs area, Nebraska-Iowa,	by K. D. Muler. Provisional geologic map of the crystalline rocks of South Carolina, by W. C. Over-	suces and menty bein, au Geology and uranium deposits of Elk Ridge and vicinity, San Juan County, Utah, by R. Q. Lewis, Sr., and R. H.	Composition Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	Precambrian and Lower Cam- brian formations in the Last Chance Range area, Inyo County, California, by J. H. Stewart.
	Location	Nevada and Califor- nia.	New York, Pennsyl- vania and New Jer- sey.	Nebraska and Iowa	do	North Carolina and South Carolina.	Nebraska and Iowa	North Carolina and South Carolina.	Utah	Nebraska and Iowa	California
Age	Former	Late Ordovician	Early or Middle De- vonian.	Pennsylvanian	do	Cambrian(?)	Pennsylvanian	Precambrian or early Paleozoic.	Middle Penn- sylvanian.	Pennsylvanian	Precambrian (?)
V	New	Middle and Late Ordovician.	Early Devonian	Late Pennsylvanian	do	Mississippian	Late Pennsylvanian	Ordovician to Devonian.	Middle and Late Pennsylvanian.	Late Pennsyl- vanian.	Precambrian
	Name	Ely Springs Dolomite	Esopus Shale, Grit, Silt- stone.	Eudora Shale Member (of Stanton Limestone) (of Lansing Group) (of Mis-	Farley Limestone Member (of Wyandotte Lime- stone) (of Kansas City Group) (of Missouri Se-	ries). Gaffney Marble	Galesburg Shale (of Kansas City Group) (of Mis-	souri Series). Henderson Gneiss	Hermosa Formation	Hickory Creek Shale Member (of Plattsburg Limestone) (of Lansing Group) (of Missouri	Series). Hines Tongue (of Reed Dolomite).

1965	1965		1964 [1965] 1064	[1965] 1965	1964 [1965] 1964	[1965] 1964	1964	1964 [1965]	1965	1965	
Bull. 1201-A	Prof. Paper 483-G	This report, p. A45	Prof. Paper 472	Bull. 1183	Prof. Paper 472 do	Prof. Paper 458-B	Prof. Paper 501-C	Prof. Paper 472	Prof. Paper 483-G	Bull. 1201-A	This report, p. A29
Petrography of evaporites from the Wellington Formation near Hutchinson, Kansas, by C. L.	Pountion and distribution of the genus Mys and Tertiary migrations of Mollusca, by F.	Paintbrush Tuff and Timber Mountain Tuff of Nye County,	Arevada, by F. F. Virtua. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	Provisional geologic map of the crystalline rocks of South Caro- # lina, by W. C. Overstreet and	Henry Bell, 3d. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	Rocks, structure, and geologic history of Steamboat Springs thermal area, Washoe County,	O. Revada, by D. E. White, G. A. Thompson, and C. H. Sandberg. Sandberg. In the stratigraphic and tectoria features of northeast Mathe, by Louis Perides, Ely Macher R. S. Navlor and A.	J. Boucot. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	Evolution and distribution of the genus Mya and Tertiary migrations of Mollusca, by F. S.	Mactiven. Petrography of evaporites from the Wellington Formation near Hutchinson, Kansas, by C.L.	South Pass Formation on the southwest flank of Wind River Moundains, Wyoming, by N. M. Denson, H. D. Zeller, and E. V. Stephens.
Kansas	California	Nevada	Nebraska and Iowa	North Carolina and South Carolina.	Nebraska and Iowa dodo	Nevada	Maine	Nebraska and Iowa	Alaska	Kansas	Eastern Wyoming and Nebraska
Permian	early Miocene	late Miocene or early Pliocene.	Pennsylvanian	Cambrian	Pennsylvaniandodo	late Pliocene or early Pleistocene.	Late Devonian	Pennsylvanian	M iocene	Permian	Pliocene
Early Permian (Leonard).	late Miocene or early Pliocene.	Miocene and Plio- cene (?).	Late Pennsylvanian. do	Ordovician to Missis- sippian.	Late Pennsylvanian. do	Pliocene (?) and Pleistocene.	early Middle Devo- nian.	Late Pennsylvanian	Oligocene or Miocene.	Early Permian (Leonard).	late Miocene and Pliocene.
Hutchinson Salt Member (of Wellington For- mation) (of Summer	Group). Imperial Formation	Indian Trail Formation.	Iola Limestone (of Kansas City Group) (of Missouri Series). Veness City, Group (of	Missouri Series). Kings Mountain Quartzite.	Lane Shale (of Kansas City Group) (of Missouri Series). Lanshæ Group (of Missouri	Series). Lousetown Formation	Mapleton Sandstone	Merriam Limestone Mem- ber (of Plattsburg Lime- stone) (of Lansing Group)	(of Missouri Series). Meshik Formation	Ninnescah Shale (of Sumner Group).	Ogallala Formation

CHANGES IN STRATIGRAPHIC NOMENCLATURE A19

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	Year of publication		1964		1965	1965	1964 [1965]	1964	1964 [1965]	1964 [1965]	1964 [1965]
tion is changed	Publication (U.S. Geol. Survey except as indicated)	This report, p. A32	Prof. Paper 501-C	This report, p. A62	Bull. 1181-R	Prof. Paper 474-B	Prof. Paper 472	Prof. Paper 483-B	Prof. Paper 472	do	do
Report in which age designation is changed	Title and authorship	Nomenclature and age of formations in the Ansonia quadrangle, Pairfield and New	by C. E. Fritts. Stratigraphic importance of corals in the Redwall Linestone.	Sando. Precambrian and Lower Cam- brian formations in the Last	County, California, hryo County, California, by J. H. Stewart. Reconnaissance geology of Ad- miratty Island, Alasska, by E. H. Lathram J. S. Pomerov.	H. C. Berg, and R. A. Loney. Geology and uranium deposits of Elk Ridge and vicinity, San Juan County, Utah, by R. Q.	Developments, St. and A. H. Camp- bell Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	The Devonian colonial coral genus Billingsastraea and its earliest known sneeles. by	W. A. Oliver, Jr. Geology of the Omaha-Council Bulffs area, Nebraska-Iowa, by R. D. Miller.	do	do
	Location	Connecticut	Arizona	California	Alaska	Utah, Arizona, and New Mexico.	Nebraska and Iowa	New York	Nebraska and Iowa	do	do
Age	Former	Ordovician or Devonian.	Mississippian	Precambrian (?)	Silurian and Devon- ian.	Pennsylvanian and Permian(?).	Pennsylvanian	Early or Middle Devonian.	Pennsylvanian	do	do
P	New	Devonian	Early and Late Mississippian.	Precambrian	Middle(?) Devonian	Permian. Permian.	Late Pennsylvanian	Early Devonian	Late Pennsylvanian	do	do
	Name	Prospect Gneiss	Redwall Limestone	Reed Dolomite	Retreat Group	Rico Formation	Rock Lake Shale Member (of Stanton Limestone) (of Lansing Group) (of	Missouri Series). Schoharie Grit	South Bend Limestone Mem- ber (of Stanton Limestone) of Lansing Group) (of	Missouri series). Spring Hill Limestone Mem- ber (of Flattsburg Lime- stone) (of Lansing Group)	(of Missouri Series). Stanton Limestone (of Lansing Group) (of Mis- souri Series).

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1964	1964 [1965]	1965	1964 [1965]	1965	1964	1964 [1965]	1965	1964	1965	1964 [1965]	1965
Prof. Paper 458-B	Prof. Paper 472	Bull. 1201-A	Prof. Paper 472	Map GQ-379	Prof. Paper 437-A	Prof. Paper 472	Bull. 1201-A	Bull. 1144-F	Bull. 1183	Prof. Paper 472	Bull. 1183
Rocks, structure, and geologic history of Steamboat Springs thermal area, Washoe County, Nevada, by D. E. White, G. A. Thompson, and C. H. Sand-	Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	Petrography of evaporites from the Wellington Formation near Hutchinson, Kansas, by C. L.	Geology of the Omaha-Council Blufts area, Nebraska-Iowa,	Geology of the Mount Aire quad- rangle, Utah, by M. D. Crit-	Alluvial fans and near-surface subsidence in western Fresno Country, California, by W. B.	Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by P. D. Milor.	Petrography of evaporites from the Wellington Formation near Hutchinson, Kansas, by C. L.	Distribution of thorium and uranium in three early Paleo- zoie plutonic series of New	Provisional geologic map of the crystalline rocks of South Carolina, by W. C. Overstreet	Geology of the Omaha-Council Buffs area, Nebraska-Iowa, by D. Mallor	Provisional geologic map of the crystalline rocks of South Carolina, by W. C. Overstreet and Henry Bell, 3d.
Ncvada	Nebraska and Iowa	Kansas	Nebraska and Iowa	Utah	California	Nebraska and Iowa	Kansas and Oklahoma.	New Hampshire	North Carolina and South Carolina.	Nebraska and Iowa	North Carolina and South Carolina.
Pliocene or Pleis- tocene.	Pennsylvanian	Permian	Pennsylvanian	Early Cambrian	Pliocenc and Pleis- tocene(?).	Pennsylvanian	Permian	do	late Carboniferous(?)	Pennsylvanian	Early Mississippian (?).
Pliocene or Pleis- tocene.	Late Pennsylvanian.	Early Permian (Leonard).	Late Pennsylvanian	Early and Middle Cambrian.	Pliocene and Pleis- tocene.	Late Pennsyl- vanian.	Early Permian (Leonard).	Early Jurassic or Late Triassic.	Ordovician to Devonian.	Late Pennsyl- vanian.	Permian
Steamboat Hills Rhyolite 82	 Stoner Limestone Member (of Stanton Limestone) (of Lansing Group) (of Miscouri Stories) 	Sumner Group	A Swope Limestone (of Kan- sas City Group) (of Miscorris Coricol)	Tintic Quartzite	Tulare Formation	Vilas Shale (of Lansing Group) (of Missouri Sories)	Wellington Formation (of Sumner Group).	White Mountain Plutonic Volcanic Series.	Whiteside Granite	Wyandotte Limestone (of Kansas City Group) (of Misconni Sovioc)	Yorkville Quartz Monzonite.

CHANGES IN STRATIGRAPHIC NOMENCLATURE

Year of	publication	1965		Year of	publication	1964	1001	1964	1965	1964			1965
instated	Publication (U.S. Geol. Sur- vey except as indicated)	Prof. Paper 525-B		bandoned	Publication (U.S. Geol. Sur- vey except as indicated)	Map GQ-303 Prof. Paner 501-D		Map GQ-303	Bull. 1194–N	Bull. 1194–B.	This report, p. A70	This report, p. A49	Bull. 1181-R
Report in which name is reinstated	Title and authorship	Suggestions for prospecting for evaporite deposits in southwest Virginia, by C. F. With- ington.	STRATIGRAPHIC NAMES ABANDONED	Report in which name is abandoned	Title and authorship	Geology of the Tyrone quadrangle, Kentucky, by E. R. Cressman. Three nee-Bull Take tills in the Wind River	Mountains, Wyoming, reinterpreted, by G. M. Richmond.	Geology of the Tyrone quadrangle, Kentucky, by E. R. Cressman. Paintbrush Tuff and Timber Mountain Tuff	of Nye County, Nevada, by P. P. Orkild. Nomenclature and correlation of lithologic sub- divisions of the Jefferson and Three Porks Formations of southern Montana and north-	Wyoming, by C. A. Sandberg, Hovey Group, a redefined stratigraphic name for the Hovey Formation of northeast	Mame, by cours Favlace. Application and use of the name Arikaree For- mation for lower and middle Miocene rocks in Granite Mountains area of central Wyo-	Paintbrush Tuff and Timber Mountain Tuff of Nye County, Nevada, by P. P. Orkild.	Reconnaissance geology of Admiralty Island, Alaska, by E. H. Lathram, J. S. Pomeroy, H. C. Berg, and R. A. Loney.
	Location	Virginia	TRATIGRAPHIC		Location	Kentucky		Kentucky Nevada	Montana, Wyoming, North Dakota, and South Dakota.	Maine	Wyoming	Nevada	Alaska
	Age.	Mississippian	Ø		Age	Middle Ordovician Pleistacene		Middle Ordovician Eocene to Pliocene or	younger. Late Devonian	Early Silurian (?)	early and middle Miocene.	Pliocene or younger	Early Cretaceous (?)
	· ·· Name	Little Valley Limestone			Name	Benson Limestone (of Lexington Group). Dinwoody Lake Till and	Glaciation.	Jessamine Limestone (of Lexington Group). Oak Spring Group.	Potlatch Member (of Three Forks Formation).	Saddleback Mountain Mem- ber (of Hovey Formation).	Split Rock Formation.	Survey Butte Member (of Piapi Canyon Formation)	(of Oak Spring Group). Symonds Formation

STRATIGRAPHIC NAMES REINSTATED

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DENNY CREEK GRANODIORITE GNEISS, BROWNS PASS QUARTZ MONZONITE, AND KROENKE GRANODIORITE, MOUNT HARVARD QUADRANGLE, COLORADO

By FRED BARKER and M. R. BROCK

Three plutonic rocks of Precambrian age in the Mount Harvard quadrangle, Colorado, are here named Denny Creek Granodiorite Gneiss, Browns Pass Quartz Monzonite, and Kroenke Granodiorite. Part of the rock assigned to the Denny Creek was previously called Pikes Peak Granite by Stark (1935) and Stark and Barnes (1935), a designation now known to be erroneous; another part and the other two intrusive rocks and all the enclosing schists and gneisses, were assigned by them to the "Sawatch schist and migmatite". The Denny Creek Granodiorite Gneiss, the oldest of the three plutonic rocks, and the Kroenke Granodiorite, the youngest, form batholiths that extend beyond the boundaries of the Mount Harvard quadrangle (fig. 1). The Browns Pass Quartz Monzonite forms two stocks in the central part of the quadrangle but thus far has not been recognized outside the quadrangle.

The Browns Pass Quartz Monzonite discordantly intrudes the Denny Creek Granodiorite Gneiss both in the Denny Creek drainage basin and at contacts of the small stock lying about 2 miles north of Cottonwood Pass (fig. 1). Xenoliths of the granodiorite gneiss lie in the quartz monzonite along parts of the contacts of these two intrusives. The Kroenke Granodiorite intrudes the Denny Creek Granodiorite Gneiss along many miles of contact (fig. 1) and sharply transects the foliation of the granodiorite gneiss. Although the Browns Pass Quartz Monzonite and the Kroenke Granodiorite are nowhere in contact, the Browns Pass is partly syntectonic whereas the Kroenke is posttectonic; moreover, on the south side of Texas Creek, the Kroenke batholith crosscuts and brecciates gneisses and schists that 1 mile to the south contain syntectonic well-foliated sills and dikes of the Browns Pass unit. The Browns Pass Quartz Monzonite evidently is the older of the two.

DENNY CREEK GRANODIORITE GNEISS

The Denny Creek Granodiorite Gneiss is named for the exposures along Denny Creek at altitudes of 10,800–11,000 feet. This area is here designated as the type area. This formation consists of gneissic biotite granodiorite and biotite-quartz diorite and underlies much of the west-central and north-central parts of the quadrangle (fig. 1). It is medium to dark gray, is well foliated, and consists of lenticular single grains and aggregates of medium- to coarse-grained plagioclase and quartz and, also, augen of pale-pink microcline perthite; the grains, aggregates, and augen are set in a continuous schistose matrix

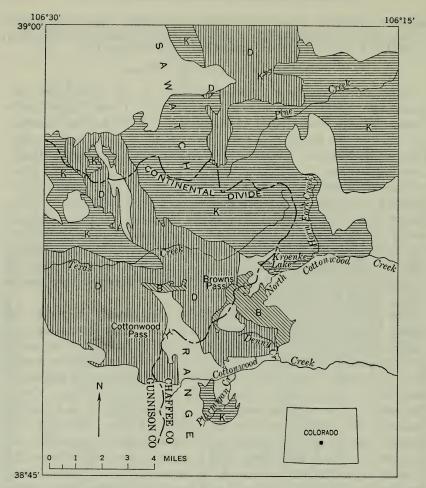


FIGURE 1.—Generalized distribution of major Precambrian intrusive rocks, Mount Harvard quadrangle, Colorado. D, Denny Creek Granodiorite Gneiss; B, Browns Pass Quartz Monzonite; K, Kroenke Granodiorite.

of fine-grained biotite that, in addition, commonly contains either hornblende or muscovite. Proportions of minerals, especially microcline, differ from one hand specimen to the next in most parts of the batholith so that quartz diorite and granodiorite are closely intermixed in many outcrops. The plagioclase commonly is andesine, rimmed with oligoclase, and is slightly to markedly sericitized. The biotite is pleochroic, X being straw yellow and Y and Z dark brown to dark green. The hornblende is pleochroic, X being golden yellow brown, Y green, and Z bluish green. Chemical and modal analyses

CHANGES IN STRATIGRAPHIC NOMENCLATURE

TABLE 1.—Chemical and modal analyses of Denny Creek Granodiorite Gneiss, Browns Pass Quartz Monzonite, and Kroenke Granodiorite, Mount Harvard quadrangle, Colorado

	Denny Creek Granodio- rite Gneiss				s Pass Monzonii		Kroenke Granodiorite			
	Bf-1	Bf-17	Bf-18	Bf-20	Bf-76	Bf-141	Bf-111	Bf-176	Bf-196	
Chemical Analyses [Analyses by E. S. Daniels]										

$\begin{array}{l} SiO_2 \\ Al_2O_3 \\ Fe_2O_3 \\ Fe_2O_4 \\ FeO \\ MgO \\ CaO \\ Na_2O \\ Na_2O \\ Na_2O \\ H_2+ \\ H_3O \\ TiO_2 \\ FeO \\ H_2 \\ H_2O \\ TiO_2 \\ FeO \\ MnO \\ CO_2 \\ CO_2 \\ CO_2 \\ \end{array}$	$58.11 \\ 16.29 \\ 3.45 \\ 5.85 \\ 2.14 \\ 5.17 \\ 2.97 \\ 2.29 \\ 1.19 \\ .12 \\ 1.52 \\ .49 \\ .19 \\ .02$	$59.14 \\ 14.65 \\ 3.06 \\ 7.13 \\ 2.54 \\ 4.23 \\ 2.58 \\ 2.57 \\ 1.46 \\ .08 \\ 1.67 \\ .45 \\ .17 \\ .02$	$\begin{array}{c} 61.73\\ 15.66\\ 2.88\\ 4.12\\ 1.48\\ 2.91\\ 3.00\\ 4.38\\ 1.83\\ .10\\ 1.05\\ .35\\ .11\\ .01\end{array}$	$\begin{array}{c} 69.\ 71\\ 14.\ 02\\ 1.\ 78\\ 2.\ 36\\ .\ 80\\ 1.\ 91\\ 2.\ 82\\ 4.\ 32\\ .\ 97\\ .\ 14\\ .\ 66\\ .\ 25\\ .\ 08\\ .\ 06\\ \end{array}$	$\begin{array}{c} 65.04\\ 15.14\\ 1.96\\ 3.60\\ 1.65\\ 3.34\\ 3.07\\ 3.40\\ .93\\ .93\\ .21\\ .95\\ .35\\ .08\\ .02\end{array}$	$\begin{array}{c} 74.17\\ 14.12\\ .41\\ .63\\ .96\\ 3.76\\ 5.01\\ .44\\ .18\\ .11\\ .04\\ .05\\ .02\\ \end{array}$	$\begin{array}{c} 76.\ 80\\ 15.\ 71\\ .\ 70\\ 1.\ 26\\ .\ 62\\ 2.\ 40\\ 5.\ 05\\ 2.\ 48\\ .\ 50\\ .\ 10\\ .\ 26\\ .\ 08\\ .\ 03\\ .\ 01\\ \end{array}$	$\begin{array}{c} 68.32\\ 16.60\\ .85\\ 1.98\\ .97\\ 2.74\\ 5.67\\ 1.51\\ .66\\ .05\\ .38\\ .11\\ .04\\ .03\end{array}$	$\begin{array}{c} 66.88\\ 16.88\\ 1.07\\ 2.20\\ 3.01\\ 5.45\\ 1.98\\ .50\\ .52\\ .15\\ .05\\ .03\end{array}$
CO ₂		. 02 . 03 . 10	.01 .03 .09	.08 .06 .02 .10	.03 .02 .02 .24	.02 .02 .02	.01 .01 .01	.03 .02 .06	. 03 . 03 . 03 . 08
Subtotal Less O	99.95 .05	99.88 .05	99.73 .05	100.00	100.00	100.11	100.07	99.99 .03	100.08
Total	99.90	99.83	99.68	99. 95	99.90	100. 11	100.04	99.96	100.04

Mode

[Analyses by R. C. Bucknam]

Plagioclase Potassic feldspar Quartz Biotite Hornblende	50 25. 5 21	49 22 18 8	$47 \\ 15 \\ 21 \\ 14 \\ .5$	31 29. 5 29. 5 5. 5	32.5 17.5 29.5 16.5	36.5 28 32 3.5	$53 \\ 10 \\ 30 \\ 4.5$	$63.5 \\ 1 \\ 25 \\ 9$	$63.5 \\ 1.5 \\ 22 \\ 12 $
Muscovite Sphene Magnetite and ilmenite Epidote	1.5 1.5	Trace 2.5	Trace 1	1.5 Trace 1.5	1.5 1 Trace	Trace Trace	1 .5 .5 Trace	Trace Trace	Trace
Allanite Apatite Zircon	Trace Trace	Trace Trace	Trace Trace Trace	.5 Trace	Trace .5 Trace	Trace	Trace Trace Trace	Trace Trace Trace	. 5 Trace

Bf- 1. Gneissic biotite-quartz diorite, 0.1 mile north of Pass Creek, 11,040 ft altitude.
Bf- 17. Gneissic biotite-hornblende-quartz diorite, 100 ft east of north fork of Denny Creek, 10,970 ft altitude.
Bf- 18. Gneissic biotite-hornblende granodiorite, 150 ft east of north fork of Denny Creek, 10,970 fet altitude.
Bf- 20. Leucoquartz monzonite, 0.25 mile southeast of Browns Pass, 11,960 ft altitude.
Bf- 76. Biotite-quartz monzonite, 0.18 mile north of Peak 12,955; 12,760 ft altitude.
Bf-141. Alaskite, ridge between north fork of Denny Creek and Delaney Gulch, 0.1 mile north of 12,320,ft closed contour, 12,300 ft altitude.
Bf-111. Leucogranodiorite, northwest side of Peak 12,778; 12,220 ft altitude; 0.4 mile west-northwest of Kreenke Lake.

Kroenke Lake.

Bf-176. Leucoquartz diorite, 0.12 mile southeast of Lake Rebecca, 12,270 ft altitude. Bf-196. Biotite-quartz monzonite, north side of canyon of Pine Creek, 0.11 mile northeast of Peak 11,391; 11,150 ft altitude.

of three typical varieties of the granodiorite gneiss are presented in table 1.

BROWNS PASS QUARTZ MONZONITE

The Browns Pass Quartz Monzonite, named for exposures west and southeast of Browns Pass (fig. 1), here designated the type area, consists of three rock types that, from oldest to youngest, are (1)

buff medium- to coarse-grained foliated homogeneous biotite-quartz monzonite (sample Bf-76, table 1), found northeast of Browns Pass; (2) pink to buff coarse-grained massive to foliated quartz monzonite and granite (sample Bf-20, table 1), found southeast to west of Browns Pass and also in the small stock north of Cottonwood Pass; and (3) buff fine- to medium-grained massive alaskite (sample Bf-141, table 1), found as dikes southeast of Browns Pass. The plagioclase of the different varieties ranges from median andesine to calcic oligoclase, and much of it has been altered to sericite. The potassic feldspar is well-twinned slightly perthitic microcline. The biotite is typically pleochroic, X being straw yellow and Y and Z dark olive brown. Chemical and modal analyses of the Browns Pass Quartz Monzonite are shown in table 1.

KROENKE GRANODIORITE

The Kroenke Granodiorite, named for excellent exposures west of Kroenke Lake, the type area, consists of quartz diorite, granodiorite, and quartz monzonite in a sharply discordant body that underlies much of the drainage areas of Pine Creek, North Cottonwood Creek, and the uppermost reaches of Texas Creek, and a small pluton in the Ptarmigan Creek drainage area in the southern part of the quadrangle. It ranges in composition from quartz monzonite to quartz diorite, but these types are so similar in outcrop and hand specimen that one cannot be distinguished from the other in the field. The rocks are mostly quartz diorite in the Pine Creek drainage area and granodiorite and quartz monzonite from Horn Fork Creek to Texas Creek. Thev are white, light gray, or buff and fine to medium grained; the rocks are commonly foliated and contain single grains and wisplike clusters of biotite but in many areas are massive. They range from homogeneous forms to banded forms consisting of alternating biotite-rich and quartz feldspar-rich laminae. The granodiorite contains swarms of amphibolite inclusions in some places.

Plagioclase in the Kroenke Granodiorite is calcic oligoclase or sodic andesine. Microcline is well twinned and slightly perthitic to nonperthitic. The biotite is pleochroic, X being straw yellow or pale brownish green and Y and Z dark olive green or greenish brown. The biotite commonly is partly altered to chlorite and typically is intergrown with apatite, epidote, sphene, magnetite, and allanite. In the vicinity of amphibolite inclusions the rocks contain hornblende. Three chemical and modal analyses are shown in table 1.

SOUTH PASS FORMATION ON THE SOUTHWEST FLANK OF WIND RIVER MOUNTAINS, WYOMING

By NORMAN M. DENSON, HOWARD D. ZELLER, and E. VERNON STEPHENS

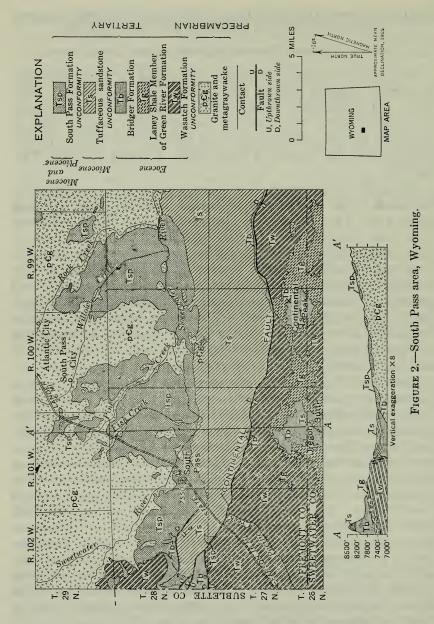
The name South Pass Formation is here used for a generally conglomeratic sequence of rocks that includes beds of sandstone, limestone, and volcanic ash; the rocks have a combined average thickness of about 350 feet. The formation unconformably overlies rocks of early and middle Miocene, early and middle Eocene, and Precambrian ages in the vicinity of South Pass, Fremont County, Wyo. (fig. 2).

The rocks in the South Pass Formation were first described by Comstock (1874) as beds of conglomerate, sandstone, and marl near South Pass, Wyo.; however, his names South Pass Group and South Pass Beds (Comstock, 1874, chart opposite p. 102 and p. 130, respectively) were never formally defined, and these names have not been used by other geologists working in the area (Wilmarth, 1938, p. 2032). A section of the rock sequence showing stratigraphic and structural relations is exposed along the Continental Divide near South Pass (alt 7,550 ft); therefore, the authors believe that the name South Pass, given to these beds by Comstock, should be retained as a formation name. The area around South Pass in Tps. 28 and 29 N. and Rs. 98 through 103 W. is herein designated as the type area.

The South Pass Formation is composed dominantly of a basal pinkish-gray pebble-to-boulder conglomerate having a matrix of fine-grained tuffaceous sandstone and siltstone that locally contains pebbles of moss agate. Locally the conglomerate is interbedded with very coarse grained arkosic sandstone and is overlain by fine- to coarse-grained laminated strongly fluorescent sandstone as much as 200 feet thick and fresh-water limestone and volcanic ash as much as 15 feet thick. The thickness of the formation ranges from 0 to more than 500 feet. More than 100 feet of this formation is exposed along many north-facing escarpments on the south side of the Sweetwater River near the center of the map area. There the South Pass Formation is heterogeneous and includes boulders and cobbles of Precambrian metamorphic and igneous rocks set in a matrix consisting of both fine-grained tuffaceous siltstone, reworked from the White River Formation (Oligocene), and tuffaceous sandstone, which is from the lower and middle Miocene sequence.

The South Pass Formation fills preexisting valleys and forms pediment and coalescing alluvial fanlike deposits along both flanks at the southern end of the Wind River Mountains. The heterogeneity of the rocks and the abrupt changes in lithology along the outcrop are

CONTRIBUTIONS TO STRATIGRAPHY



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distinguishing characteristics of the unit and probably represent torrential, short-duration deposition following major uplift of the Wind River Mountains at the end of middle Miocene time. The formation has been mapped as irregular, discontinuous patches in an area of about 350 square miles which extends from the mouth of Rock Creek northwestward to a point near the head of Big Sandy Creek in T. 30 N., 104 W.—a distance of about 40 miles. The rocks herein referred to as the South Pass Formation have been designated as post-Bridgerian (McGrew and others, 1959, fig. 1), as basal part of the White River (Nace, 1939, pl. 1), and as middle or upper Eocene (Love and others, 1955).

Although the exact age of the South Pass Formation is unknown, it is no older than the lower and middle Miocene rocks which it overlies near the west-central edge of the map area. At various places in the area, the South Pass also overlies the Bridger (middle Eocene) and Wasatch (early Eocene) Formations and in addition Precambrian granite and metagraywacke. Northwest of the area in figure 2, in T. 30 N., R. 104 W., and along the northeast flank of the Wind River Mountains in Tps. 29 and 30 N., Rs. 98 and 99 W., the formation rests on the beveled edge of as much as 800 feet of the White River Formation (Oligocene). The South Pass Formation is overlain by pediment gravel and alluvium of Quaternary age.

Four miles west of the map area, near the McCann Ranch (SW $\frac{1}{4}$ T. 28 N., R. 103 W., and NW $\frac{1}{4}$ T. 27 N., R. 103 W.), and in the SW cor. T. 28 N., R. 102 W. (fig. 2), the formation is offset by movement along the Continental fault and the beds are tilted as much as 45°. The cementation of the basal conglomerate is noticeably greater along the Continental fault whereas elsewhere the conglomerate is generally semiconsolidated. If the latest movement along the Continental fault spot-middle Pliocene, which regional evidence indicates (Love, 1954, p. 1312), the South Pass Formation may be considered as late Miocene to middle Pliocene in age. Because the South Pass Formation can be traced eastward discontinuously into fossiliferous rocks assigned to the Ogallala Formation (upper Miocene and Pliocene) in southeastern Wyoming and because it has lithologic, stratigraphic, and structural relations similar to those of the Ogallala, a correlation between the two formations is suggested.

NOMENCLATURE AND AGE OF FORMATIONS IN THE ANSONIA QUADRANGLE, FAIRFIELD AND NEW HAVEN COUNTIES, CONNECTICUT

By CRAWFORD E. FRITTS

The names and geologic ages of most of the formations in this area are discussed by Fritts (1962a, b). Additions to the nomenclature and revisions of geologic ages are discussed briefly here.

The Oronoque Member of the Derby Hill Schist is named here for the unincorporated community of Oronogue (pronounced Or-o-noke') in the northwestern part of the adjacent Milford guadrangle. The type locality of this member is a long roadcut at the intersection of the Merritt Parkway and State Route 110 on the west bank of the Housatonic River about 1,300 feet south of the Ansonia quadrangle. The rocks exposed there are in the kyanite zone of regional metamorphism and are similar to rocks mapped as part of the Oronoque Member just north of the Far Mill River in the western part of Pine Rock Park in the Ansonia quadrangle. Abundant quartz-rich paragneiss, or siliceous metatuff, distinguishes this member from the predominant schist of the Derby Hill Schist, which is also well exposed just north of the Far Mill River. Farther north near Shelton and Derby the entire formation is highly sheared; the Oronoque Member there is difficult to recognize but presumably is in the upper part of the formation, and grades into the predominant schist. Rocks of the Oronoque Member in zones of low-grade metamorphism in the eastern part of the Ansonia and Milford quadrangles are also difficult to recognize, but in at least some places there the rocks appear to contain more primary (nonintroduced) silica than the underlying phyllites, which probably are equivalent to the main part of the Derby Hill Schist. The Oronogue Member, especially the upper part, also contains subordinate impure limestone of metasedimentary origin and amphibolite and greenschist, which probably represent mainly metavolcanic rocks similar to those found in the overlying Maltby Lakes Volcanics.

The names Maltby Lakes Volcanics and Allingtown Metadiabase were chosen in consultation with John Rodgers and J. E. Sanders of Yale University, whose students, Holdaway¹ and Burger,² mapped rocks of pre-Triassic age in the western part of the New Haven quadrangle during the period 1957–62 when work by the present writer was in progress in the Mount Carmel, Southington, Ansonia, and Milford quadrangles. The rocks mapped by Holdaway and

¹Holdway, M. J., 1958, The bedrock geology of the Maltby Lakes area [Conn.]: Yale Univ. unpub. senior thesis.

² Burger, H. R., 1962, Stratigraphy and structure of the Milford Group, New Haven quadrangle, Connecticut: Yale Univ. unpub. senior thesis.

Burger formerly were included in the Milford Chlorite Schist (Rice and Gregory, 1906), a name recently abandoned (Fritts, 1962a).

A threefold division of the Milford Chlorite Schist was made by Burger. A southeastern unit, which Burger called the Savin Schist, consists mainly of chloritic phyllite similar to that exposed near Savin Rock in the New Haven quadrangle. This unit extends southwestward across the northwest corner of the Woodmont quadrangle and into the Milford quadrangle, where it is interpreted as low-grade Derby Hill Schist (Fritts, 1965). A central unit, which Burger called the Allingtown Formation, consists of abundant metadiabase intruded into phyllitic metasedimentary rocks. This unit extends southwestward into the southeast corner of the Ansonia quadrangle and into the adjacent Milford quadrangle, where the metasedimentary rocks are mapped separately as low-grade Oronoque Member of the Derby Hill Schist. Burger's northwestern unit, which includes rocks mapped previously by Holdaway near the Maltby Lakes Reservoirs in the New Haven quadrangle, consists mainly of metavolcanic rocks but contains subordinate metasedimentary rocks and minor intrusive metadiabase similar to that characteristic of Burger's Allingtown Formation. This northwestern unit of predominantly metavolcanic rocks also extends southwestward into the Ansonia and Milford quadrangles, where it lies above the Derby Hill Schist and unconformably below the Wepawaug Schist. The present writer uses the name Maltby Lakes Volcanics for the metavolcanic rocks and the name Allingtown Metadiabase only for intrusive metadiabase or metabasalt which is probably younger than both the Derby Hill Schist and the Maltby Lakes Volcanics.

The unit of rocks mapped as Maltby Lakes Volcanics is named here for the Maltby Lakes Reservoirs near State Route 34 in the New Haven quadrangle; this locality is also designated the type locality. The unit contains a distinctive basal pyroclastic schist, characterized by numerous lapilli and bomblike masses of metabasalt, but consists mainly of a thick sequence of greenschists and amphibolites. The rocks probably represent metamorphosed marine tuffs and(or) lava flows. The formation also contains minor metasedimentary schists and impure limestones. The stratigraphic position occupied by the Maltby Lakes Volcanics is comparable to that of the Barnard Volcanic Member of the Mississquio Formation of Doll, Cady, Thompson, and Billings (1961) in Vermont. The age of the Maltby Lakes Volcanics, therefore, probably is Ordovician(?). The basal pyroclastic schist is well exposed at the type locality and was mapped by Holdaway and Burger (see footnotes 1 and 2, p. A30) as a separate unit, although they did not identify it as a metamorphosed pyroclastic rock. They also mapped several thin but apparently distinctive greenschist and amphibolite units now known to be stratigraphically above the pyroclastic schist and below the Wepawaug Schist. The thin stratigraphic units beneath the Wepawaug, however, have not been mapped separately in the Ansonia and Milford quadrangles mainly because of a scarcity of outcrops in these areas and in the southwestern part of the New Haven quadrangle. In the Ansonia and Milford quadrangles, therefore, the name Maltby Lakes Volcanics is used for all metavolcanic (extrusive) rocks and subordinate metasedimentary rocks that are underlain by the Derby Hill Schist and are overlain unconformably by the Wepawaug Schist.

The Allingtown Metadiabase is named here for the community of Allingtown (spelled Allington on recent topographic maps of the New Haven quadrangle) just southwest of the city of New Haven on U.S. Highway 1. This is also the type locality of the formation. The rock is a metamorphosed porphyritic diabase or basalt, which forms numerous closely spaced sills and gently inclined dikes in the metasedimentary rocks interpreted by the writer as low-grade Oronoque Member of the Derby Hill Schist. Dikes of similar porphyritic rocks also intruded the Maltby Lakes Volcanics sometime before metamorphism but are not known to have intruded the Wepawaug Schist. The age of the metadiabase therefore is probably Ordovician.

The age of the Prospect and Ansonia Gneisses formerly was reported as Ordovician or Devonian (Fritts, 1962a); however, the age of the Ansonia now is believed to be Devonian, because in this quadrangle the rock appears to be a gneissic equivalent of the Woodbridge Granite. The Woodbridge intruded the Wepawaug Schist of Silurian and Devonian age before the climax of progressive regional metamorphism in Middle to Late Devonian time (Fritts, 1962b). The regional distribution of the Prospect Gneiss (Fritts, 1962a, fig. 128.1) suggests that this intrusive rock was emplaced mainly along the eastern flank of a north-trending line of domes, such as the Waterbury dome, after initial folding of the Wepawaug Schist and underlying metasedimentary rocks but before emplacement of the Ansonia Gneiss, which crosscuts the Prospect. Thus the age of the Prospect Gneiss also is reported here as Devonian.

BELTED RANGE TUFF OF NYE AND LINCOLN COUNTIES, NEVADA

By K. A. SARGENT, D. C. NOBLE, and E. B. EKREN

As the result of mapping about 2,000 square miles north and west of the Nevada Test Site during the past two years, problems in naming the volcanic rock units have developed that cannot be solved by use of the previous nomenclature applied to rocks within the Test Site. Volcanic rocks, more than 10,000 feet thick and equivalent in age to the Indian Trail Formation (Poole and McKeown, 1962), have been divided into more than 20 mappable units, some of which are several thousand feet thick. In contrast, the Indian Trail Formation in the Test Site is only about 1,000 feet thick and is subdivided into an informal lower member and the Tub Spring and Grouse Canyon Members (Hinrichs and Orkild, 1961; Poole and McKeown, 1962). Though some age equivalence is certain, use of the name Indian Trail Formation for the thick sequence of volcanic rocks north of the Test Site would imply lithologic and genetic relations that do not exist and would preclude the development of a reasonable nomenclature for the rocks north of the Test Site. For these reasons, the name Indian Trail Formation is here restricted geographically to the eastern part of Nevada Test Site, that is, east of long 116°15' (fig. 3).

The recent geologic mapping has shown that the Tub Spring and Grouse Canyon Members underlie a large area north and west of the Test Site. As the Indian Trail Formation is restricted to the Test Site, the lateral extensions of the rocks of the Tub Spring and Grouse Canyon Members are no longer parts of a formal stratigraphic unit outside the Test Site; therefore, north and west of the Nevada Test Site, the Tub Spring and Grouse Canyon Members are placed in a new formation herein named the Belted Range Tuff (fig. 4).

BELTED RANGE TUFF

The Belted Range Tuff crops out over a broad area in the southern parts of the Belted and Kawich Ranges and in the vicinity of Pahute Mesa (fig. 3). In addition, the Grouse Canyon Member is present over an extensive area west of that shown on figure 3. The formation is best exposed in the southern part of the Belted Range, which is designated the type area. At most places the Belted Range Tuff unconformably overlies local informal units of tuff and lava and is unconformably overlian by tuff, lava, the Paintbrush Tuff, the Timber Mountain Tuff, or the Thirsty Canyon Tuff (p. A44; Noble and others, 1964). Many of the lava flows that underlie or overlie the tuff in the southern part of the Belted and Kawich Ranges have a soda-rhyolite composition and are genetically related to the tuff.

Both the Tub Spring and Grouse Canyon Members are compound cooling units of comenditic (peralkaline soda-ryholite) ash-flow tuff. Rocks of both members are peralkaline. Soda-rich sanidine is the dominant phenocryst mineral in rocks of both members; fayalite, apatite, zircon, and sodic iron-rich clinopyroxene are minor ubiquitous phenocrysts. The Grouse Canyon Member is virtually quartz

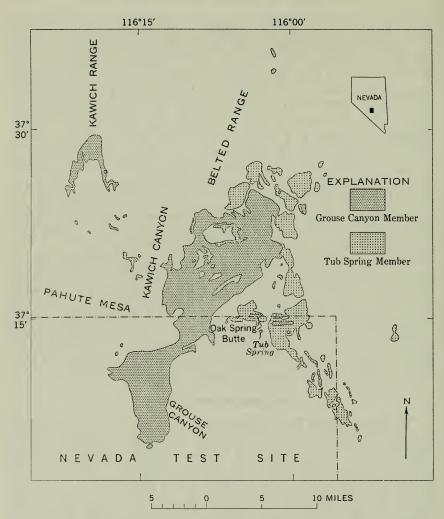
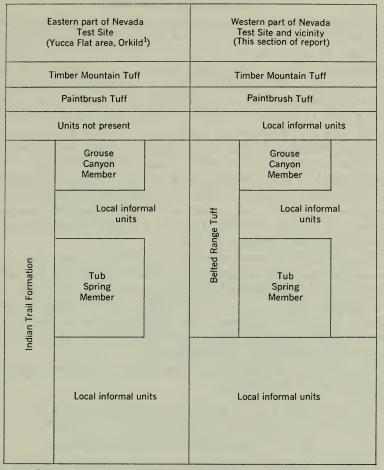


FIGURE 3—Distribution of the Tub Spring and Grouse Canyon Members of the Belted Range Tuff in the northeastern part of the Nevada Test Site and the southeastern part of the Las Vagas Bombing and Gunnery Range.

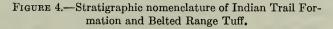
free, whereas the Tub Spring Member contains abundant quartz phenocrysts.

TUB SPRING MEMBER

At the type locality, Tub Spring (fig. 3), the Tub Spring Member is approximately 250 feet thick; north of Oak Spring Butte the member is locally 300 feet thick. Rocks of the member are typically buff or bluish gray but locally are brick red. The degree of welding



¹ See p. A44-A51.



ranges from poor to dense. At most outcrops the member is devitrified, but poorly welded vitric tuff, in many places partly to completely zeolitized, commonly occurs at the base and top of the member; vitrophyre is locally present where the member is thick and densely welded. Phenocrysts compose approximately 20-25 percent of the member. Acicular crystals of acmite and sodic amphibole of vapor-phase origin are typical of poorly welded parts. The member locally contains large fragments of reddish-brown or buff pumice and lithic fragments of rhyolite, welded tuff, and Paleozoic sedimentary rocks.

GROUSE CANYON MEMBER

At Grouse Canyon, the type locality (fig. 3), the Grouse Canyon Member is 80 feet thick; north of Oak Spring Butte it has a maximum thickness of approximately 300 feet. The member, as originally described by Hinrichs and Orkild (1961), is composed of an upper part of ash-flow tuff and a lower part of ash-fall and reworked tuff. Although the upper part of the member was described as a simple cooling unit (Poole and McKeown, 1962), further work has demonstrated that it is a compound cooling unit. In most places the unit is densely welded and, with the exception of a thin basal vitrophyre, is almost everywhere devitrified. Gas cavities containing vaporphase crystals of sodic amphibole are common. Rocks of the cooling unit range from greenish to bluish gray to gray buff and brown and locally are brick red. The various ash flows composing the compound cooling unit contain from less than 0.1 to approximately 20 percent phenocrysts, and average 5 percent.

The ash-fall and reworked tuffs composing the lower part of the Grouse Canyon Member are indistinguishable from ash-fall and reworked tuffs locally intercalated with lavas both above and below the member. Correlation of such tuffs over distances of several miles is uncertain; the lower part of the member is, therefore, no longer retained anywhere as a part of the member; these tuffs are treated as informal units of the formation.

On the basis of the data given on page A51, the Belted Range Tuff is of Miocene or Pliocene age.

SALYER AND WAHMONIE FORMATIONS OF SOUTHEASTERN NYE COUNTY, NEVADA

By F. G. POOLE, W. J. CARR, and D. P. ELSTON

A Tertiary sequence of lava flows, volcanic breccia, tuff, and sandstone is exposed in the southern part of the Nevada Test Site. Part of this sequence of rocks is divided into two new formations here named the Salyer and Wahmonie Formations. The Salyer Formation, which crops out in an area exceeding 300 square miles that is centered near Mount Salyer and Cane Spring (fig. 5), is best exposed in the Cane Spring quadrangle. Its maximum exposed thickness is nearly 2,000 feet at its type area in the vicinity of Mount Salyer and Cane Spring (fig. 6). The original volume of the Salyer was at least 20 cubic miles.

Overlying the Salyer is the Wahmonie Formation, which crops out in an area exceeding 500 square miles centered near its type area at Wahmonie Flat (fig. 5). The Wahmonie is best exposed in the Cane Spring and Skull Mountain quadrangles. Its maximum exposed thickness is 3,500 feet near Wahmonie Flat (fig. 6). The original volume of the Wahmonie was at least 25 cubic miles.

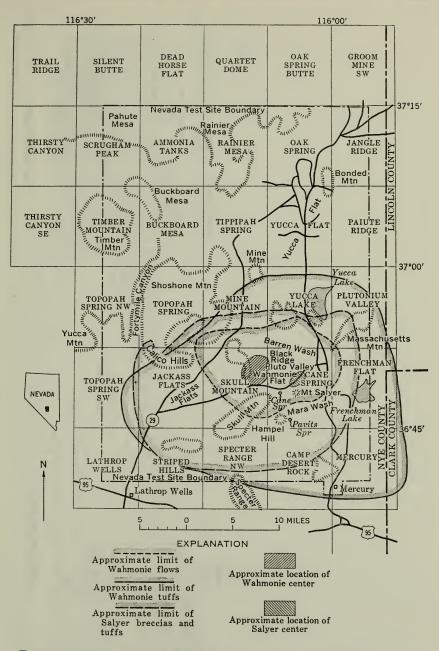
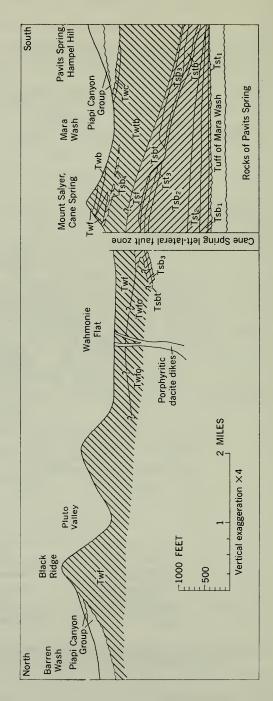


FIGURE 5.—Map of Nevada Test Site showing topographic quadrangles and approximate distribution of Salyer and Wahmonie Formations.

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IGURE 6.—Partly restored generalized north-south section in the Cane Spring quadrangle showing relations between volcanic units. Explanation of symbols given below:	Salyer Formation (ruled pattern slanted to right): Tinit 10 (Tsh.): breecia flow	9 (Tsbt): interbedded lithic breecia, breecia flow, lithic tuff-breecia, tuff, and sandstone	8 (Tsf): lava flow 7 (Tsb ₃): breccia flow	6 (Tstb): biotite tuff 5 (Tst3): interbedded tuff, sandstone, siltstone, and	claystone 4 (Tsb ₂): breccia flows	3 (Tst ₂): interbedded tuff, sandstone, lithic tuff- breccia, and lithic breccia	2 (Tsb ₁): breccia flow 1 (Tst ₁): tuffaceous claystone	
IGURE 6.—Partly restored generalized north-south section in the Explanation of s	Zahmonie Formation (ruled pattern slanted to left): TT-2 = 7	4 (Twc): conglomerate and lithic tuff-breecia 3 (Twb): lithic breecia	2 (Twtb): interbedded tuff, sandstone, lithic tuff- breccia, and pumice agglomerate	1 (Twfa): older lava flows				

FIGURE 6.—Partly rest

Wahmonie Formation

CHANGES IN STRATIGRAPHIC NOMENCLATURE

The type section of the Salyer Formation is a composite of partial sections in the area from Cane Spring to Mount Salyer to Hampel Hill (figs. 5, 6). The type localities are as follows: Units 1 to 3, a north-trending ridge near the head of Mara Wash about $1\frac{1}{4}$ miles south of Cane Spring; units 4, 5, 9, 10, a northeast-trending ridge about half a mile southeast of Cane Spring; unit 6, about 1 mile northeast of Hampel Hill; units 7 and 8, about three-fourths of a mile west of Mount Salyer. As only part of unit 9 is present on the ridge half a mile southeast of Cane Spring, a well-exposed section extending for about half a mile northeast from Mount Salyer also is considered the type locality for this unit.

The type section of the Wahmonie Formation is also a composite of partial sections found on the east and south sides of Wahmonie Flat, at Mount Salyer, and in the area between Cane Spring and Hampel Hill (figs. 5, 6). The type localities are as follows: Unit 1, a small exposure on a low ridge about 1¼ miles northwest of Cane Spring; unit 2, on a northeast-trending ridge about half a mile southeast of Cane Spring and around Hampel Hill; unit 3, at Mount Salyer; unit 4, on the northwest side of Hampel Hill; unit 5, incomplete sections at the east end of Wahmonie Flat and on the north side of Skull Mountain.

STRATIGRAPHIC RELATIONS

In general the Salyer and Wahmonie Formations occur in the middle part of the Tertiary volcanic sequence at the Nevada Test Site (Poole and McKeown, 1962, fig. 80.2). In the type area the Salyer Formation unconformably overlies yellowish zeolitized rhyolitic tuffs of Mara Wash and underlies the Wahmonie Formation (fig. 6). The Wahmonie is older than most of the Paintbrush Tuff (see p. A44) of the Piapi Canyon Group. On Skull Mountain the Wahmonie is directly overlain by the Topopah Spring Member of the Paintbrush. Northwest of Wahmonie Flat a late flow of the Wahmonie occurs between the Topopah Spring and Tiva Canyon Members of the Paintbrush.

Although the Salyer Formation is in general overlain by the Wahmonie, one unit of the Salyer (Tsb₄, fig. 6) intertongues with the tuffs of the Wahmonie south of Cane Spring.

LITHOLOGIC DESCRIPTION

In some areas the Salyer and Wahmonie are not easily separated, but generally rocks of the Wahmonie are more mafic, contain primary hornblende, are less altered, and consist mainly of lava flows and related tuffs, whereas the rocks of the Salyer Formation are more

A40

acidic, contain primary pyroxene and secondary hornblende, are more altered, and consist of breccia flows and interstratified tuff, sandstone, and volcanic breccia.

The Salyer Formation is divided into 10 informal units and the Wahmonie into 5 informal units, as described and illustrated on figure 6. Unless otherwise stated, contacts between units are sharp or abruptly gradational.

The following definitions are for terms used in the stratigraphic sections of the Wahmonie and Salyer Formations.

- Lithic tuff-breccia. An extrusive multilithic breccia that is composed of fragments of angular to rounded older volcanic rock and sparse pre-Cenozoic sedimentary rock contained in a matrix of lapilli and ash tuff.
- Lithic breccia. A breccia similar to lithic tuff-breccia except that it has a matrix of finely fragmented previously formed volcanic rock. In addition, lithic breccia includes some monolithic breccia.
- **Pumice agglomerate.** An extrusive monolithic rock that is composed mainly of angular to rounded pumice blocks and lapilli in a matrix of ash tuff.
- **Breccia flow.** Lava flows that are almost entirely composed of angular to subrounded fragments of altered previously solidified lava contained in a less altered lava matrix of similar composition. Most fragments probably formed prior to, rather than during, emplacement of unit. Fragments occur throughout entire flow body.
- Flow breccia. An extrusive breccia on the margins of lava flows, formed by the breaking up of lava that was solidifying as it moved. The matrix is solidified lava and comminuted flow rock.

Composite sections of Wahmonie and Salyer Formations [Maximum measured thickness is given]

Wahmonie Formation:

- 4. Conglomerate or lithic tuff-breccia consisting of porphyritic stony and glassy flow-banded rhyodacite and dacite cobbles and boulders in a gray tuffaceous sandy matrix. Unit 4 is present only in the southwestern part of the Cane Spring quadrangle_____50
- Gray to pinkish-gray lithic breccia consisting of dense to vesiculated, glassy to devitrified blocks crudely stratified in 1- to 2 ft-thick zones. This unit has a few thin local interbeds of tuff and sandstone_______270

Thickness (feet)

Wahmonie Formation—Continued

- 2. Interbedded tuff, sandstone, lithic tuff-breccia, and pumice agglomerate. The tuff is white, gray, pink, green, and brown; locally contains abundant lithic inclusions of volcanic rock and is partly zeolitized and is generally well bedded. The sandstone is white, gray, or red, very fine to coarse grained, thinly laminated to thick bedded. Lithic tuff-breccia consists of red, purple, and grav blocks of volcanic rock in grav to green tuff matrix. Many breccia zones are crudely stratified, and some are glassy and similar to some of the breccia of unit 3. The pumice agglomerate consists of gray to tan pumice in a lapilli and ash matrix. The pumice contains 25 percent phenocrysts of plagioclase, biotite, and some hornblende and quartz. Silicified conifer wood is common at several places in unit 2. Unit 2 contains a lens of unit 10 of the Slayer, and its basal contact is locally gradational with unit 9 of the Salver. Relations with unit 5 of the Wahmonie north of the Cane Spring fault are not clear_____1,700 1. Red, purple, and brown rhyodacitic lava flows, commonly hydrothermally altered. This unit probably does not occur south of the Cane Spring fault_____ 100 +Salyer Formation: 10. Purple and red breccia flow, petrographically like other breccia flows of the Salyer. The breccia contains 30 percent phenocrysts of andesine-labradorite, partly altered biotite, altered pyroxene, and secondary hornblende, hematite, quartz, and plagioclase. Crudely sorted and locally layered. The unit occurs as a tongue within the Wahmonie Formation in the Cane Spring area 50 9. Interbedded lithic breccia, breccia flow, lithic tuff-breccia, tuff, and sandstone. Top of unit is a grav to green glassy breccia flow as thick as 320 ft. The rocks of this unit contain about 30 percent phenocrysts of labradorite, hypersthene, and biotite. Middle of unit 9 is a lithic breccia as thick as 280 ft that contains large blocks of volcanic rock and a few pieces of Paleozoic carbonate rock. Lower part of unit is gray, pink, and vellow lithic tuff-breccia, crudely bedded locally. At the base of the unit is a widespread gray and green interbedded tuff and sandstone as thick as 90 ft. A major erosional unconformity is at the base. Upper contact is gradational in Cane Spring area 700 8. Purplish lava flow and minor flow breccia. Flow is holocrystalline-porphyritic rock that contains 25 percent phenocrysts of labradorite and altered biotite. Lower contact is locally gradational with unit 7_____ 50 7. Purplish- to reddish-brown breccia flow, similar to unit 10 but thicker. Major erosional unconformity at base transects some earlier structures_____ 450
 - 6. Gray, pink, and orange biotite tuff that contains volcanic rock fragments. Unit is probably two ash-flow tuffs separated by a persistent thin stratified tuff zone. Widespread redbrown sandstone layer is at top. Unit is not present in the Salyer type section near Cane Spring______

200

Salyer Formation—Continued

- 5. Interbedded tuff, sandstone, siltstone, and elaystone. The tuff is mainly olive, green, yellow, and red porphyritic ash and pumice. Reddish-brown and green tuff, containing conspicuous white altered pumice lumps, marks base of unit. Sandstone and siltstone are gray, green, red, and brown, tuffaceous, pebbly, very fine to very coarse grained. Claystone is pinkish gray, silty, sandy, and pebbly, and contains abundant small fragments of volcanic rocks, and a few small pieces of pre-Cenozoic sedimentary rocks in the lower part______
- 4. Breccia flows, purple and red in lower part, grading upward through pale red, pale purple, and pale yellow zones into yellow and olive. Lower part similar to units 7 and 10. Reddish blocks as large as 20 ft across occur in upper part of this unit. Several red-brown breccia zones are in upper part of unit____ 1,000
- Pink, brown, gray, and red interbedded tuff, sandstone, lithic tuff-breccia, and lithic breccia. A blood-red sandstone 1-3 ft thick marks the top of unit 3. The lithic tuff-breccia and lithic breccia contain blocks of volcanic rock and some pre-Cenozoic sedimentary rocks. A widespread breccia zone as much as 15 ft thick in the upper part of unit 3 contains pre-Cenozoic rock fragments_______250
 Mottled purplish, reddish, and brownish breccia flow similar to

CHEMICAL COMPOSITION

Rocks of the Salyer and Wahmonie sequence grade upward chemically and mineralogically from acidic to intermediate composition. Plots of the major elements against silica for rocks of the Wahmonie through the Salyer generally form smooth curves along a differentiation trend common to many volcanic sequences. Chemical analyses and norms of several breccia flows and a lava flow indicate that the Salyer Formation is dellenitic to rhyodacitic in composition. See Nockolds (1954). Lavas of the Wahmonie are modally dacite and andesite, but chemical analyses and norms of several lava flows indicate that the lower part of the Wahmonie is rhyodacite and that the upper part is dominantly dacite. The chemical and mineralogical similarities of the Salyer and Wahmonie strongly suggest that the two formations are comagmatic.

AGE

The Salyer and Wahmonie Formations are dated as late Miocene and as late Miocene and early Pliocene(?), respectively, primarily on potassium-argon age determinations. The Topopah Spring Member of the Paintbrush Tuff (see p.A45–A47), which directly overlies the Wah-

Thickness (feet)

150

monie, has been dated at 13.3 million years ± 0.5 million years (R. W. Kistler, written commun., 1963). The lower and upper part of the Wahmonie, however, have been dated by the same method by Kistler (written commun., 1963) at 12.9 million years ± 0.5 million years and 12.5 million years ± 0.5 million years, respectively. The precision error in these dates may account for the apparent age inversion, as the dates of the Topopah and Wahmonie are only 0.4–0.8 million years apart, less than the possible error. Present data, therefore, place the Wahmonie very close to the Miocene-Pliocene Epoch boundary (Kulp, 1961), and as the Salyer underlies the Wahmonie its age is tentatively considered to be late Miocene.

PAINTBRUSH TUFF AND TIMBER MOUNTAIN TUFF OF NYE COUNTY, NEVADA

By PAUL P. ORKILD

Various names have been applied to the Tertiary rocks in the Nevada Test Site since the turn of the century. In 1907 Ball (p. 31-34) described some of the Tertiary rocks as rhyolite and latite flows and correlated other with the "Siebert tuff (lake beds)" (Spurr, 1905, p. 51-55). Fifty years later the Tertiary volcanic rocks in the eastern part of the Nevada Test Site were named the Oak Spring Formation by Johnson and Hibbard (1957, p. 367) (fig. 7). In 1957 Hansen, Lemke, Cattermole and Gibbons (1963, p. A7) divided the Oak Spring Formation at Rainier Mesa into numbered units (fig. 8). Hinrichs and Orkild (1961, p. D96) divided the Oak Spring Formation in the Yucca Flat area into seven named members and one informal member. In 1962 Poole and McKeown (p. C60, C61) raised the Oak Spring in the Nevada Test Site to the rank of a group and included two formations in it (fig. 8).

Additional detailed mapping of the Test Site and vicinity since 1962 has resulted in new stratigraphic information that requires another revision of the stratigraphic nomenclature of the Tertiary rocks. The use of Oak Spring Group for all Tertiary rocks in the Test Site is no longer practicable because the Tertiary rocks have been divided into a large number of units which have complex relations among the units. This group, therefore, is herein abandoned. Rockstratigraphic units that have similar chemical and mineralogical composition and mode of deposition constitute the major genetic units and are given formational rank. New data have shown that the Piapi Canyon Formation includes two lithologically dissimilar sequences of volcanic rocks that have different source areas. As each sequence is comprised of several units, the Piapi Canyon Formation herein is redefined and raised to the rank of group. It includes two new forma-

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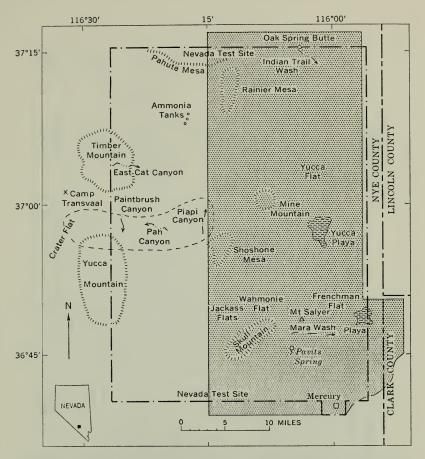
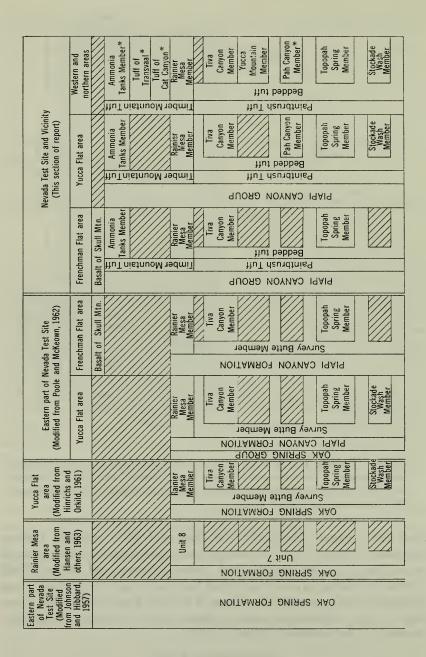


FIGURE 7.—Nevada Test Site and vicinity showing localities referred to in text. Area mapped by Johnson and Hibbard (1957) shown by stipple. The surface and subsurface distribution of Pah Canyon Member of the Paintbrush Tuff is enclosed by dashed line.

tions here named the Paintbrush Tuff and the overlying Timber Mountain Tuff (fig. 8). The rank of the older Indian Trail Formation is unchanged. The names Piapi Canyon Group and Indian Trail Formation, however, are restricted to the areas where they were originally used: namely, the Yucca Flat and Frenchman Flat areas (figs. 7, 9). This is the area (eastern part of Nevada Test Site) mapped by Johnson and Hibbard (1957).

PAINTBRUSH TUFF

The Paintbrush Tuff includes the Stockade Wash, Topopah Spring, and Tiva Canyon Members of Hinrichs and Orkild (1961), the Yucca



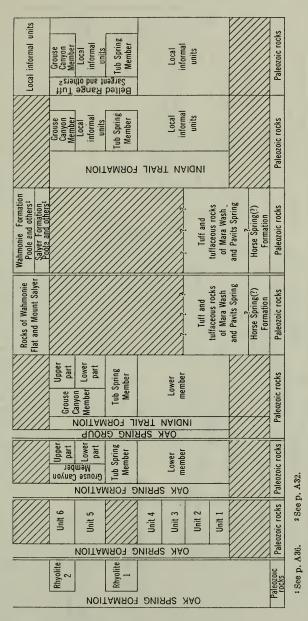


FIGURE 8.--Stratigraphic nomenclature for Tertiary volcanic rocks of the Nevada Test Site. Asterisk (*) indicates No age relations implied between columns. new stratgraphic unit; crosshatching, units not present.

CONTRIBUTIONS TO STRATIGRAPHY

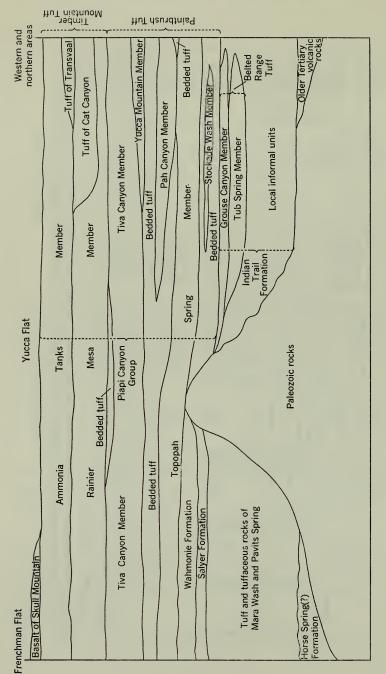


FIGURE 9.-Schematic diagram of the Tertiary volcanic rocks.

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Mountain Member of Lipman and Christiansen (1964), and a new member herein defined as the Pah Canyon Member. These rocks are all similar in mineralogy, chemistry, and areal distribution. The type locality of this formation is in Paintbrush Canyon, which is located 2½ miles northeast of Yucca Mountain (fig. 7).

The Survey Butte Member has been a useful term for certain rocks in the eastern part of the Nevada Test Site. As additional data have become available, however, positive correlation of similar rocks in the same stratigraphic interval has become increasingly difficult as rocks are found farther from the type locality of the member. This difficulty results partly from the fact that the rocks of this member are derived from many sources. To avoid the connotation of positive correlation and implies genetic relations, therefore, the name Survey Butte Member is here abandoned. Rocks in this stratigraphic interval will hereafter be designated as bedded tuff of Paintbrush Tuff.

PAH CANYON MEMBER

The Pah Canyon Member of the Paintbrush Tuff is here named for Pah Canyon (fig. 7). The member is a simple cooling unit of rhyolitic ash-flow tuff and is 280 feet thick at the type section at Pah Canyon. It has a maximum thickness of 300 feet at Yucca Mountain, but generally ranges from 40 to 180 feet in thickness. In most outcrops the member includes a basal zone of ash-fall tuff; this tuff is generally less than 3 feet thick but is locally as much as 15 feet thick. In much of the area of outcrop, the member consists of light-gray to light-brown partly to densely welded devitrified ash-flow tuff containing 5–15 percent phenocrysts of biotite, alkali feldspar, and plagioclase; quartz and clinopyroxene are very rare. The tuff of this member is characterized by abundant small pumice and felsic lithic inclusions and by abundant biotite. The member is conformable with both the underlying Topopah Spring Member and the overlying Yucca Mountain Member.

TIMBER MOUNTAIN TUFF

The well-exposed Timber Mountain Tuff attains its maximum thickness in the Timber Mountain region (fig. 7), which is the type area for the formation. It is unconformable on the Paintbrush Tuff. The formation includes the Rainier Mesa Member formerly of the Oak Spring Formation described by Hinrichs and Orkild (1961); this member is overlain by two informal members and one formal member, which are, in ascending order: tuff of Cat Canyon, tuff of Transvaal, and Ammonia Tanks Member.

The Rainier Mesa Member and Ammonia Tanks Member have greater areal extent than any other members of the Timber Mountain Tuff. They are nearly coextensive and crop out in approximately a

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330° arc around Timber Mountain. The tuff of Cat Canyon is limited to Timber Mountain, and the tuff of Transvaal is restricted to an area southwest of Timber Mountain near Camp Transvaal (fig. 7).

Preliminary study indicates that the four members of the Timber Mountain Tuff have similar chemical and mineralogical compositions. These similarities suggest that the members are comagmatic. Phenocrysts, which comprise 10-40 percent of the rock, include quartz, alkali feldspar, plagioclase, and biotite.

RAINIER MESA MEMBER

The Rainier Mesa Member is well exposed in Piapi Canyon and on Pahute Mesa (fig. 7). The member is a compound cooling unit of rhyolitic to quartz latitic ash-flow tuff containing abundant phenocrysts. Mafic minerals are generally most abundant in the upper or quartz latite part of the cooling unit.

TUFF OF CAT CANYON

The tuff of Cat Canyon is best exposed in Cat Canyon on the east flank of Timber Mountain (fig. 7). The tuff consists of a lower composite ash-flow sheet as much as 2,850 feet thick and an upper compound ash-flow cooling unit 0-150 feet thick (terminology of Smith, 1960). Despite its considerable thickness, tuff of Cat Canyon is confined to the collapsed area of the Timber Mountain caldera (Byers and others, 1963).

TUFF OF TRANSVAAL

The tuff of Transvaal crops out south and southwest of Camp Transvaal, an abandoned mining camp 5 miles southwest of Timber Mountain (fig. 7). The tuff consists of a lower nonwelded rhyolitic ash-flow tuff 0-200 feet thick and an upper compound cooling unit of rhyolitic ash-flow tuff which locally is as much as 300 feet thick.

AMMONIA TANKS MEMBER

The newly named Ammonia Tanks Member is well exposed both a quarter of a mile north of Ammonia Tanks (fig. 7), the type locality for the member, and for several miles along the south rim of Pahute Mesa (fig. 7). The member is a composite cooling unit of rhyolitic to quartz latitic ash-flow tuff which locally reaches a thickness of almost 300 feet in the vicinity of Pahute Mesa but in most places ranges in thickness from 50 to 200 feet. At the type locality the member is 250 feet thick. The lower 5–20 feet of the member is composed of light-gray or pink vitric ash-fall tuff containing about 15 percent white pumice fragments. The basal vitric zone is overlain in much of the area by pink to reddish-purple moderately to densely welded generally devitrified ash-flow tuff containing abundant phenocrysts. The Ammonia Tanks Member is characterized by wedge-shaped sphene crystals as large as 1 mm and by locally abundant reddishbrown porphyritic lithic fragments derived mainly from the underlying Rainier Mesa Member.

AGE

The age of the Oak Spring Group of Poole and McKeown (1962, p. C61), as indicated by plant, invertebrate, and vertebrate fossils, ranges from Eocene to Pliocene.

Potassium-argon age determinations of biotite and sanidine from various units in the Piapi Canyon Group have been made by R. W. Kistler (written commun., 1963), but the data are incomplete; adequate discussion of the results is beyond the scope of this paper. The data do provide, however, an interim and tentative basis for dating the formations. The base of the Indian Trail Formation in the northern part of the Test Site is about 16 million years old; the age of its top is unknown but is older than 13.5 million years. The Paintbrush Tuff ranges from about 13.5 to 12.5 million years in age, and the Timber Mountain Tuff has an age of about 10.5 million years. Translation of these potassium-argon dates to epochs in the geologic scale must be done somewhat arbitrarily because of the inconsistent placement of epochs recorded in the literature (Holmes, 1947, p. 145; Kulp, 1961; and Evernden and others, 1964). Using Kulp's geologic time scale, in which the Miocene-Pliocene boundary is between 12 to 14 million years before the present, the Indian Trail Formation is Miocene and Pliocene(?) rather than Miocene or Pliocene as given in Poole and McKeown (1962, p. C62), the Paintbrush Tuff is Miocene(?) and Pliocene, and the Timber Mountain Tuff is Pliocene.

AGE OF THE ELEANA FORMATION (DEVONIAN AND MIS-SISSIPPIAN) IN THE NEVADA TEST SITE

By F. G. POOLE, P. P. ORKILD, MACKENZIE GORDON, JR., and HELEN DUNCAN

Field investigations in the Nevada Test Site by the authors and colleagues in August 1961 provided considerably more extensive fossil collections from the Eleana Formation than were available when an earlier paper on the formation was prepared (Poole and others, 1961). These collections provided a more reliable basis for dating than the previous collections from which the Eleana was dated as Mississippian and Early Pennsylvanian.

The Eleana was divided into ten informal units lettered A (at the base) through J (at the top). As mapped, unit A contains faunal assemblages that cannot be younger than early Late Devonian (Frasnian). The Devonian corals originally collected were frag-

mentary specimens in the matrix of a limestone conglomerate. Laboratory study did not establish that these specimens were not derived from older rocks and redeposited in beds of Mississippian age; hence, the age of unit A was given as Mississippian, even though it was suggested that unit A might be as old as Lake Devonian. Field study later revealed, however, that layers containing *Atrypa* occur between the conglomeratic coral beds. Inasmuch as *Atrypa* is not known in rocks younger than the lower Upper Devonian and as no fossils characteristic of the Mississippian have been found in unit A, it seems safe to conclude that unit A is of Devonian age.

Units B to G of the Eleana Formation have not yielded fossil collections that provide evidence for precise dating. The authors have recognized no undoubtedly Early Mississippian assemblages in the Eleana; however, several collections which contain fossils that cannot be older than Late Mississippian also contain some fossils that probably were derived from beds of Early Mississippian age.

In 1964 Gordon and Duncan completed a detailed study of the fossils collected from the upper part of the Eleana and the lower beds of the overlying Tippipah Limestone. Information gained from their work indicates that units H, I, and J are of Late Mississippian age and that no rocks of Pennsylvanian age are included in the formation. The basal beds of the Tippipah Limestone, at a locality in the C. P. Hills northwest of Frenchman Flat, 2.2. miles southwest of Yucca Pass in the Yucca Lake quadrangle, contain goniatites (*Branneroceras, Bisatoceras*, and *Boesites?*) indicative of Early Pennsylvanian age.

One collection from unit H or I in the Eleana in the foothills of Quartzite Ridge, in the Oak Spring quadrangle, was originally considered to be of Pennsylvanian age because the only identifiable fossils belonged to a species of the horn coral Lophophyllidium, a genus generally considered to be diagnostic of Pennsylvanian and Permian age. The author's later collection from the same locality yielded, among other fossils, the lithostrotionoid coral Siphonodendron (Cionodendron) sp. and the brachiopod Reticulariina cf. R. spinosa (Meek and Worthen), which are characteristic Late Mississippian fossils. Moreover, at another locality specimens of Lophophyllidium were associated with an extensive assemblage of corals, bryozoans, and brachiopods among which the critical elements for dating purposes are forms of early Late Mississippian age (Zaphrentites aff. Z. californicus (Tischler) Cyathaxonia n. sp., Antiquatonia sp., Auloprotonia? sp., and Semicostella n. sp.).

Several collections made from conglomeratic limestone beds in unit I of the Eleana indicate that at least some of the fossils were derived from erosion of considerably older rocks. A stromatoporoid and corals of Devonian aspect associated with corals and brachiopods indicative of both Early and Late Mississippian age provide evidence of such reworking; however, the conglomeratic layers containing mixed associations occur well below unit J, which carries go itatites (*Cravenoceras hesperium* Miller and Furnish and *C. merriami* Youngquist, which are typical of the *Eumorphoceras bisulcatum* zone) and other fossils characteristically found in late Late Mississippian faunas in the Great Basin. The faunal assemblages in which these goniatites occur seem to be natural associations that are not contaminated by elements diagnostic of the Pennsylvanian or of the pre-Late Mississippian.

MARINETTE QUARTZ DIORITE AND HOSKIN LAKE GRANITE OF NORTHEASTERN WISCONSIN

By WILLIAM C. PRINZ

Much of northern Wisconsin is underlain by granitic rocks of Precambrian age; these rocks compose the so-called Wisconsin batholith. The U.S. Geological Survey, as part of its restudy of the Menominee district, has mapped a small part of the northern edge of the batholith in northern Marinette County and Florence County, Wis. Two units recognized within the batholith—Marinette Quartz Diorite and Hoskin Lake Granite (Prinz, 1959)—are here adopted for use by the Geological Survey. Cain (1963) has mapped the extension of these two units to the south and has distinguished several other granitic units.

Marinette Quartz Diorite forms a pluton of about 12 square miles, mainly in Marinette County. The diorite is here named for that county. It had previously been referred to as biotite-hornblendequartz diorite by Lyons (Emmons and others, 1953, pl. 16 and p. 108). Good exposures of Marinette Quartz Diorite occur south of Niagara, Wis., in the NW¼ sec. 22 and SW¼ sec. 15, T. 38 N., R. 20 E., and along the Chicago, Milwaukee, St. Paul and Pacific Railroad in the S½NE¼ sec. 18, T. 38 N., R. 20 E.

Typical Marinette Quartz Diorite is medium grained, is massive or slightly foliated, and consists of 40-50 percent oligoclase, 0-20 percent potassium feldspar, 10-30 percent quartz, and 20-30 percent prochlorite, biotite, or hornblende. The texture of the rock is hypidiomorphic to allotriomorphic granular consisting of large subhedral to anhedral plates of plagioclase and flakes of biotite or chlorite and consisting of finer grained interstitial grains of quartz and potassium feldspar. Commonly, the oligoclase shows normal zoning and has cores as calcic as andesine; most plagioclase is partly altered to sericite or epidote. Most potassium feldspar is unaltered and shows microcline grid twinning. The major mafic constituent of the quartz diorite is either prochlorite or biotite, and some biotitic samples also contain hornblende. Ubiquitous accessory minerals are magnetite, zircon, and apatite; pyrite, calcite, and tourmaline are found locally.

Hoskin Lake Granite forms an arucate-shaped intrusion of at least 20 square miles that in most places lies between metabasalts of the Quinnesec Formation to the north and Marinette Quartz Diorite to the south. It is here named for Hoskin Lake which lies within the granite pluton in the NE¼ sec. 23, T. 38 N., R. 19 E. Previous workers referred to this unit as porphyritic granite or granite porphyry. Good exposures are found almost everywhere within the mapped limits of the unit; those exposures near the county secondary roads in the $S\frac{1}{2}$ sec. 12 and $N\frac{1}{2}$ sec. 13, T. 38 N., R. 19 E., and the NE¼ sec. 17, T. 38 N., R. 20 E. are readily accessible.

Most Hoskin Lake Granite is coarse grained and porphyritic, having large subhedral, white, gray, or flesh-colored microcline phenocrysts as long as 2½ inches; the phenocrysts are set in a matrix of partially sericitized oligoclase, quartz, microcline, and biotite. In many places the phenocrysts are more or less oriented to form a crude foliation. Chlorite and some hornblende occur with the biotite, and epidote is a sparse alteration product of oligoclase. Zircon, apatite, and subhedral to euhedral sphene, in places altered to leucoxene, are ubiquitous accessory minerals; allanite, monazite, tourmaline, calcite, magnetite, and pyrite are also present in some of the granite. Commonly, half the granite consists of microcline phenocrysts; locally, however, the number of phenocrysts diminishes, and the granite grades into almost nonporphyritic coarse-grained quartz monzonite identical with the matrix of the porphyritic granite.

Hoskin Lake Granite intrudes Marinette Quartz Diorite, and both units intrude metabasalts of the Quinnesec Formation. The contacts between metabasalt and granite or quartz diorite are sharp. The contacts between granite and quartz diorite are generally gradational, because near most of them the granite contains many quartz diorite inclusions and the quartz diorite is cut by many stringers and dikes of granite.

The age of neither the Marinette Quartz Diorite nor the Hoskin Lake Granite is known with certainty. Intrusive relations show that these rocks are younger then the Quinnesec Formation, which is considered to be of early Precambrian age, and Hoskin Lake Granite is intruded by unmetamorphosed diabase dikes of Keweenawan age (late Precambrian). Their age relation to the Animikie metasedimentary rocks (middle Precambrian) is controversial, Prinz (1959) considered them both to be pre-Animikie, whereas R. W. Bayley (written commun., 1960), who has restudied the eastern part of the Menominee district in Dickinson County, Mich., concluded that they are post-Animikie.

MASHEL FORMATION OF SOUTHWESTERN PIERCE COUNTY, WASHINGTON

By KENNETH L. WALTERS

The name Mashel Formation is given herein to a sequence of unconsolidated fluvial and lacustrine deposits of Miocene age that unconformably underlie Pleistocene deposits and overlie consolidated rocks in southwestern Pierce County, Wash.

J. E. Sceva proposed in 1955 (written commun.) that these deposits be named after the Mashel River (fig. 10), in whose valley walls they are exposed.

CHARACTER AND THICKNESS

The Mashel Formation consists of a predominantly fine-grained upper part and a coarse-grained lower part. The upper part is composed mostly of clay and sand. The clay is predominantly

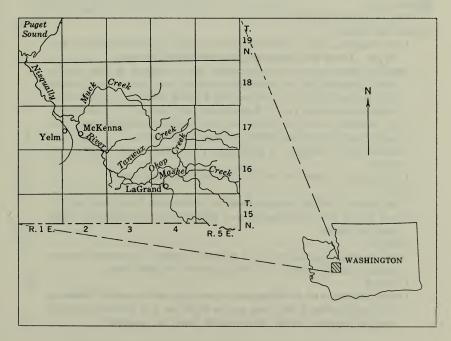


FIGURE 10.—Southwestern Pierce County area, Washington.

light colored and commonly contains plant material ranging in degree of preservation from unidentifiable fragments to whole leaves and sections of logs. The sandy phases of the upper part of the formation contain tuffaceous material, pumice, and volcanic ash.

The lower part of the formation is composed mostly of iron-stained medium- to coarse-grained poorly cemented gravel of predominantly dark volcanic rock types. Most of the granitic pebbles, where present, are badly decomposed and crumble if struck with a hammer.

The greatest thickness of the Mashel observed is about 225 feet. The total thickness may be in excess of 500 feet. Locally, the finegrained upper part may be more than 400 feet thick.

The following section, measured by the author and Grant E. Kimmel, is designated the type section. An excellent but rather inaccessible section of the formation was measured by J. E. Sceva and B. A. Liesch in the north bluff of the Mashel River about 1,350 feet east of a power transmission line, near the center of sec. 20, T. 16 N., R. 4 E.

Thickness

Vashon Drift:

Mashel Formation ·

Gravel, coarse-grained.

u	sher i of mation.	(Jeel)
	Clay, massive, rusty; weathers white to cream; contains organic ma-	
	terial. Lower few inches contains streaks of fine sand	7
	Sand, medium, angular to subangular, tan; composed of clear quartz,	
	biotite, muscovite, and miscellaneous dark rock fragments; includes	
	thin beds of silty clay that contain a small amount of organic ma-	
	terial and lenses of partly decomposed pebbles having a purple hue_	14
	Clay, massive, cream to white; has rusty streaks; contains bits of or-	
	ganic material; becomes silty and sandy near base	7
	Sand, medium, angular to subangular; contains bits of organic ma-	
	terial	2
	Clay, massive, cream to white; contains organic material	1.5
	Sand	1.9
	Clay, cream to white; contains wood and leaves near base (fossil loc. 4)_	3
	Sand, gray if wet; contains silt and clay; grades laterally into coarse	
	sand and gravel	3.8
	Lignite	0.7
	Clay, massive; appears waxy on fractures; contains scattered organic	
	material and fragments of tuff; has same thin sand beds	14
	Covered	22
	Gravel, coarse, and boulders, rusty: contains interstratified lenses of	
	sand and blocks of clay that are as much as 5 ft across; has some	
	granitic boulders, badly decomposed; base not exposed	75

Type section of Mashel Formation measured along Weyerhaeuser Road descending from Mashel Prairie to Mashel River, SE4SW4 sec. 20 and NE4NW4 sec. 29, T. 16 N., R. 4 E.

Section of Mashel Formation measured in north bluff of Mashel River near the center of sec. 20, T. 16 N., R. 4 E.

[Measured September 12, 1953 by J. E. Sceva and B. A. Liesch]

Mashel Formation:	ness (feet)
Sand, silty, stratified, light-brown	
Sand and fine gravel, stratified	
Sand, fine, cemented, light-brown	15
Conglomerate; composed of small white pumice pebbles	6
Sand, stratified, rusty-red	6
Clay and silt, gray	12
Sand, crossbedded, gray; contains many streaks of white pun	aice
pebbles	25
Sand, silty, brown	
Sand and gravel, rusty; composed predominately of pink and g	
andesite	
Sand, silt, and ash	
· ·	
Ash, hard, brittle, brown	
Tuff, hard, green	
Sand, rusty, brown	
Clay, and fine sand, silty, laminated	5
Clay, gray, and fine sand	12
Silt, sand, and clay	5
Lignite	1
Sand, silt, and clay, stratified, gray	
Gravel, basaltic, cemented, rusty; base not exposed	

mb : . l.

AGE AND CLIMATIC IMPLICATIONS-

Floras collected by the author and Grant E. Kimmel from the Mashel at the localities indicated in the following table were identified by Jack A. Wolfe of the U.S. Geological Survey:

	Locality					
Flora	1	2	3	4		
Pinus ponderosa Lawson			×			
Populus trichocarpa Torrey and Gray Salix hesperia (Knowl.) Condit						
sp., n. sp Carya simulata (Knowl.) Brown				- ×		
Pterocarya mixta (Knowl.) Brown	×	××		-		
sp., n. sp Alnus relata (Knowl.) Brown sp., n. sp				- ×		
sp., cones		××		×		
Betula lacustris MacGinitie Fagus sanctieugeniensis Hollick		×				
Quercus chrysolepis Liebmann Ulmus speciosa Newberry		×××××	×			
Zelknova oregoniana (Knowl.) Brown		Ŷ				
Mahonia reticulata (MacG.) Brown		×××				
Cinnamomum sp., n. sp Persea lanceolata (Berry) Brown	X	×	X			
Platanus dissecta Lesquereux	×		×			
Acer sp., n. sp		×		_ ×		
macrophyllum Pursh Paulownia columbiana Smiley						
Cornus sp., n. sp Fraxinus sp		\times				

Tanwax Lake quadrangle, SW¼SE¼ sec. 22, T. 17 N., R. 4 E., on east shoulder of Clear Lake highway, 478 ft north of Golden Road junction. Altitude about 675 ft.
 Eatonville quadrangle, SE¼NE¼ sec. 25, T. 16 N., R. 3 E., on southeast side of Weyerhaeuser logging road, in bluff southeast of Ohop Creek, 260 ft northeast of bend in road at creek level. Altitude about 525 ft.

Batonville quadrangle, NW¼SE¼ sec. 18 T. 16 N., R. 4 E., on west side of State highway 7, in bluff west of Ohop Valley.
 Eatonville quadrangle, NW¼NW¼ sec. 29, T. 16 N., R. 4 E., on northwest shoulder of Weyerhaeuser logging road, in bluff west of Mashel River. Altitude about 525 ft.

Wolfe (written commun., 1961) assigned the flora from locality 2, and also that from locality 1, a probable late middle Miocene age. He considered the flora from locality 3 to be no older than late Miocene and that from locality 4 to be either late middle or late Miocene.

The climate at the time of deposition of the Mashel Formation, according to Wolfe, was probably warmer than at present, and there may have been at least as much precipitation. The flora from locality 3 represents a climate somewhat cooler than that represented by the flora from locality 2.

DISTRIBUTION

The Mashel Formation is well exposed in the bluffs along the lower Mashel River, the Ohop Valley, the south valley wall of Tanwax Creek, and along the Nisqually River valley from near LaGrande to the mouth of Tanwax Creek. The formation is typically exposed in sec. 20 and 29, T. 16 N., R. 4 E., along a logging road descending from Mashel Prairie to the Mashel River. Near McKenna, in the Nisqually River valley, the Mashel disappears under younger formations.

Similar unconsolidated Miocene deposits of fluvial and lacustrine origin containing pumice and volcanic ash are reported in the vicinity of Voight Creek southeast of Orting (Mullineaux and others, 1959) and in the valley walls of the Green River in King County, Wash. (Glover, 1941, p. 138).

MODE OF DEPOSITION

The Mashel Formation probably was deposited in a piedmont environment during and after the uplift of the ancestral Cascade Range. The areal distribution of the lower part of the formation is not known, but the deposition of much of the gravel was probably limited to areas marginal to the upland, except along northwesttrending drainage lines that may have been in existence on the lowland before the uplift took place. Early in the development of the piedmont plain underlain by the Mashel, sand was deposited in the area midway between the upland and the center of the lowland; silt and clay were deposited near the center of the lowland. As the upland area was eroded and the lowland was aggraded, the area in which gravel was being deposited was reduced in size, and finegrained material was deposited closer to the upland front.

PRECAMBRIAN AND LOWER CAMBRIAN FORMATIONS IN THE LAST CHANCE RANGE AREA, INYO COUNTY, CALIFORNIA

By JOHN H. STEWART

About 8,000 feet of conformable strata of Precambrian and Early Cambrian age are exposed in the Last Chance Range area (fig. 11). The strata in the northern part of the range and nearby areas are assigned to the Deep Spring Formation of Precambrian age, the Campito Formation of Precambrian and Early Cambrian age, the Poleta, Harkless, Saline Valley, Mule Spring Formations of Early Cambrian age, and the Emigrant(?) Formation of Middle and Late Cambrian age (fig. 12). The strata exposed farther south in the range are assigned to the Wood Canyon Formation of Precambrian and Early Cambrian age, the Zabriskie Quartzite of Early Cambrian age, and the Carrara Formation of Early and Middle Cambrian age (fig. 12).

The nomenclatural differences between the northern and southern parts of the area reflect a change in the lithologic character of the

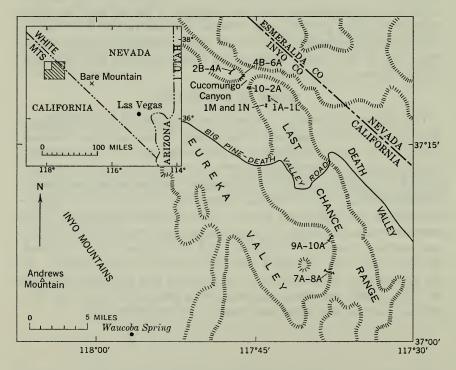


FIGURE 11.—Map showing location of measured sections in the Last Chance Range area, Inyo County, Calif. Numbers and letters refer to units in figures 13 and 14.

CHANGES IN STRATIGRAPHIC NOMENCLATURE

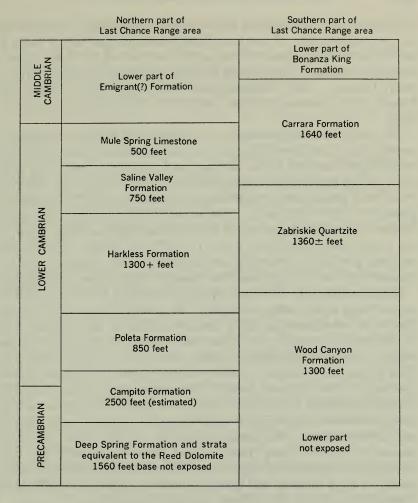


FIGURE 12.—Precambrian and Lower Cambrian formations identified in the Last Chance Range area, Inyo County, Calif.

strata. The strata in the northern part of the area contain more siltstone and limestone, and less quartzite, than comparable strata in the southern part of the area. The terms "Deep Spring" "Campito," "Poleta," "Harkless," "Saline Valley," and "Mule Spring," indicate the lithologic similarity of the strata in the northern part of the Last Chance Range area to the rocks in the stratigraphic sections to the west and north of the Last Chance Range in the Inyo and White Mountains in California (Nelson, 1962) and in Esmeralda County, Nev. (Albers and Stewart, 1962; McKee and Moiola, 1962) where these terms are used. The terms "Wood Canyon," "Zabriskie," and "Carrara" indicate the close lithologic similarity of the strata in the southern part of the Last Chance Range area to the rocks in the stratigraphic sections to the south and east of the Last Chance Range, in southern Nevada and adjoining parts of California (Nolan, 1929; Hazzard, 1937; Cornwall and Kleinhampl, 1961; Barnes and Palmer, 1961; Barnes and others, 1962) where these terms are generally used.

The author's study is based on reconnaissance geologic mapping of much of the Last Chance Range area and on measurement of detailed stratigraphic sections in the southern and northern parts of the area. Parts of the Last Chance Range area have been mapped by Whetten (1959) and McKee (1962), and their studies were helpful in the present geologic study in the area. McKee and Moiola (1962) mentioned briefly some of the exposures of Precambrian and Cambrian strata in the range directly northwest of the Last Chance Range. The present study is part of a continuing investigation of the correlation of the Precambrian and Lower Cambrian strata in the southern Great Basin.

Strata assigned to the Deep Spring Formation (fig. 13), including at the base some strata considered to be equivalent to the Reed Dolomite, are exposed in badly faulted sections in the northern part of the Last Chance Range. These are the oldest rocks exposed in the range, and the total section exposed amounts to about 1,560 feet. Unit 1A (fig. 13) consists of yellowish-brown fine- to medium-grained quartzite, minor siltstone, and, in the upper part, some limestone. Units 1B through 1L consist of gray and grayish-orange limestone, sandy limestone containing scattered fine to medium quartz grains, and minor amounts of dolomite, quartzite, and siltstone. Unit 1M consists of yellowish-gray quartzite and greenish-gray siltstone; unit 1N, of limestone and minor amounts of dolomite and siltstone; unit 1O, mostly of dark siltstone and quartzite similar to that in the Campito Formation; and unit 1P, of limestone and dolomite.

Correlation of units 1A to 1L with described units (Nelson, 1962) in the Inyo and White Mountains is uncertain. Some or all of units 1A to 1L must correlate with Nelson's (1962, p. 141; 1963) lower member of the Deep Spring Formation. Units 1A through 1L total 913 feet in the Last Chance Range as compared with only 496 feet (C. A. Nelson, written commun., 1963) for the lower member near Andrews Mountain (fig. 11) in the Inyo Mountains; thus the section exposed in the Last Chance Range probably extends down into rocks laterally equivalent to the Reed Dolomite, the formation underlying the Deep Spring Formation in the Inyo and White Mountains. This correlation is supported by the lithologic similarity of unit 1A with the Hines Tongue of the Reed Dolomite as described by Nelson (1962, p. 141). The dolomite that forms the top of the Reed Dolomite in the Inyo and White Mountains may be represented by limestone in the section in the Last Chance Range.

Units 1M and 1N are considered to correlate with Nelson's (1962, p. 141, 1963) middle member of the Deep Spring Formation, and units 1O and 1P with his upper member of that formation.

The Campito Formation consists of olive-gray, greenish-gray, darkgreenish-gray, and medium-gray very fine grained quartzite and siltstone and conformably overlies the Deep Spring Formation. The formation is extensively exposed in the northern part of the area. The thickness is estimated to be about 2,500 feet, although little significance can be placed on this figure because faulting in the area makes accurate measurements of the thickness impossible. Both the quartzitic Andrews Mountain Member and the overlying thinner and silty Montenegro Member (Nelson, 1962, p. 141) were recognized in the northern part of the area.

The Poleta Formation conformably overlies the Campito Formation and is divided into a lower member (unit 3A, fig. 14) of archeocyathidbearing limestone; a middle member (unit 3B-3E, fig. 14) of greenishgray and olive-gray siltstone, minor limestone, and, in the upper part, some quartzite; and an upper member (unit 3F, fig. 14) of limestone. Division of the formation into three members follows the practice of McKee and Moiola (1962, p. 534-535) although Nelson (1962, p. 142) originally divided the formation into only two members; the upper two members of McKee and Moiola correspond to the upper member of Nelson. The middle member contains the trilobite *Nevadella* (identified by A. P. Palmer, written commun., 1963) in the basal 181 feet of unit 3C. Rare *Scolithus* (worm borings) occur in the upper part of the middle member. The Poleta Formation is about 850 feet thick.

The Harkless Formation is not completely exposed in any one section in the Last Chance Range area. The lower 580 feet (unit 4A, fig. 14) of the formation is exposed 1 mile northwest of Cucomungo Canyon (fig. 11). There the strata which lie conformably on the Poleta Formation consist of greenish-gray and gravish-olive siltstone containing layers of quartzite in a few places. Trilobites that probably can be designated Fremontia(?) and Paedeumias(?) (identified by C. A. Nelson, written commun., 1964) occur in this section about 280 feet above the base of the Harkless Formation. Along Cucomungo Canyon, about 1,300 feet of the Harkless Formation are exposed conformably below the Saline Valley Formation. Parts of this section are highly faulted. The strata consist mostly of gravishpurple, pale-red, yellowish-gray, and greenish-gray fine- to mediumgrained quartzite containing numerous siltstone layers from onequarter inch to several feet thick. Unit 4C is a particularly conspicuous siltstone unit. Unit 4D is composed of dolomite and lime-

UNITS 1A-1N MEASURED IN SEC. 31 AND 32 (UNSURVEYED), T. 7 S., R. 39 E. UNITS 10-2A MEASURED IN SEC. 24, T. 7 S., R. 38 E.

CAMPITO FORMATION	{	Unit	Thick- ness (feet)	Description
5 C		2A		Olive-gray to medium-dark-gray siltstone
		1P	61	Yellowish-gray and light-gray dolomite in basal 35 ft; over- lain by very pale orange limestone
FEET	UPPER	10	223	Greenish-gray, olive-gray, and medium-gray siltstone, yellowish-gray, greenish-gray, and medium-gray very fine to fine-grained quartzite; silty limestone, siltstone and dolomite from 69 to 85 ft above base
4G FORMAI		1N	55	Gray limestone, very pale orange dolomite in top 6 ft, some siltstone in middle
- 300 DEEP SPRING FORMATION	MBER			
-400	MIDDLE MEMBER	1M	310	Yellowish-gray very fine to fine-grained quartzite and greenish-gray siltstone. Amount of quartzite decreases upward

FIGURE 13.—Columnar section of Deep Spring Formation and related

CHANGES IN STRATIGRAPHIC NOMENCLATURE

		ſ	1L	38	Gray limestone, very pale orange dolomite in top 8 ft
			1K	43	Grayish-yellow limy sandstone, gray sandstone and olive- gray siltstone
			1J	110	Grayish-orange limestone containing pisolites in top 5 ft; rare greenish-gray silt s tone
			11.	54	Olive-gray very fine grained quartzite, rare gray siltstone
		LOWER MEMBER	1H	128	Gray limestone, gray and yellowish-brown dolomite as 5 ft layer 46 ft above base and replacing limestone in top 45 ft; rare sandy limestone in middle
FEET	ATION		1G	35	Gray sandy limestone and limestone, yellowish-gray quart- zite, and olive-gray siltstone
-100	DEEP SPRING FORMATION		1F	155	Gray lìmestone and sandy limestone, minor siltstone, and rare quartzite
		1	1E	35	Yellowish-brown fine- to medium-grained quartzite and greenish-gray siltstone
-200	LOMITE	1D	77	Gray limestone, oolitic in top 10 ft ; rare siltstone	
- 300		ED DO	10	25	Olive-gray siltstone and yellowish-gray fine-grained quartzite
		TO THE REED DOLOMITE	1B	87	Gray limestone and sandy limestone, minor quartzite and siltstone
400		STRATA EQUIVALENT	1A	127+	Yellowish-brown fine- to medium-grained quartzite minor greenish-gray siltstone, and minor gray limestone in upper half
L500			 		

strata in northern Last Chance Range area, Inyo County, Calif.

NORTHERN PART OF THE LAST CHANCE RANGE AREA, UNITS 2B-4A MEASURED IN SEC. 15, T. 7 S., R. 38 E., UNITS 4B-6A MEASURED IN SECS. 14 AND 23, T. 7 S., R. 38 E.

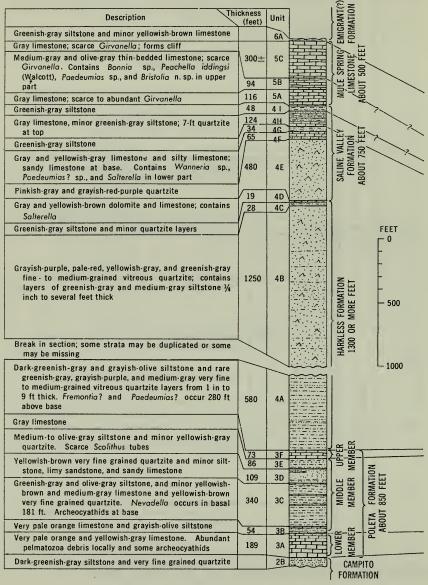


FIGURE 14.—Columnar sections and correlations of Lower Cambrian

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SOUTHERN PART OF THE LAST CHANCE RANGE AREA, UNITS 7A-8A MEASURED IN SECS. 7 AND 8 (UNSURVEYED), T. 10 S., R. 40 E., UNITS 9A-10A MEASURED IN SEC. 29, 30, 31, AND 32 (UNSURVEYED), T. 9 S., R. 40 E.

7		Unit	Thickne (feet)	SS Description
BONANZA KING ORMATIOI	E S	10A	unmeas- ured	Gray limestone
- <u>-</u>		SN	115	Gray and yellowish-brown limestone and minor silty lime- stone and limy siltstone
	臣幸	9M	176	Gray limestone; conspicuous white band at base
		9L	163	Yellowish-gray and medium-gray limestone, minor silty limestone and siltstone
TION		9K	250	Gray limestone; white band at top, locally oolitic. Common Girvanella in lower half
SMA 0 FI	물물물	L6	23	Gray limy siltstone
A PORMATION ABOUT 1640 FEET		91	44	Gray limestone
CARRA		9H	240	Greenish-gray phyllitic siltstone, scarce thin yellowish-brown or gray limestone layers
No No No		9G	200	Gray limestone. Common Girvanella throughout; forms prominent cliff
12] 9F	82	Gray limestone; mostly phyllitic siltstone in lower half
		9E	112	Gray limestone and minor phyllitic siltstone in lower part. Some Girvanella
	^``^	9D	71	Gray limestone with minor siltstone in lower part. Girvanella at top
	· · · · ·	90	47	Greenish-gray phyllitic siltstone, minor limestone
TZITE		98	32	Gray limestone and dolomitic limestone; sandy limestone at top
QUART FEET		9A	88	Yellowish-brown very fine to fine-grained quartzite, minor phyllitic siltstone
ZABRISKIE QUARTZITE 1360± FEET		8A	1360±	Medium - dark-gray to grayish-purple fine- to medium- grained rarely coarse-grained vitreous quartzite; lami- nated and minor small-scale cross-strata. Minor siltstone and conspicuous cross-strata in lower 88 ft. Cut by fault about 950 ft above base
		/7F	124	Dark-greenish-gray siltstone. Quartzite layers minor in lower part and dominant at top. Contains Paedeumias?
		/7E	28	Gray limestone
		70	194	Yellow-brown and yellow-gray very fine grained quartzite. Abundant Scolithus tubes
CANYON FORMATION		70	424	Greenish-gray to medium-gray phyllitic siltstone, and minor yellowish-brown and medium-gray very fine sandstone and yellowish-brown and medium-gray limestone. Lime- stone contains pelmatozoa debris and, in basal 140 ft, archeocyathids. A few Scolithus tubes. Nevadella? occurs in basal 10 ft
D CAN		7B	123	Very pale orange and yellowish-gray limestone. Archeo- cyathids and pelmatozoa debris
DOOW		7A	407	Yellowish-brown very fine grained quartzite, minor greenish- gray and gray phyllitic siltstone. Two thin limestone beds in upper half, the lower limestone bed contains archeocyathids and pelmatozoa debris

strata in the Last Chance Range area, Inyo County, Calif.

stone and contains *Salterella* in places; the unit is considered to correlate with a fairly persistent limestone found at the top of the Harkless Formation directly below the Saline Valley Formation in the Waucoba Spring area of the Inyo Mountains.

Unfortunately the strata of the Harkless Formation exposed 1 mile northwest of Cucomungo Canyon cannot be matched with any of the strata of that formation exposed along Cucomungo Canyon, and a small amount of strata may be missing between the two stratigraphic sections. The two sections may possibly overlap, however, and the difference in lithologic character may be due to abrupt facies changes coupled with thrust faulting; such faulting would bring unrelated facies into close juxtaposition. Thrust faults have been recognized in the vicinity of Cucomungo Canyon and one, at least, may occur between the section in Cucomungo and the section a mile to the northwest; critical outcrops are, however, covered by younger strata.

The Saline Valley Formation as exposed in Cucomungo Canyon is about 750 feet thick and consists, in the lower part, of pinkish-gray and grayish-red-purple quartzite (unit 4E) and, in the upper part (units 4F-4I), of gray limestone and greenish-gray siltstone, a few layers of sandy limestone near the base, and a layer of quartzite near the top. The quartzite in the lower part of the Saline Valley Formation is very similar to that in the Harkless Formation, and the two formations apparently can be separated only if the intervening carbonate layer (unit 4D) is between them. *Salterella* and the trilobites *Wanneria* sp. and *Paedeumias*? sp. have been collected and identified by A. R. Palmer (written commun., 1963) from the lower part of the u it 4F of the Saline Valley Formation in Cucomungo Canyon.

The Mule Spring Limestone (units 5A-5C, fig. 14) conformably overlies the Saline Valley Formation. It is about 510 feet thick and consists of gray limestone commonly containing algal structures (Girvanella). The trilobites Bonnia sp., Peachella iddingsi (Walcott), Paedeumias sp., and Bristolia n. sp. have been collected and identified by A. R. Palmer (written commun., 1963) from the upper part of unit 5B of the Mule Spring Limestone in Cucomungo Canyon.

The Mule Spring Limestone is overlain by, and is partly or entirely in thrust contact with a thick and badly faulted sequence of generally thin-bedded limestone and cherty limestone containing some dolomite and siltstone units. A siltstone unit may occur at the base of this sequence directly above the Mule Spring Limestone (fig. 14), although the outcrop showing this relationship is small and faulted. This thick sequence above the Mule Spring Limestone has not been studied critically but is probably assignable to the Emigrant Formation of Middle and Late Cambrian age, a formation overlying the Mule Spring Limestone in Esmeralda County, Nev. (Albers and Stewart, 1962, p. D27; McKee and Moiola, 1962, p. 536-537). Trilobites collected in this limestone sequence near Cucomungo Canyon by A. R. Palmer and the author have been identified as Middle Cambrian in age (A. R. Palmer, written commun., 1963).

In the southern Last Chance Range, the upper 1,300 feet of the Wood Canyon Formation (units 7A-7F, fig. 14) is exposed. The exposed section consists of very fine grained quartzite, greenish-gray and gray siltstone, and yellowish-brown and gray limestone. Archeocyathids and pelmatozoan debris occur in a thin limestone bed 177 feet below the top of unit 7A, throughout much of unit 7B, and in the basal part of unit 7C. One specimen of a poorly preserved trilobite, identified as possibly Nevadella (A. R. Palmer, written commun., 1963), was found in the basal 10 feet of unit 7C. Trilobite remains, possibly in part Paedeumias (A. R. Palmer, written commun., 1963), are locally common in unit 7F. Scolithus (worm borings) occur in the upper half of unit 7C and are abundant in unit 7D. Units 7B through 7E are correlative with the Poleta Formation of the northern Last Chance Range. This correlation is substantiated by similar thicknesses and sequences of units, as well as by similar fossil occurrences.

The Zabriskie Quartzite (unit 8A, fig. 14) in the southern Last Chance Range consists of medium-dark-gray to grayish-purple fine-to medium-grained vitreous quartzite that is coarse grained in a few places. The quartzite is mostly laminated but locally contains smallscale cross strata. The lower 88 feet contains a minor amount of siltstone, and the quartzite is conspicuously cross stratified. The measured thickness of the Zabriskie Quartzite is about 1,360 feet, although a prominent fault cuts the unit about 950 feet above its base and makes the exact thickness uncertain.

The Zabriskie Quartzite correlates clearly with the quartzite in the Harkless Formation and with that in the lower part of the Saline Valley Formation of the northern part of the Last Chance Range area. Much of the quartzite in the northern part of the area, however, contains thin layers of siltstone, which are not present in the Zabriskie Quartzite. These siltstone layers represent a change to a more silty facies within this part of the sequence to the north. This facies change is most noticeable in the lower part of the Harkless Formation. The basal 580 feet (unit 4A) of the Harkless Formation 1 mile northwest of Cucomungo Canyon is dominantly siltstone and contains only a few thin layers of quartzite. These quartzite layers are tongues of the Zabriskie Quartizite, and the siltstone, except in the basal 100-200 feet, is probably a lateral equivalent of the Zabriskie Quartzite, but of a different facies. The Carrara Formation (units 9A–9N, fig. 14) is 1,640 feet thick in a measured section in the southern part of the Last Chance Range area. It consists of gray limestone, commonly containing *Girvanella*, greenish-gray phyllitic siltstone, and yellowish-brown silty limestone. Units 9G and 9K form persistent cliffs, whose tops mark vertical lithologic changes that divide the Carrara Formation into three mappable units or members. Unit 9A is intermediate in lithologic type from the Zabriske Quartzite below to the Carrara Formation above. It is included with the Carrara Formation because a similar, and probably correlative unit, has been included in the type Carrara Formation (Cornwall and Kleinhampl, 1961) on Bare Mountain. The Carrara Formation is overlain by the Bonanza King Formation, a thick formation of Middle and Late Cambrian age composed of dolomite and limestone.

The Carrara Formation of the southern part of the Last Chance Range area contains correlatives of the upper part of the Saline Valley Formation, all the Mule Spring Limestone, and the lower part of the Emigrant(?) Formation of the northern part of the Last Chance Range area (fig. 14). Some of the quartzite in unit 9A of the Carrara Formation in the southern part of the Last Chance Range area may correlate with the quartzite at the top of unit 4H of the Saline Valley Formation in Cucomungo Canyon in the northern part of the area. Part of unit 9E and all of units 9F and 9G of the Carrara Formation probably correlate with the Mule Spring Limestone in Cucomungo Canyon. The remainder of the Carrara Formation is correlative with the lower part of the Emigrant(?) Formation, of the northern part of the area.

MIOCENE AND PLIOCENE ROCKS OF CENTRAL WYOMING By Norman M. Denson

A widespread succession of light-gray fossiliferous tuffaceous siltstone and fine-grained sandstone extends from the vicinity of the Alcova Dam westward along the north side of the Seminoe, Ferris, and Green Mountains to the southern terminus of the Wind River Range a distance of about 85 miles. Love (1961) applied the term Split Rock Formation to this succession and some associated rocks because fossiliferous Miocene rocks in eastern Wyoming are difficult to correlate with named formations and groups in Nebraska and because the lower and middle Miocene rocks of the Granite Mountains area can not be mapped continuously into the type area of the Arikaree in northwestern Nebraska.

Recent studies and regional mapping by the author from northwestern Nebraska into central Wyoming and local studies by Harshman (1964) in the Shirley Basin, Stephens (1964) at Crooks Gap, Rich (1962) at Clarkson Hill, and Zeller, Soister, and Hyden (1956) in the Gas Hills indicate that most of the rocks originally assigned to the Split Rock Formation in the Granite Mountains area are laterally equivalent and remarkably similar lithologically and chemically to the rocks in eastern Wyoming and northwestern Nebraska described and first assigned by Darton (1899) to the Arikaree Formation. Furthermore, the upper part of the Split Rock Formation at its type locality includes rocks lithologically very similar to the lower part of the Ogallala Formation and, for this reason, is here assigned to that part of the Ogallala. In other areas the lower part of the Split Rock includes the upper part of the White River Formation (Oligocene). A correlation chart of Miocene and Pliocene rocks of central Wyoming follows:

Love (1961)				This report				
Series	Age	Formation	Subdivision	Formation	Age	Series		
Pliocene	Early or middle Pliocene	Moonstone			Pliocene and late Miocene	Pliocene and Miocene		
			Upper porous	Ogallala				
			Local fauna ¹					
		Split Rock	sandstone					
	Middle Miocene		sequence		Middle and early Miocene	Miocene		
			Silty	Arikaree				
			sequence					
			Clayey sandstone sequence					
			Vertebrate for	sils ²				
	Early Miocene		Lower porous sandstone sequence	White River (upper part)	Late Oligocene	Oligocene		

¹"Split Rock local fauna " of middle Miocene age (Love, 1961, p. 19)

²Merycoides cursor Douglass of early Miocene (Gering) age (Rich, 1962, p. 506)

Rocks of widely different lithologies and ages were originally assigned to the Split Rock Formation in two measured sections at the type locality. They include middle Miocene rocks (sec. 36, T. 29 N., R. 90 W.), assigned in this report to the Arikaree, as well as Pliocene rocks (secs. 25 and 36, T. 29 N., R. 89 W.) that are here referred to as the Ogallala Formation. The lower of these two distinct stratigraphic units contains many middle Miocene vertebrate fossils ("Split Rock local fauna" of Love and others) and was assigned by Love (1961, p. 19) to the upper porous sandstone sequence of the Split Rock Formation. Rocks below the local fauna are composed predominantly of wind-blown buff and tan fine- to medium-grained poorly bedded sandstone having abundant tiny rounded grains of bluish-gray magnetite. Lateral persistence in lithology and the general absence of coarse detritus and locally derived debris from the surrounding highland are outstanding characteristics of rocks below the local fauna at the type locality of the Split Rock and elsewhere over the whole region.

Conformably overlying the middle Miocene windblown sandstone at the type locality are Pliocene rocks which Love (1961, p. 17, 18) also assigned to the Split Rock. The rocks comprising this stratigraphic unit consist mostly of thin beds of relatively pure white pumicite, pumiceous limestone, sandstone, claystone, and tuff which grade mountainward into fanlike deposits of coarse-grained sandstone, conglomerate, and gravel. The basal contact becomes an unconformity near the mountains. Most of the rocks contain a preponderance of volcanic ash. These volcanic-rich rocks have yielded many species of diatoms and spores (Love, 1961, p. 17, 20, 21) and have been traced eastward to the vicinity of the Pathfinder Reservoir where they have yielded vertebrate fossils determined by P. O. Mc-Grew to be of Pliocene age (J. D. Love, oral commun., 1964). These rocks range from a few feet to at least 300 feet in thickness and are assigned on the basis of lithologic similarity and age to the Ogallala Formation; some vertebrate paleontologists (Wood and others, 1941, p. 27) consider the oldest of several Ogallala local faunas to be of latest Miocene age, whereas others (for example, Schultz and Falkenbach, 1949, p. 80, 83) believe that all the Ogallala is Pliocene.

The Oligocene-Miocene (White River-Arikaree) contact along the eastern and northeastern margins of the Granite Mountains area was drawn by Rich (1962, pl. 7, p. 503-506) at the base of a persistent and wide-spread conglomerate that is 150-600 feet thick. Love (1961, p. 9-12) referred to it as the lower porous sandstone sequence. Areal mapping by the writer from the Granite Mountains eastward into the Shirley Basin now indicates that this widespread conglomeratic succession is a lateral equivalent of a conglomerate, 300-350 feet thick, that some previous workers (Love and others, 1955) assigned to the Miocene and Pliocene, but that Harshman (1964) correctly mapped as an upper coarse-grained member of the Oligocene White

River Formation. At several localities in the Shirley Basin, this conglomeratic sequence has yielded Oligocene (Brule) vertebrate fossils from its lower and middle parts (Harshman, oral commun. 1965; Whitmore, F. C., Jr., and Lewis, G. E., written commun., 1962, 1963). Rocks identical to the Arikaree in the Clarkson Hill area overlie the conglomerate and were mapped by Rich (1962, pl. 7) as Miocene. In the eastern and northeastern parts of the Granite Mountains area, these thick conglomerates directly overlie fossiliferous lower Oligocene (Chadron) rocks and are assigned herein for the first time to the upper part of the White River. At many places in the Shirley Basin, light-gray calcareous sandstone at the base of the Arikaree directly overlies Harshman's upper coarse-grained member of the White River Formation. In the eastern part of the Granite Mountains and in the Shirley Basin, conglomerates occur at the base of the Arikaree but are not widespread; these conglomerates are largely channellike lenticular deposits generally less than 60 feet thick. At many localities the conglomerates are interbedded with fine- to medium-grained light-gray calcareous sandstone lithologically similar to that in the lower part of the Arikaree, from which early Miocene fossils have been reported in the Granite Mountains (Rich. 1962, p. 506).

The following analyses are presented to show the striking similarity in chemical composition of the Miocene rocks on the west and east sides of the Laramie Range.

Rapid rock analyses of very fine grained tuffaceous sandstone of Miocene age

	stone in o	middle Mie entral Wyon ness 1,000 ft)		Arikaree Formation in southeast Wyoming and northwest Nebras- ka (average thickness 700 ft)		
	Range (5	analyses)	Average	Range (10 analyses)		Average
	From-	То		From-	То—	
$\begin{array}{c} {\rm SiO}_{2-} \\ {\rm Al}_2 {\rm O}_{3-} \\ {\rm K}_2 {\rm O}_{-} \\ {\rm Fe}_2 {\rm O}_{3-} \\ {\rm CaO}_{-} \\ {\rm CaO}_{-} \\ {\rm Na}_2 {\rm O}_{-} \\ {\rm MgO}_{-} \\ {\rm TiO}_{2-} \\ {\rm P}_2 {\rm O}_{5-} \\ \end{array}$	$2.0 \\ 2.2 \\ 1.6 \\ 1.3 \\ 1.3 \\ .21$	$74. \ 3 \\ 12. \ 5 \\ 4. \ 2 \\ 3. \ 1 \\ 2. \ 6 \\ 2. \ 0 \\ 2. \ 6 \\ . \ 32$	$\begin{array}{c} 69.\ 7\\ 11.\ 7\\ 3.\ 1\\ 2.\ 5\\ 2.\ 3\\ 1.\ 6\\ 1.\ 9\\ .\ 26\end{array}$	$\begin{array}{c} 66.\ 8\\ 10.\ 5\\ 2.\ 3\\ 2.\ 2\\ 1.\ 2\\ 1.\ 4\\ .\ 9\\ .\ 30\\ .\ 02\\ \end{array}$	$75.8 \\ 13.2 \\ 3.8 \\ 3.4 \\ 3.1 \\ 2.4 \\ 1.7 \\ .46 \\ .12$	$71.9 \\ 11.9 \\ 2.9 \\ 2.7 \\ 2.3 \\ 2.0 \\ 1.3 \\ .37 \\ .07$
MnO H ₂ O Sum		11. 0	7.3 100.33	. 04 2. 7	. 08 7. 3	. 06 4. 50 100. 00

[Method described by Shapiro and Brannock (1956)]

The Moonstone Formation defined by Love (1961, p. 25–35) for some lower or middle Pliocene rocks in the Granite Mountains area is a sufficiently distinctive lithologic unit to warrant a separate rockstratigraphic designation; however, the white uranium-bearing shale, green tuff, bedded chalcedony, and finely-laminated tuffaceous arenites, algal reefs, and lenticular beds of conglomerate to which Love applied the term Moonstone are present only locally in the Granite Mountains. These rocks occur principally in an area of about 50 square miles in and adjacent to T. 30 N., R. 89 W. Elsewhere in the region a succession of rocks lithologically similar to those assigned to the Moonstone is not present. The term Moonstone, therefore, has limited use as a rock-stratigraphic designation and is referred to in this report as part of the Ogallala.

In summary, the lower and middle Miocene rocks of the Granite Mountains area of central Wyoming are composed largely of windblown fine- to medium-grained tuffaceous sandstone having thin and relatively unimportant interbeds of limestone, tuff, and conglomerate. These rocks average about 1,000 feet in thickness and constitute an easily recognized lithogenetic unit. This widespread unit has been mapped discontinuously from the vicinity of Oregon Buttes along the southwest flank of the Wind River Mountains through central and southeastern Wyoming into northwestern Nebraska. Because this unit is strikingly similar lithologically to the Arikaree Formation as defined by Darton and because the rocks that constitute it are unconformably overlain and underlain at most places by rocks that can properly be assigned to the Ogallala and White River Formations, respectively, the term Arikaree is applied here in central Wyoming with the same meaning given it by Darton in 1899. Since 1899, Arikaree has been used for Miocene (unrestricted) rocks in Wyoming by the U.S. Geological Survey, although the stratigraphic and chronologic range elsewhere has been restricted. As indicated in this report, the term "Split Rock Formation" is not regionally useful or meaningful and, therefore, is abandoned.

The Moonstone, originally defined to include some lower or middle Pliocene rocks in the Granite Mountains area, is referred to in this report as part of the Ogallala because of its limited areal extent.

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